

---

This is the **accepted version** of the journal article:

Sorbelli, Leonardo; Cherin, Marco; Kostopoulos, Dimitris S.; [et al.]. «Earliest bison dispersal in Western Palearctic : insights from the Eobison record from Pietrafitta (Early Pleistocene, central Italy)». *Quaternary Science Reviews*, Vol. 301 (February 2023), art. 107923. DOI 10.1016/j.quascirev.2022.107923

---

This version is available at <https://ddd.uab.cat/record/270336>

under the terms of the  license

1 **Earliest bison dispersal in Western Palearctic: insights from the *Eobison* record from Pietrafitta**  
 2 **(Early Pleistocene, Central Italy)**

3  
 4 Leonardo Sorbelli<sup>a,\*</sup>, Marco Cherin<sup>b,\*</sup>, Dimitris S. Kostopoulos<sup>c</sup>, Raffaele Sardella<sup>d</sup>, Beniamino Mecozzi<sup>d</sup>,  
 5 Valerii Plotnikov<sup>e</sup>, Maria Prat-Vericat<sup>a</sup>, Beatrice Azzarà<sup>b</sup>, Saverio Bartolini-Lucenti<sup>a,f</sup>, Joan Madurell-  
 6 Malapeira<sup>a,f</sup>

7 <sup>a</sup>*Institut Català de Paleontologia Miquel Crusafont, Universitat Autònoma de Barcelona, Edifici ICTA-ICP,*  
 8 *c/ Columnes s/n, Campus de la UAB, 08193, Cerdanyola del Vallès, Barcelona, Spain.*

9 <sup>b</sup>*Dipartimento di Fisica e Geologia, Università degli Studi di Perugia, Via A. Pascoli, 06123 Perugia, Italy*  
 10 <sup>c</sup>*Laboratory of Geology and Paleontology, School of Geology, Aristotle University of Thessaloniki,*  
 11 *54124, Thessaloniki, Greece.*

12 <sup>d</sup>*Dipartimento di Scienze della Terra (PaleoFactory), Sapienza Università di Roma, Piazzale A. Moro 5,*  
 13 *00185 Roma, Italy.*

14 <sup>e</sup>*Academy of Sciences of the Republic of Sakha (Yakutia), 677007 Yakutsk, Lenin ave. 33, Russia.*

15 <sup>f</sup>*Dipartimento di Scienze della Terra, Università di Firenze, 50121 Firenze, Italy*

16  
 17 \* Corresponding author.

18 *E-mail address:* leonardo.sorbelli@icp.cat

19  
 20 **Abstract:** The late Villafranchian is one of the pivotal time-spans in the succession of Pleistocene

21 European faunal assemblages, setting the bases for the major faunal renewal that characterized the continent  
 22 during the Epivillafranchian. *Bison* is one of the most important and successful large mammals to spread in  
 23 Europe at the latest stages of the Early Pleistocene. Here we describe the remains of a large bovid from the  
 24 late Villafranchian site of Pietrafitta (ca. 1.5 Ma), previously attributed to *Leptobos*. Our analyses allow to  
 25 refer this sample to the genus *Bison*. The primitive characters featured by this sample suggest the attribution  
 26 to the subgenus *Eobison*, a long-debated taxon which includes all the earliest forms of *Bison*. At Pietrafitta,  
 27 thus, one of the largest and most complete record of primitive *Bison* is recorded. The vague diagnosis and  
 28 confused taxonomic history of *Eobison* called for a reappraisal of its status. We present a re-definition of the  
 29 diagnostic characters and a review of the Eurasian record of *Eobison* with a focus on the late Villafranchian  
 30 samples from the Mediterranean area. We recognize at least three valid species of *Eobison* remarking,  
 31 however, the extreme morphological variability of this group. The comparative analysis of *Eobison* and its  
 32 closest relatives (i.e., *Leptobos* and *Bison* s.s.), confirms that both the leptobovine and bisontine clades

