

Halogens Enhance Haze Pollution in China

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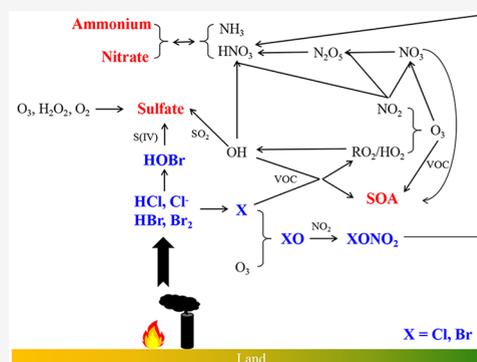
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ABSTRACT: Severe and persistent haze events in northern China, characterized by high loading of fine aerosol especially of secondary origin, negatively impact human health and the welfare of ecosystems. However, current knowledge cannot fully explain the formation of this haze pollution. Despite field observations of elevated levels of reactive halogen species (e.g., BrCl, ClNO₂, Cl₂, HBr) at several sites in China, the influence of halogens (particularly bromine) on haze pollution is largely unknown. Here, for the first time, we compile an emission inventory of anthropogenic bromine and quantify the collective impact of halogens on haze pollution in northern China. We utilize a regional model (WRF-Chem), revised to incorporate updated halogen chemistry and anthropogenic chlorine and bromine emissions and validated by measurements of atmospheric pollutants and halogens, to show that halogens enhance the loading of fine aerosol in northern China (on average by 21%) and especially its secondary components (~130% for secondary organic aerosol and ~20% for sulfate, nitrate, and ammonium aerosols). Such a significant increase is attributed to the enhancement of atmospheric oxidants (OH, HO₂, O₃, NO₃, Cl, and Br) by halogen chemistry, with a significant contribution from previously unconsidered bromine. These results show that higher recognition of the impact of anthropogenic halogens shall be given in haze pollution research and air quality regulation.

KEYWORDS: Reactive halogen species, Secondary aerosol, Haze pollution, Anthropogenic bromine emission, WRF-Chem



1. INTRODUCTION

Haze pollution has been threatening the health of millions of people in China for the past two decades. Haze pollution is characterized by an extremely high loading of fine aerosol (also known as PM_{2.5}, particulate matter with aerodynamic diameter $\leq 2.5 \mu\text{m}$), and especially that of secondary aerosol (formed from gases in the atmosphere as opposed to the primary aerosols that are particles directly emitted from sources).¹ Secondary aerosols are mainly secondary organic aerosol (SOA), sulfate, nitrate, and ammonium. SOA comprises many oxidation products of various volatile organic compounds (VOCs) by oxidants, that is, OH, O₃, and NO₃.² Sulfate is formed from the gaseous oxidation of SO₂ by OH, aqueous reactions by H₂O₂, O₃, and O₂ catalyzed by transition metal ions (TMIs) in cloud droplets, and heterogeneous reactions on aerosol.^{3–5} Nitrate is produced from NO_x chemical processes involving OH and O₃.⁶ Ammonium aerosol (in the form of ammonium sulfate and ammonium nitrate) is determined by the availability of sulfuric acid, nitric acid (HNO₃), ammonia (NH₃), and aerosol pH. Details of traditional formation channels of secondary aerosols are summarized in [Supporting Information \(SI\) Text S1](#). The current aerosol mechanisms are not able to fully explain the observed secondary aerosols in haze pollution in China calling for a better understanding of their formation processes.⁷

In addition to the well-known role of ocean-emitted halogens in photochemistry in the polar and marine troposphere,^{8,9} anthropogenic halogens are increasingly recognized to play a role in continental air pollution,^{10–17} by perturbing the levels of oxidants and atmospheric pollutants, including HO_x (OH and HO₂), NO_x (NO and NO₂), VOCs, O₃, sulfur species, etc. Briefly, halogens enhance the production of oxidants (OH, O₃, and NO₃) by igniting photochemistry (R1–R7). Halogens also directly destroy O₃ via reactions of halogen atoms with O₃ (R8) and indirectly destroy O₃ via reducing the level of NO₂ (R9–R10, R5). The presence of abundant VOCs and NO_x species in the polluted region facilitates the reactions R1–R7 leading to an enhancement effect of halogens on oxidants; meanwhile, in the clean environment, reactions R8–R9 directly destroy O₃ (the main source of OH radicals) and dominate the effect of halogens on oxidants, decreasing oxidation capacity.

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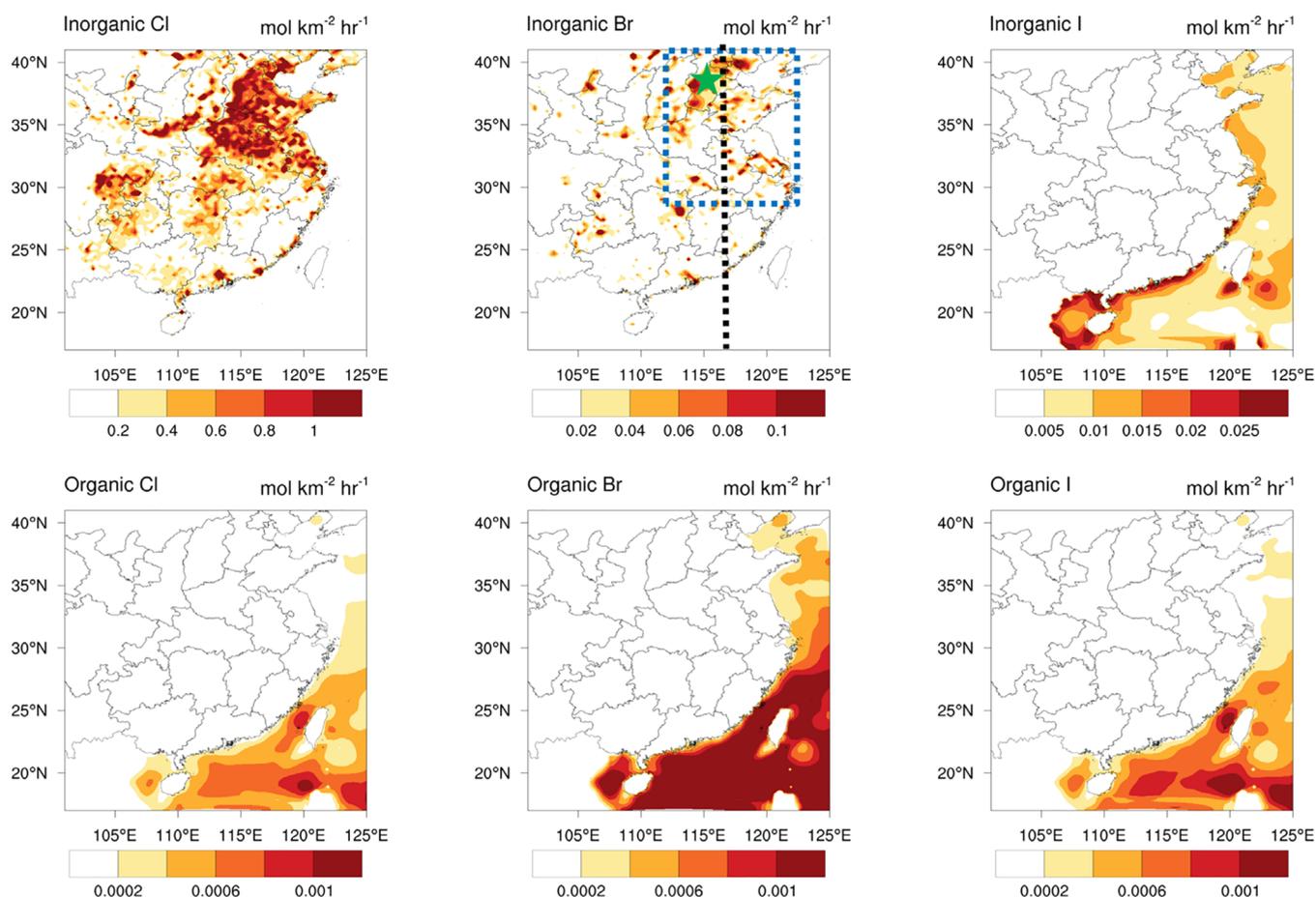
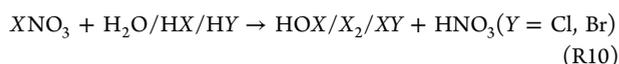
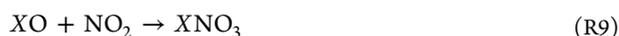
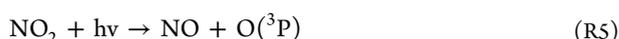
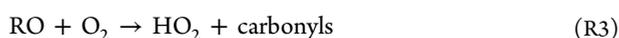
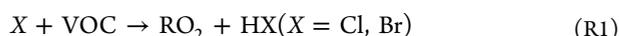