33 underwent to an increase of stoutness of the appendicular skeleton in response to the shrinking of forested  
34 habitats and the onset of colder, arid climate during the Pleistocene.  
35 **Keywords:** *Bison*, Bovidae, Europe, Quaternary, Villafranchian

## 36 **1 Introduction**

### 37 1.1. *The genus Leptobos and its historical and taxonomic background*

38 According to most scholars the origin of *Bison* probably lies within the extinct genus *Leptobos*  
39 (McDonald, 1981; Duvernois and Guérin, 1989; Masini, 1989; Geraads, 1992; Martínez-Navarro et al.,  
40 2007; Tong et al., 2013; Cherin et al., 2019 among others). This group of mid- to large-sized and slender  
41 bovines is a constant faunal element during the whole Villafranchian European Land Mammal Age (ELMA;  
42 about 3.3–1.2 Ma; Cherin et al., 2019). Fossils of these ungulates have been found in all the Palearctic of  
43 Eurasia, from the Iberian Peninsula to China (Dong, 2008; Garrido, 2008), and from northern India to,  
44 possibly, the British Isles (Pilgrim, 1937; Breda et al., 2010). Since the establishment of the genus by  
45 Rüttimeyer (1877-1878), a large number of remains was attributed to *Leptobos* and several species were  
46 erected. According to the most recent review, *Leptobos* includes the species: *Leptobos brevicornis* (Asia),  
47 *Leptobos crassus* (Asia), *Leptobos falconeri* (Asia), *Leptobos stenometopon* (Europe), *Leptobos elatus*  
48 (Europe), *Leptobos merlai* (Europe), *Leptobos furtivus* (Europe), *Leptobos etruscus* (Europe), and *Leptobos*  
49 *vallisarni* (Eurasia) (Cherin et al., 2019).

50 Masini (1989) and Masini et al. (2013) divided the European species of *Leptobos* into two groups or  
51 lineages, the first including *L. stenometopon*, *L. elatus*, *L. merlai*, and *L. furtivus*; the second *L. etruscus* and  
52 *L. vallisarni*. Duvernois (1990, 1992), on the basis of horn and tooth morphologies established two  
53 subgenera for the aforementioned groups: *Leptobos (Leptobos)* and *Leptobos (Smertiobos)*, respectively,  
54 adding to the latter the species *Leptobos bravardi*. These two subgenera, however, suffer from nomenclatural  
55 inconsistencies (see Croitor and Popescu, 2011; Cherin et al., 2019), thus this subdivision is not commonly  
56 used in literature (Masini et al., 2013). The Asian group includes the species *L. falconeri* (type species), *L.*  
57 *brevicornis* (which embraces also *L. amplifrontalis* and *L. laochihensis*), and *L. crassus* (Mead et al., 2014).  
58 Over the years, several authors tried to find relationships between the Asian and European forms (Dong,  
59 2008). However, up to date, it is still not clear if these groups are actually related or if, for instance, the  
60 Asian taxa are part of a third, distinct, clade (Mead et al., 2014). This incapability of linking the European  
61 and Asian groups might find a solution in the species *L. vallisarni*, which is the only occurring in both  
62 continents (Zheng et al., 1985; Tong et al., 2016). The existence of advanced *Leptobos* forms in the Early  
63 Pleistocene of eastern Asia, along with the fact that the earliest *Bison sensu lato* (s.l.) species are known from

64 India and China roughly in the same time span, led most authors to suggest that *Bison* could have originated  
65 in Asia, emerging from a derivate stock of *Leptobos* (Sala, 1986; Masini, 1989; Duvernois, 1990; Martinez-  
66 Navarro et al., 2007; Sorbelli et al., 2021a among others).

67

### 68 1.2 Earliest Eurasian bison in the paleontological literature, a review

69 The earliest forms of *Bison* most probably emerged in the eastern Palearctic and, as soon as they reached  
70 Europe at the end of the Early Pleistocene, started to replace *Leptobos* in all the large herbivore guilds.  
71 Indeed, a possible co-existence between *Bison* and *Leptobos* is reported only in few late Villafranchian  
72 European localities of south-eastern Europe (Duvernois, 1990; Agadzhanyan et al., 2017; Kostopoulos et al.,  
73 2018; Lopatin et al., 2019). In Asia, on the contrary, *Bison* and *Leptobos* seem to have shared the same  
74 geographic areas for a prolonged time, possibly for almost two million years (Tong et al., 2016). This long  
75 overlap could be explained (i) by the large size of suitable habitats in Asia, which may have triggered niche  
76 partitioning among these bovines in a context of wide trophic availability, and/or (ii) by bias in the  
77 taxonomic identification and/or stratigraphic range of fossils.

78 The first occurrence of *Bison* s.l. is represented by the scanty remains of *Bison* (*Eobison*) *sivalensis* and  
79 *B. (Eobison) palaeosinensis* from India and China, respectively. Flerov (1979) referred to the newly erected  
80 subgenus *Eobison* these two Asian species, as well as *B. (Eobison) tamanensis* from Eastern Europe. The  
81 author, however, did not deepen the diagnosis of the subgenus, which is currently flawed by several issues.  
82 *Bison (Eobison) georgicus* from Dmanisi (Georgia) is putatively the earliest *Bison* s.l. approaching Europe at  
83 ca. 1.77 Ma (Burchak-Abramovich and Vekua, 1992; Bukshianidze, 2005). From the late Villafranchian on,  
84 a large number of localities in Mediterranean and Eastern Europe are characterized by the presence of these  
85 mid-sized bison-like bovids (Fig. 1), still commonly grouped in the subgenus *Eobison* (Masini et al., 2013;  
86 Sorbelli et al., 2021a). After 1.2 Ma (i.e., at the beginning of the Epivillafranchian stage), with the onset of  
87 new climatic conditions in the context of the Early-Middle Pleistocene Transition (EMPT; Head and  
88 Gibbard, 2005; Clark et al., 2006), *Eobison* was replaced by the larger-sized *Bison sensu stricto* (s.s.) (i.e., *B.*  
89 *schoetensacki*, *B. menneri*, and *B. priscus*, in chronological order of appearance), which rapidly colonize  
90 most of the Holarctic (Kahlke, 1999; Sorbelli et al., 2021a).

91 In the time interval between the age of the Dmanisi fauna and the beginning of the Epivillafranchian,  
92 crucial information on the evolution of early bison can be sought, but, unfortunately, paleontological records  
93 are scarce. Here, we describe the rich sample from Pietrafitta (central Italy), which can shed new light on this  
94 intriguing topic. In a preliminary analysis, this sample was attributed to *Leptobos* aff. *vallisarni* based on  
95 cranial similarities with the holotype of *L. vallisarni* from the Upper Valdarno (Italy; Masini, 1989). Later  
96 on, this attribution was confirmed by Masini and Gentili (2005), although not supported by quantitative  
97 comparative analysis. Nevertheless, the authors noticed that the material from Pietrafitta was characterized  
98 by some features more advanced than in *L. vallisarni*, hence the choice of open nomenclature. Building on  
99 that insight, this paper presents the first detailed description of the Pietrafitta record, in the context of the  
100 arrival of the first bison in Europe and their replacement of *Leptobos*.