Figure 1. Average emissions of Cl, Br, and I species ($\text{mol km}^{-2} \text{hr}^{-1}$). The blue rectangle represents the region of interest, northern China. The green star stands for the location of the Wangdu campaign (used in model evaluation in SI Text S4). The black line (longitude = 116.3°) is the location of the vertical cross-section shown in Figure 4.



In the past few years, observational studies have reported elevated levels of halogen species in China. The majority of these reports focused on ClNO_2 ^{18–22} and a few on Cl_2 ^{23,24} and bromine species (BrCl , HOBr , Br_2 , and HBr).^{25,26} A few modeling studies evaluated the impacts of chlorine chemistry on O_3 and $\text{PM}_{2.5}$.^{11,14} However, the overall impact of anthropogenic halogens (particularly bromine) on aerosols and haze pollution in China remains poorly understood.

In this study, for the first time, we quantify the collective role of halogens in haze pollution with a regional model, Weather Research and Forecasting model coupled with Chemistry (WRF-Chem), updated to include comprehensive reactive halogen chemistry and natural halogen sources,²⁷ as well as anthropogenic chlorine²⁸ and bromine emissions. The bromine emission presented in this work is the first compilation of anthropogenic inorganic bromine sources in the world. The bromine emission compilation method, the WRF-Chem model, and WRF-Chem simulations are summarized in Section 2. The simulation results of the halogen impacts on secondary aerosols are discussed in Section 3.

2. MATERIAL AND METHOD

2.1. Anthropogenic Bromine Emission Inventory. A bottom-up emission inventory of reactive bromine species from coal combustion activities, including power plants, industrial processes (e.g., cement, iron and steel, brick, lime production), industrial boiler, and residential burning, in China in the year 2017 is proposed in this study. An emission factor method was applied to estimate total bromine emissions, based on the following equations:

$$E_{\text{Br}} = \sum_{i,j,k} A_{i,j,k} \text{EF}_{i,j,k} \quad (1)$$

$$EF_{i,j,k} = C R_{j,k} \sum_l (1 - f_{(\text{SO}_2)_{i,j,k,l}} \eta_{(\text{SO}_2)_{i,j,k,l}}) \sum_m (1 - f_{(\text{PM})_{i,j,k,m}} \eta_{(\text{PM})_{i,j,k,m}}) \quad (2)$$

where A is coal consumption, which was obtained from Chinese official energy statistics²⁹ or calculated as the products of industrial outputs and the corresponding coal intensities, as described in Fu et al.²⁸ Residential coal consumptions have been revised based on previous surveys,³⁰ due to high uncertainties in statistic data for this sector. EF is the emission factor. C is Br content in coal. Peng et al.³¹ reported values ranging from 0.12 to 69.66 $\mu\text{g/g}$ in 305 coal samples from 27 provinces in China. The highest value, 69.66 $\mu\text{g/g}$, was used for raw coal in this study, to account for other potential unknown Br sources which are not included in this study. Even though the current estimation represent an upper limit for Br emission from coal combustion, the total bromine emission over China is not overestimated, as indicated by the validation results of reactive bromine species in our study (Section 2.3). Br contents in cleaned coal and briquette coal were calculated as described in Fu et al.²⁸ The Br release rates (R) for different combustion technology and processes were derived from previous measurements and studies, as listed in SI Table S1. $f_{(\text{SO}_2)}$ and $f_{(\text{PM})}$ are application rates of conventional SO_2 and PM emission control technologies, obtained from the emission database established in previous studies.^{32,33} The Br removal efficiencies ($\eta_{(\text{SO}_2)}$ and $\eta_{(\text{PM})}$) were set as 13.4% for electrostatic precipitator and fabric filter, 50% for wet scrubber, and 92.5% for flue gas desulfurization, based on the previous measurements for different control devices.^{34–40} i, j, k, l, m represent the province, sector, technology type, SO_2 emission control technology and PM emission control technology, respectively. HBr and Br_2 were reported to be the two dominant emitted Br species,^{37,41} whose proportions were set as 70% and 30%, respectively, in this study. The annual emissions were distributed into each month based on the factors in Wang et al.⁴² for residential burning and equally for other sectors.

The overall emission intensity of halogen species during the simulation period is shown in Figure 1, in which anthropogenic inorganic chlorine ($>0.2 \text{ mol km}^{-2} \text{ hr}^{-1}$ in polluted areas) and bromine species ($>0.02 \text{ mol km}^{-2} \text{ hr}^{-1}$) dominate the halogen emission over China. The bromine emission inventory compiled in the current work is the first anthropogenic inorganic bromine source and there is no relevant emission data set to compare. We evaluate this newly proposed emission data set by comparing the WRF-Chem simulated bromine species, BrCl, HOBr, and Br_2 , with the respective observed values, as shown in SI Text S4.

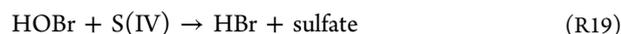
2.2. WRF-Chem Model. WRF-Chem is a regional air quality model commonly used to investigate air pollution and its formation mechanism.⁴³ The SI spreadsheet file shows the chemical mechanism (photolysis, gas-phase reactions, and multiphase reactions) used in the present study.^{27,44–48} As to the aerosol phase mechanism, there are three main processes, including nucleation, coagulation, and gas-particle partitioning. The details of the chemical processes in the standard WRF-Chem model that are relevant to the secondary aerosol formation can be found in SI Text S1. Here we describe the new reactions and chemical processes that we added in WRF-Chem in the present study.

a. SOA. We followed the volatility basis set (VBS) framework^{49–51} and utilized the Cl-initiated SOA yield

data reported in previous chamber studies^{52–57} to propose a new set of Cl-initiated SOA yield parametrizations for their use in chemical transport models (SI Text S2; Figure S1; Table S2). Two Br-initiated VOCs oxidations (R17–R18) are also considered to result in SOA formation using the same yield of Cl-initiated reactions for the lack of Br-initiated SOA yield data. In the following reactions, BIGALK, ISOP, APIN, BPIN, and LIMON represent alkanes with more than 3 carbon atoms, isoprene, α -pinene, β -pinene, and limonene, respectively.



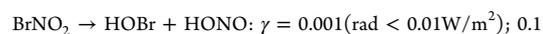
b. Sulfate. The formation of sulfate aerosol through the uptake of HOBr on the aqueous phase S(IV)-containing aerosol.^{58–61} We followed Chen et al.⁶⁰ and the references therein to adopt a rate of $5.0 \times 10^9 \text{ M}^{-1} \text{ s}^{-1}$ for $\text{HOBr} + \text{SO}_3^{2-}$ reaction, and we used the recently proposed rate of $2.6 \times 10^7 \text{ M}^{-1} \text{ s}^{-1}$ by Liu and Abbatt⁶¹ for $\text{HOBr} + \text{HSO}_3^-$ reaction.



c. We added a few heterogeneous reactions of chlorine and bromine species on the aerosol surface. According to Abbatt et al.,⁶² the solubility and reactivity of BrNO_2 are substantially higher than that of ClNO_2 , which potentially enables BrNO_2 to be taken up on the aerosol surface. In a polluted environment during winter as in northern China ($\sim 10 \text{ ppbv O}_3$ and $\sim 30 \text{ ppbv NO}_2$), BrNO_2 is the dominant form of gaseous bromine if its uptake reaction is not activated. Two multiphase channels of BrCl formation from HOCl and BrNO_3 are also added.



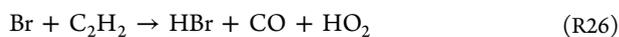
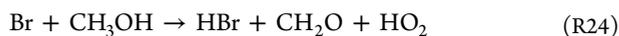
d. We also introduced a photoenhanced uptake coefficient for the heterogeneous bromine reactions to mimic the good correlation between the photolysis rate and the level of bromine species (see SI Text S4).



$$(0.01\text{W}/\text{m}^2 < \text{rad} < 400\text{W}/\text{m}^2); 0.1 * \text{rad}/400(\text{rad} > 400\text{W}/\text{m}^2)$$

e. Several Br atom initiated VOC oxidation channels. MVK and MACR represent methyl vinyl ketone and methacrolein, respectively. Note that the reactions of Br atom with ISOP and LIMON are not included in the current study and should be considered in future works

which will further enhance the halogen effects on VOC oxidation and SOA formation.



Here we acknowledge that a few previously reported potential pathways of secondary aerosols are not included in our chemical scheme. For instance, Cheng et al.⁴ proposed an aqueous pathway of sulfate formation from NO₂ oxidation but we expect such a pathway would not have a large effect in northern China during our simulated period due to the low RH (SI Figure S2). Recent works^{5,63} also suggested that the pH level in northern China might be too low for the NO₂ oxidation pathway to occur. The research on SOA formation from various chemical pathways (e.g., oxidation of primary organic aerosol) has received attention in the past few years^{64,65} but these new channels are not considered in the present work. The omission of these very recent developments of SOA could be the cause of the underestimated SOA in our study (see Section 2.3). However, we note that current chemical transport models are subject to large uncertainty in secondary aerosol simulation particularly under polluted conditions. The purposes of this study are (1) to implement the most up-to-date halogen sources and chemistry, (2) to quantify the relative impact of halogens in the formation of fine aerosol, and (3) to motivate further research oriented to reproduce the observed high levels of secondary aerosols during haze pollution in China.

2.3. WRF-Chem Simulations. The domain of the WRF-Chem simulation covers the major city clusters in China with a grid size of 27 km. The simulation period is the same as the Wangdu campaign, during which comprehensive reactive halogen species were measured,²⁵ i.e. December 9 to 31, 2017, with an extra spin-up period of 20 days (November 20 to December 8, 2017). The physical and chemical parametrizations adopted in this study are listed in SI Table S3.

Several data sets were used to initiate and drive the WRF-Chem model. Data set ds083.2 provided by Nation Centres for Environmental Prediction (NCEP) was used as the meteorological input. Data sets ds351.0 and ds461.0 were used for the data assimilation of meteorological simulation in the WRF model. The output from the CAM-Chem model, a global chemistry and climate model, was used as chemical initial and boundary conditions following Li et al.¹⁵ As to the anthropogenic emissions, a widely used emission inventory, MEIC (www.meicmodel.org), is adopted for the routine air pollutants (NO_x, SO₂, CO, VOC, NH₃, PM_{2.5}, PM₁₀). We also applied an open fire emission inventory, the Fire Inventory from NCAR (FINN).⁶⁶ Natural halogen emissions were estimated online using the mechanism described in Badia et al.²⁷ Anthropogenic chlorine emissions from the burning activities of coal, biomass, and municipal solid waste compiled by Fu et al.²⁸ are adopted. The anthropogenic bromine

emission inventory developed for this work (Section 2.1) is also employed.

We conducted two major simulations, BASE and HAL, whose results are discussed in detail. BASE case includes no halogen sources nor chemistry and only the reactions in the standard WRF-Chem configuration.^{44–47} The HAL case includes a full set of comprehensive halogen sources and chemistry (Archer-Nicholls;⁴⁸ Badia et al.;²⁷ this work) described in Section 2.2 and SI Text S1. The difference in atmospheric compositions between BASE and HAL is the impact of the overall halogen sources and chemistry. We also ran five additional sensitivity cases, namely, no_Cl_emi, no_Br_emi, no_BrNO₂_uptake, fixed_gamma, no_ClBr_SOA, to identify the key factors determining the overall halogen effect on secondary aerosol production. The difference between these scenarios is shown in SI Table S4 and Text S3.

We use the observational data reported in Peng et al.²⁵ to evaluate the model performance. The average observed bromine species (BrCl+HOBr+2 × Br₂) during the campaign is 105 pptv; our WRF-Chem model generally simulated the magnitude and temporal variation with an average of 71 pptv (SI Figure S8; Text S4). The simulated ClNO₂ reproduces well the level and temporal trend of the observation, although with a mean underestimation of 25%. Our modeling results present some discrepancies on the hourly concentration profile due to the omission of local sources, which are a common issue of regional modeling studies. However, the general underestimation in chlorine and bromine abundances in WRF-Chem suggests that the halogen effects reported in this study should be considered as a lower limit in this region. Note that our model setup provides a good representation of the gas phase chemistry (SI Figure S9) because the simulated gaseous pollutants (NO₂, SO₂, and O₃), total reactive nitrogen (NO_y), and oxidant precursors (HONO and H₂O₂) are very close to the observations, although the simulated NH₃ is moderately lower than the measurement. As to the fine aerosol and its composition (SI Figure S10), the simulated PM_{2.5}, sulfate, and black carbon are in line with the observation, and the modeled SOA is underestimated, and the simulated nitrate and ammonium aerosols are significantly overestimated. The underestimation of SOA could be due to the omission of local sources and other chemical formation channels which are not included in this study. The good simulation of NO₂ and NO_y, the underestimation of NH₃, and the overestimation of nitrate and ammonium suggest that the gas-particle partitioning between ammonium nitrate aerosol and HNO₃ and NH₃ might be the cause for the overestimated nitrate and ammonium aerosol. We explore this issue in more detail in SI Text S4.

Overall, this updated WRF-Chem model reasonably simulates the level and temporal variation of the reactive halogen species, oxidants, and air pollutants. The overestimation of nitrate and ammonium and underestimation of SOA is not related to halogens and will not affect the aim of the present work, that is, the halogen impacts on haze pollution. Therefore, the current setup is suitable to quantify the relative impact of reactive halogen chemistry on secondary aerosol production and haze pollution in China.

3. RESULTS AND DISCUSSION

3.1. Simulated Halogens in China. The simulated spatial (horizontal and vertical) distribution and partitioning of halogens in the HAL scenario are depicted in Figure 2. The

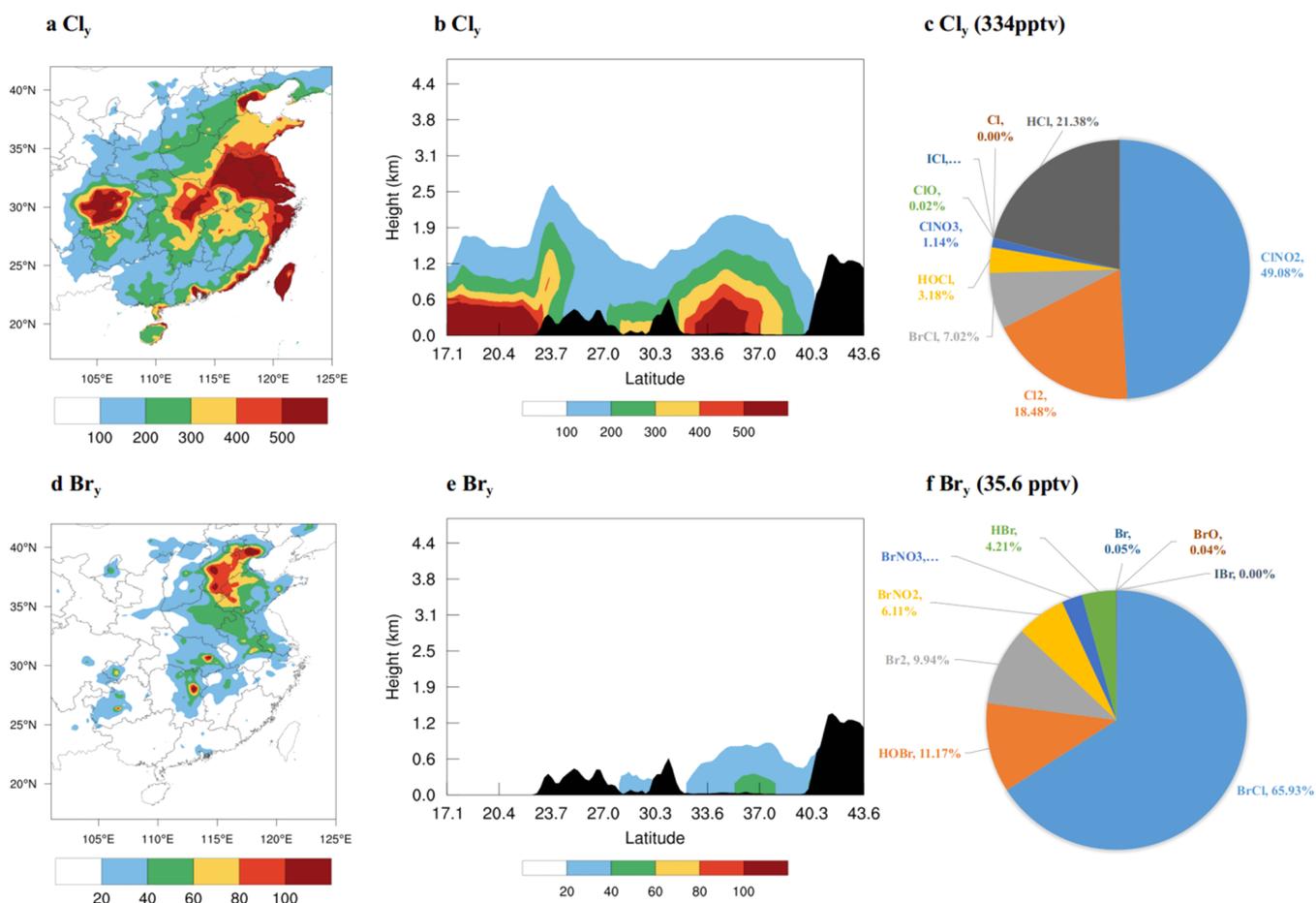


Figure 2. Average simulated Cl_y (pptv; a, surface; b, cross-section) and Br_y (pptv; d; e) in the HAL scenario. Partitioning of Cl_y (c) and Br_y (f) in the HAL scenario in northern China and the mixing ratios of Cl_y and Br_y shown in the brackets are the average values in northern China (location shown in Figure 1). The black shaded area in insets b and e is the mountainous region.

modeled Cl_y (total gaseous inorganic chlorine = $\text{BrCl} + \text{ClNO}_2 + 2 \cdot \text{Cl}_2 + \text{HOCl} + \text{ClNO}_3 + \text{ICl} + \text{ClO} + \text{Cl} + \text{HCl}$) is elevated (>300 pptv) in northern China, along the coast, and in southwestern China; whereas the estimated Br_y (total gaseous inorganic bromine = $\text{BrCl} + \text{BrNO}_2 + 2 \cdot \text{Br}_2 + \text{HOBr} + \text{BrNO}_3 + \text{IBr} + \text{BrO} + \text{Br} + \text{HBr}$) shows large mixing ratios (>60 pptv) in northern and central China and much lower in other regions. The average partitioning of Cl_y in northern China shows that ClNO_2 is the dominant species ($\sim 50\%$), whereas HCl , Cl_2 , BrCl , and HOCl contribute about 50%. A large scale observation campaign of reactive halogen species was conducted at Wangdu site (location shown in Figure 1) in northern China,²⁵ reporting the first set of inorganic bromine species in China. During the Wangdu campaign, BrCl is the predominant chlorine species,²⁵ and our simulation (in the same period) also predicted higher BrCl levels than other chlorine species on most days (SI Text S4). As to Br_y , BrCl is also the most abundant species which contributes $\sim 60\%$, whereas HOBr and Br_2 take up $\sim 20\%$. Our simulated halogen species are in line with the available observational reports of reactive bromine species in China by Peng et al.²⁵ and Fan et al.²⁶ and also within the ranges in previous studies of chlorine species in China.^{19,23,24}

3.2. Halogen Impact on Oxidants. The spatial distribution of oxidants and their changes due to halogens are shown in SI Figure S3. The addition of halogens induces large increases of OH ($+54\%$ or $+0.01$ pptv) and HO_2 ($+160\%$

or $+0.56$ pptv) in northern China (R1–R4). Similar increases ($+26\%$ to $+73\%$ in OH , HO_2 , and RO_2 radicals) due to halogens were reported at Wangdu site in northern China during the same period using a box model constrained with the observed halogen species.²⁵ Note that the enhancement of HO_x by halogens helps to close the gap between observed and the under-predicted HO_x radicals by many models.⁶⁷ O_3 is enhanced by halogens in northern China ($+27\%$ or $+5.4$ ppbv) suggesting that the O_3 enhancing pathways (R1–R6) exceed those that destroy O_3 (R8–R10, R5). The large increase of O_3 due to the reactive halogen chemistry presented here, which is not fully considered in models, could partially explain the “fast photochemistry” observed in winter in northern China.⁶⁸ NO_3 is also significantly increased due to halogens by 23% or 0.41 pptv in northern China (R7). The enhancement in these oxidants is mostly caused by the Cl and Br-initiated VOCs oxidation (R1) leading to additional production of OH and HO_2 (R2–R4) and subsequently O_3 (R5–R6) and NO_3 (R7). The presence of halogen species transforms a fraction of NO_x into halogen nitrates (R9) which also contributes to the increase in O_3 in this polluted region due to the nonlinear O_3 chemistry. A previous modeling study¹⁵ provided a more detailed analysis of the halogen impacts on oxidants in China.

3.3. Halogen Impact on Secondary Aerosol. Figure 3 shows the simulated halogen impacts (in percentage) on fine aerosol and its secondary components (SOA, sulfate, nitrate, and ammonium) over China. The simulated levels of aerosols

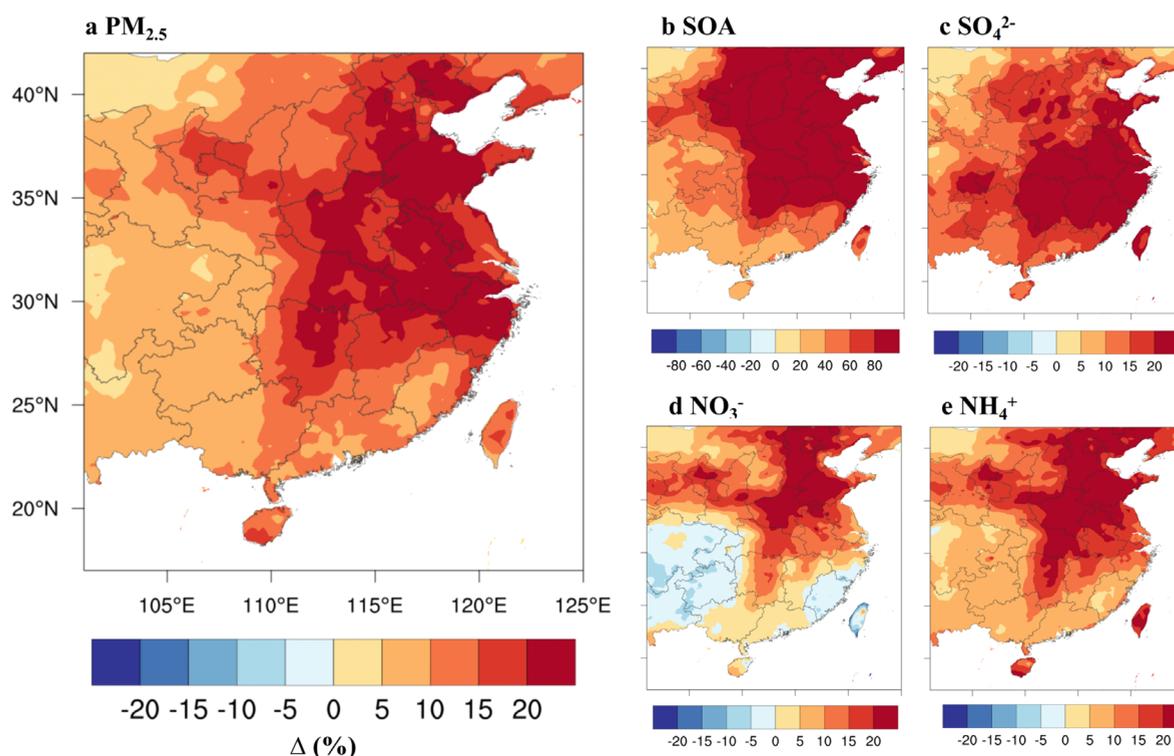


Figure 3. Average simulated impact (in percentage) of reactive halogen chemistry on (a) fine aerosol and its major components (b SOA; c sulfate; d nitrate; e ammonium) at the surface in China from December 9 to December 31, 2017. Note the difference in the color scales of SOA changes. Results for the BASE and HAL simulations and their difference (in $\mu\text{g m}^{-3}$) are shown in SI Figure S4. The change (%) in the total fine aerosol is smaller than those (%) in the secondary aerosol because some parts of the fine aerosol are primary aerosol which remains constant in the BASE and HAL cases.

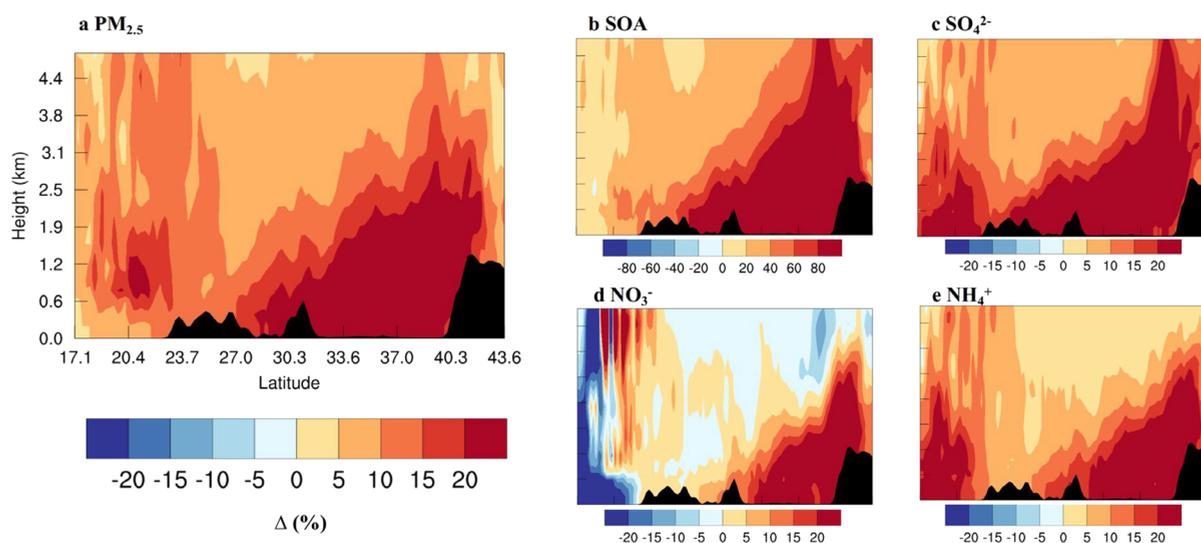


Figure 4. Average simulated impact (%) of reactive halogen chemistry on (a) fine aerosol and its major components (b SOA; c sulfate; d nitrate; e ammonium) in the south-north vertical cross-section (see location in Figure 1). The black shaded area is the mountainous region. Note the difference in the color scale of SOA changes. Results for the BASE and HAL simulations and their difference (in $\mu\text{g m}^{-3}$) are shown in SI Figure S5.

in BASE and HAL cases and their difference (in $\mu\text{g m}^{-3}$) are shown in SI Figure S4. The average increase due to halogen chemistry in northern China (the area shown as the blue rectangle in Figure 1) is 21% ($21.7 \mu\text{g m}^{-3}$) for the total fine aerosol, including 136% ($8.7 \mu\text{g m}^{-3}$) for SOA, 21% ($2.2 \mu\text{g m}^{-3}$) for sulfate, 19% ($6.4 \mu\text{g m}^{-3}$) for nitrate, and 23% ($3.2 \mu\text{g m}^{-3}$) for ammonium. Of particular interest is the modeled SOA. Although our simulated SOA in the HAL case is still

lower than the observation (see SI Text S4), the addition of halogens brings the simulated SOA concentration closer to the observations. Indeed, the halogen-mediated SOA formation helps with the current model under predictions of urban SOA.⁶⁴

We find that halogens influence the formation of fine aerosols throughout the planetary boundary layer (PBL) and extending well into the free troposphere (see Figure 4 for

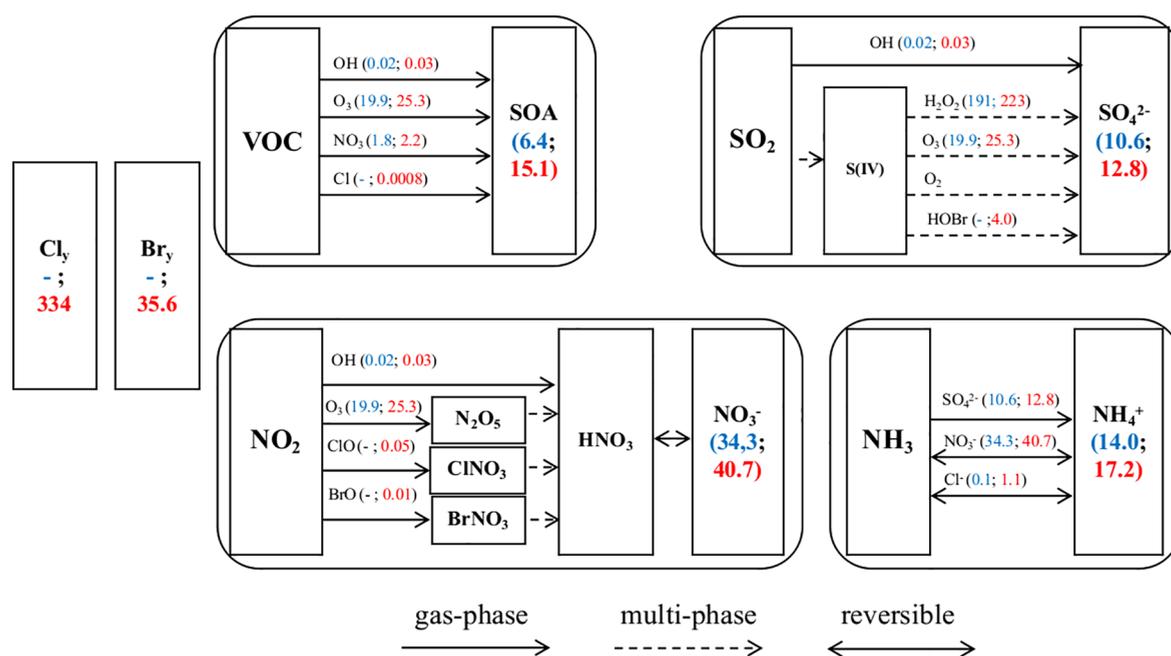


Figure 5. Formation pathways of secondary aerosols (SOA, sulfate, nitrate, and ammonium). The mixing ratios of gaseous species are the average values in northern China (location shown in Figure 1) in pptv except for O_3 which is in ppbv; the concentrations of particulate species are the average values in northern China in $\mu\text{g m}^{-3}$. Different colors represent results from the BASE (in blue) and HAL (in red) cases, e.g., OH (0.02; 0.03) represents that the OH mixing ratio in the BASE case is 0.02 pptv and that in HAL case is 0.03 pptv. Note that the level of O_2 is not affected by halogens. Note that the HOBr+S(IV) reaction takes place in aerosol water.

percentage changes and SI Figure S5 for absolute changes and the levels of aerosols in the BASE and HAL cases). Thereby, the halogen effect on secondary aerosol formation is not restricted to the emission region and boundary layer because chemical recycling processes (both gaseous and multiphase) occurring during the transport extend the area and altitude of influence. While halogen chemistry induces the largest absolute changes near the surface in northern China (SI Figure S5), the higher percentage difference is observed in the upper parts of the PBL (Figure 4). This is because northern China is “flooded” of air pollutants (e.g., NO_x) which suppress the surface levels of O_3 , NO_3 , N_2O_5 , and hence X ($=\text{Cl}, \text{Br}$) and XO radicals; whereas in the upper PBL and the free troposphere, the lower NO_x and increased O_3 allows for the more efficient recycling of gas-phase halogens, thereby further enhancing secondary aerosol production. In the cleaner environment (open ocean and free troposphere), halogens reduce O_3 and OH, therefore, decreasing nitrate aerosol production; halogens also redirect some of the N_2O_5 heterogeneous reactions to form $\text{ClONO}_2/\text{BrONO}_2$ ($\text{N}_2\text{O}_5 + \text{Cl}^-(\text{aq}) \rightarrow \text{NO}_3^-(\text{aq}) + \text{ClONO}_2$; RS11 in the SI Text S1) instead of HNO_3 when no halogens are present ($\text{N}_2\text{O}_5 + \text{H}_2\text{O}(\text{aq}) \rightarrow 2\text{HNO}_3(\text{aq})$; RS7), and hence reduce the production of nitrate aerosol within the HAL case. Such a decrease of nitrate aerosol due to chlorine reactions with N_2O_5 has been reported recently.⁶⁹

Halogens also modify the diurnal variation of the aerosols and their fractions in $\text{PM}_{2.5}$ at the surface in northern China (SI Figure S6). In particular, the fraction of SOA in the fine aerosol increases from 7.6% in the BASE case to 14.3% in the HAL case. The relative contributions of primary aerosols to the total fine aerosol, on the other hand, are decreased (from 14.7% to 11.7% for primary OA and 7.3% to 5.8% for BC). The addition of halogen chemistry also substantially increases the fraction of chloride in the fine aerosol (from 0.1% to 1.1%).

The increased chloride has recently been reported to sustain aerosol growth by enhancing ammonium aerosol and water uptake in a polluted area in India.¹⁶

3.4. Chemical Pathways for Halogen Impact on Secondary Aerosol Formation. Halogens impact the concentrations of the conventional oxidants (OH, O_3 , and NO_3) and thus the rates of the reactions of these oxidants with gas-phase pollutants (e.g., VOC, SO_2 , NO_x) ultimately altering the formation of the secondary aerosols. In addition to this dominant indirect effect, halogens also participate in the direct production processes of all secondary aerosols. Figure 5 summarizes the main chemical channels to form secondary aerosol together with the average levels of oxidants and secondary aerosols in the BASE and HAL scenarios in northern China.

Halogens enhance the production of SOA by increasing the mixing ratios of OH, O_3 , and NO_3 , and adding two new radicals (Cl and Br) (Figure 5). Sulfate aerosol is increased through enhancing OH, H_2O_2 (product of HO_x self-productions), and O_3 , and the introduction of a new sulfate formation channel initiated by HOBr (Figure 5). The nitrate aerosol concentrations are enhanced both directly through halogen oxides (R9–R10) and indirectly via their halogen-mediated enhancement of the levels of OH and O_3 (R1–R6). The increases of sulfate and nitrate and the addition of chloride, with the presence of NH_3 , induce a noticeable increase of ammonium via gas-particle partitioning (e.g., $\text{NH}_3 + \text{HNO}_3 \rightarrow \text{NH}_4\text{NO}_3$ aerosol). Liu et al.⁷⁰ suggested that $\text{PM}_{2.5}$ in China could be reduced by 11%–17% provided that a 50% reduction of NH_3 and a 15% reduction of SO_2 and NO_x are achieved. Here we note that by omitting reactive halogens, one might leave out an equivalent fraction of the aerosols formation rates in air quality model forecasting (~20% of ammonium and ~20% of $\text{PM}_{2.5}$).

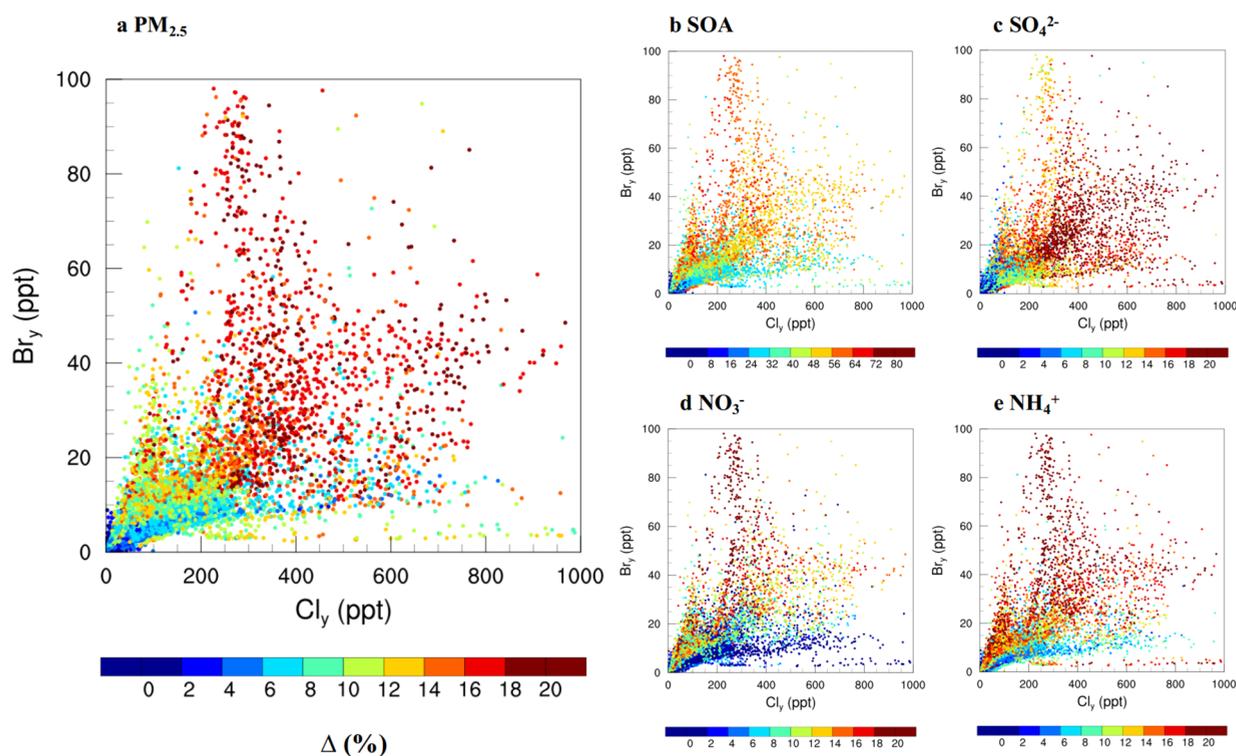


Figure 6. Relationship between the aerosols changes (a, $\text{PM}_{2.5}$; b, SOA; c, sulfate; d, nitrate; and e, ammonium) and the level of chlorine and bromine species at the surface in China. Note that the negative changes in nitrate in southeast and southwest China is omitted in this figure. The scale of SOA is different from other panels.

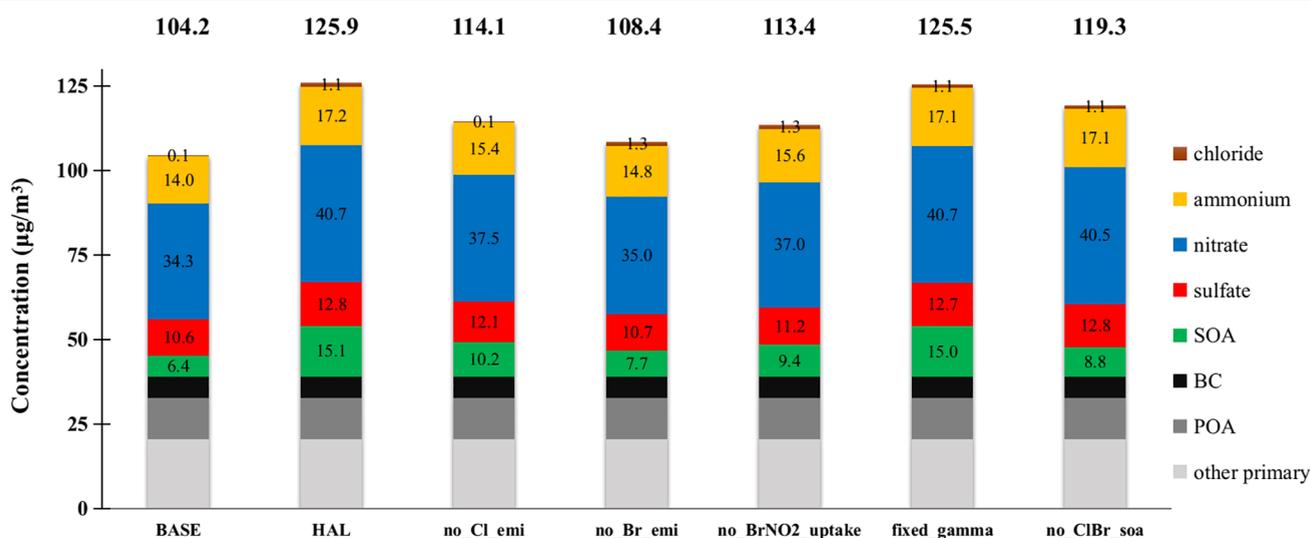


Figure 7. Average simulated $\text{PM}_{2.5}$ and its components in northern China in all simulations. The number on top of each column represents the averaged $\text{PM}_{2.5}$ concentration ($\mu\text{g}/\text{cm}^{-3}$) from each simulation.

Figure 6 further shows the relationship of aerosol changes (in all locations within the simulation domain) to the corresponding levels of Cl_y and Br_y . Our results show that larger Cl_y levels at a given Br_y level lead to larger changes of aerosols although this sensitivity is reduced for Cl_y levels over 300 pptv; larger Br_y mixing ratios at a given Cl_y level result in larger enhancements on secondary aerosol, though the changes in the aerosol are insensitive to Br_y larger than 50 pptv. An exception is sulfate, whose largest enhancement does not correlate with highest Br_y , because the multiphase reaction of $\text{HOBr} + \text{S(IV)} \rightarrow \text{HBr} + \text{Sulfate}$ requires humid conditions

(Section 2.3; R19) but a large fraction of northern China (the region with the highest bromine) has a low RH during the simulation period (SI Figure S2). An empirical connection of Cl_y to fine aerosol changes could be drawn from our simulation, that is, > 300 pptv Cl_y leads to $\sim 20\%$ enhancement in $\text{PM}_{2.5}$ and similarly, > 50 pptv Br_y leads to $\sim 20\%$ enhancement in fine aerosol implying that per unit of mixing ratio, bromine leads to a larger impact on aerosols compared to that of chlorine. Such phenomena will be discussed in more detail in the next section using sensitivity simulations.

3.5. Sensitivity of Halogen-Driven Aerosol Changes to Key Factors. In light of the complexity and the uncertainty of the halogen effects on aerosol concentration and composition, we conducted five additional sensitivity simulations (SI Table S4) to quantify the sensitivity of halogen impact on aerosol to the key parameters and processes. The simulated distribution and partitioning of chlorine and bromine in these sensitivity cases are shown in SI Figure S7. Cl_y in no_Cl_emi case and Br_y in no_Br_emi case are significantly lower than their counterpart in the HAL case. Note that Br_y (34.3 pptv) in the no_Cl_emi case is only slightly lower than that in HAL case (35.6 pptv) because all bromine species are present in the gas phase and the chlorine abundance mostly impact bromine partitioning but does not imply any direct bromine source. Meanwhile, Cl_y in the no_Br_emi case (235.0 pptv) is significantly lower than that in the HAL case (334 pptv) because the presence of bromine species (HOBr, BrNO_2 and BrNO_3) in HAL transforms HCl to BrCl leading to evaporation of semivolatile aerosol chloride ($\text{NH}_4\text{Cl} \rightarrow \text{NH}_3(\text{g}) + \text{HCl}(\text{g})$) thereby increasing the total gaseous chlorine (Cl_y). In the no_BrNO₂_uptake, the dominant species becomes BrNO_2 instead of BrCl in the HAL case. The simulated level, distribution, and partitioning of Cl_y and Br_y in fixed_gamma and low_yield cases are similar to those in the HAL case.

Figure 7 shows the average concentration of $\text{PM}_{2.5}$ and its components in northern China for all cases. Interestingly, anthropogenic chlorine emissions alone (no_Cl_emi v.s. HAL) increase $\text{PM}_{2.5}$ by $11.8 \mu\text{g}/\text{m}^3$ (10.3%), whereas the $\text{PM}_{2.5}$ increase by anthropogenic bromine emissions alone (no_Br_emi v.s. HAL) is $17.5 \mu\text{g}/\text{m}^3$ (16.1%). Our results show, for the first time, a comparable and even more important role of anthropogenic bromine emission than that of anthropogenic chlorine in producing secondary aerosols because the presence of bromine increases both the production of OH and O_3 and the loading of the total gaseous chlorine. A recent modeling study by Wang et al.¹² suggested an increase of $3.2 \mu\text{g}/\text{m}^3$ in annual average $\text{PM}_{2.5}$ due to anthropogenic chlorine emission in China. Zhang et al.⁷¹ reported an enhancement of $7.5 \mu\text{g}/\text{m}^3$ (9.1%) in $\text{PM}_{2.5}$ in China in winter after considering anthropogenic chlorine emission. Here we show that another halogen species (bromine) can have a similar or even larger contribution compared to that of chlorine alone.

The heterogeneous uptake of BrNO_2 on aerosols (forming HOBr or BrCl) efficiently recycles the halogen atoms (Br and Cl) and induces a $12.5 \mu\text{g}/\text{m}^3$ (11.0%) increase of $\text{PM}_{2.5}$ (no_BrNO₂_uptake v.s. HAL) which is comparable to that caused by anthropogenic chlorine. Wang et al.⁷² proposed that the role of BrNO_2 could be vital for sustaining active bromine chemistry in low O_3 regions. Note that BrNO_2 was not observed above the detection limit at Wangdu site²⁵ which could be due to its rapid chemical loss (photolysis and heterogeneous uptake) in the atmosphere. This is in line with SI Figure S7, which shows that if the BrNO_2 uptake does not happen (the no_BrNO₂_uptake case), BrNO_2 is the dominant Br_y species and would have been detected together with BrCl, HOBr, and Br_2 at Wangdu. In light of the modeled significant effect of BrNO_2 and its heterogeneous uptake process on Br_y partitioning and secondary aerosol production, we call for further laboratory and field studies to investigate this species and its role in aerosol formation.

Apart from BrNO_2 , other halogen species (HOCl, HOBr, ClNO_2 , and BrNO_3) also undergo heterogeneous processes. In the HAL case, we use dynamic uptake coefficients (depending on the light intensity) to mimic the good correlation between photolysis rate and bromine levels. The small difference in bromine species (SI Figure S7) and aerosols (Figure 7) between fixed_gamma ($\gamma = 0.1$, used in previous studies^{15,27}) and HAL (dynamic γ from 0.001 to >0.1) cases suggest that the halogen effects on aerosol are not sensitive to the two uptake coefficient configurations. Such a small change is observed because the heterogeneous reactions serve as recycling processes imposing larger changes to the partitioning of Br_y (e.g., larger BrCl and lower HOBr fractions in the fixed_gamma case) than to the total bromine levels.

When neglecting the direct formation of SOA from Cl- and Br-initiated VOC oxidation, the SOA enhancement due to indirect halogen effect (BASE v.s. no_ClBr_SOA) is $2.4 \mu\text{g}/\text{m}^3$; the direct effect (no_ClBr_SOA v.s. HAL) is $6.3 \mu\text{g}/\text{m}^3$. This result suggests that the direct halogen effect is comparable and even larger than the indirect halogen effect on SOA. Here we note that we compile the Cl-initiated SOA yield parametrizations based on the limited chamber studies and we call for the need for future chamber studies to further identify the key parameters in SOA formation from Cl- and Br-initiated reactions. Choi et al.⁷³ reported that 5–10 times larger anthropogenic chlorine emissions were required to reproduce the observed chlorine species in China in winter, and such large chlorine emission could lead to an SOA increase of $0.7\text{--}3.0 \mu\text{g m}^{-3}$ in terms of direct effect and $2.5\text{--}3.0 \mu\text{g m}^{-3}$ in terms of indirect effect. The present study considers both anthropogenic chlorine (no scaling is applied) and bromine emissions and more comprehensive halogen chemistry (Cl, Br, and I) and shows a comparable increase ($8.7 \mu\text{g}/\text{m}^3$) in SOA due to the full consideration of halogen sources and chemistry.

Despite the noted uncertainties in emissions and chemical mechanisms, this comprehensive sensitivity analysis shows the fundamental role of anthropogenic halogens in increasing haze pollution and the need to consider their emissions and chemistry in air pollution research and air quality regulation.

We demonstrate that reactive halogens (chlorine and bromine) substantially increase the loading of secondary aerosols ($\sim 130\%$ for SOA; $\sim 20\%$ for sulfate, nitrate, ammonium, and $\text{PM}_{2.5}$) in northern China during the winter season where and when haze pollution has been having severe health and visibility impacts in the past two decades. Our model exercise adopts a widely used model system with updated halogen chemistry and comprehensive halogen emissions. We find that bromine emissions, which have been previously unaccounted for in continental aerosol formation, exert a significant impact (larger than that of chlorine) on aerosols in northern China. The significant effect of bromine on aerosol formation arises for three reasons: directly via oxidation of aerosol precursors by bromine species, and indirectly by increasing the level of oxidants (OH, HO_2 , O_3 , and NO_3), and by enhancing gaseous chlorine levels through heterogeneous reactions.

Finally, albeit our study is restricted to China, the significant range of halogen impacts on haze pollution found here is likely widespread in other continental regions with potential anthropogenic halogen emissions from coal burning, biomass burning, waste burning, etc. For instance, a recent study in India¹⁶ reported that chlorine increased fine aerosol via gas-particle partitioning (HCl to chloride and NH_3 to ammonium)

and water uptake on existing aerosol, implying that the role of halogens in the chemical production of secondary aerosol could also be important in South Asia as well as other continental regions where anthropogenic halogens (chlorine and bromine) emissions and chemistry have also been shown to impact gas-phase atmospheric compositions^{74,75}

We call for more studies to improve the understanding of emission, chemistry, and impacts of anthropogenic reactive halogens, for example, direct measurement of bromine emission factors of various potential sources, laboratory experiments on Cl- and Br-initiated SOA formation, and field observations of the multiphase processes of halogen species, etc.

■ ASSOCIATED CONTENT

SI Supporting Information

The Supporting Information is available free of charge at <https://pubs.acs.org/doi/10.1021/acs.est.1c01949>.

Processes related to secondary aerosol formation in standard WRF-Chem, Chlorine-initiated SOA yield parametrization, WRF-Chem sensitivity design, and WRF-Chem model evaluation (PDF)

(XLSX)

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A.S.L. and T.W. initiated the research. A.S.L. devised the model simulations and analysis. Q.L., with the help of A.B. and X.F., conducted WRF-Chem development and simulation. X.F. and T.W. developed the anthropogenic chlorine and bromine emissions. X.F., with contributions from Q.L., T.W., and A.S.L., evaluated the performance of WRF-Chem model. X.P., W.W. and T.W. provided observation data of halogen species. Y.M. and J.M. provided observations of routine air pollutants. Q.L. and A.S.L. wrote the manuscript with significant contributions from X.F., J.L.J., R.P.F., A.B., C.A.C., and T.W., and contributions from all authors.

Notes

The authors declare no competing financial interest.

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