101

## 102 **2. Geological setting**

103 The Tavernelle-Pietrafitta Basin is located in the upper valley of the Nestore River. During the early  
104 stages of the Pleistocene, after an intense phase of tectonic activity, the Tavernelle-Pietrafitta Basin became  
105 part of a lateral branch of the main Tiberino Basin and started to be characterized, alternatively, by  
106 lacustrine/palustrine and alluvial plain conditions (Conti and Esu, 1981; Cherin et al., 2012; Pazzaglia et al.,  
107 2013). The lignite deposition was a result of the elevation of the delta of the Nestore paleo-river causing the  
108 formation of several small basins filled with stagnant water (Gentili et al., 1996). At the end of the  
109 Pleistocene, new tectonic events occurred causing an uplift of the area and the desiccation of the basin,  
110 which subsequently underwent an erosive phase (Ambrosetti et al., 1989). The site of Pietrafitta  
111 ( $42^{\circ}59'33''\text{N } 12^{\circ}12'42''\text{E}$ ) is located close to the homonymous village in the municipality of Piegaro. The  
112 lignite seam had a thickness comprised between 5 and 8.5 m and was exploited industrially for almost half a  
113 century. Lignite was intercalated with thin layers of organic clay yielding freshwater mollusks and  
114 oligohaline ostracods (Gliozzi et al., 1997). Facies analysis suggests that the lignite was deposited in a  
115 swampy area located on the edge of a lake (Conti and Girotti, 1977; Ambrosetti et al., 1992; 1995;  
116 Martinetto et al., 2014). This hypothesis is furtherly confirmed by paleobotanical studies, which report the  
117 presence of a large water body surrounded by broad-leaved deciduous forests in a humid climatic context.  
118 The plant assemblage of Pietrafitta is dominated by aquatic genera such as *Azolla*, *Najas*, *Nymphaea*, and

119 *Potamogeton* (Martinetto et al., 2014). Palynological data show a variation in the forest composition  
120 surrounding the swamp, due to a progressive harshening of the climatic conditions. In particular, moving  
121 along the stratigraphic succession from bottom to top, the disappearance of *Taxodium*-type and its  
122 replacement by “Quercetum” taxa is observed, followed by the spread of herbaceous plants and *Pinus* in the  
123 upper layers (Lona and Bertoldi, 1972; Kottek et al., 2006; Martinetto et al., 2017). The depositional and  
124 environmental conditions of the swamp allowed the preservation of one of the richest continental vertebrate  
125 samples in the Italian Early Pleistocene, including some thousand specimens belonging to almost 40 taxa of  
126 freshwater fishes, anurans, reptiles, birds, and mammals (see revised faunal lists in Martinetto et al., 2014  
127 and Sorbelli et al., 2021b). The vertebrate assemblage is referred to the Farneta Faunal Unit (FU; approx.  
128 1.6–1.4 Ma) of the late Villafranchian ELMA, corresponding to late MNQ 18 in the French mammal  
129 biochronology scheme (Guérin, 1982; Torre et al., 1992; Gliozzi et al., 1997; Gentili et al., 2000; Rook and  
130 Martínez-Navarro, 2010). The taphonomic evidence and the presence of numerous skeletons of large  
131 mammals found in their original anatomical position or slightly scattered, suggest a low degree of post-  
132 mortem movement (Ambrosetti et al. 1992).

133

### 134 **3. Materials and Methods**

135 More than 400 remains of a large bovid have been recovered from the lignite deposits of Pietrafitta (Table  
136 S1). Most of the material come from the “Miniera Vecchia” and few from “Poderetto” and “Poderone”  
137 excavations (Gentili et al., 1996). The studied sample is stored in the “Museo Paleontologico Luigi Boldrini”  
138 at Pietrafitta except for a metacarpal (IGF 1011), a cast of the original fossil which is lost, housed in the IGF  
139 (see abbreviations below). The fossils present different state of preservation and colours, suggesting that the  
140 bones undergone to different taphonomic processes. Most of the specimens share the same kind of  
141 preservation with dark grey to dark brown colour, mildly to strongly deformed volumes, and slightly to  
142 strongly cracked surfaces testifying a certain degree of subaerial exposure before final burial. Few of the  
143 cranial and postcranial remains are characterized by dark greyish colour and are less deformed or cracked.

144 In this review of the primitive forms of *Bison* we attempt to use the largest comparative sample available,  
145 including both the putative ancestors and descendants of this group (i.e., *Leptobos* and *Bison* s.s.). However,  
146 due to the difficulties in studying the original material, lack of data in the literature, debated taxonomy, or

147 scanty fossil record, we kept out from the discussion some samples, e.g., *Leptobos* spp. from China and *L.*  
148 *furtivus* and *L. bravardi* from Europe. The two *Leptobos* groups considered in this work are thus, the one  
149 including *L. stenometopon*, *L. elatus*, and *L. merlai* (gr. LSEM hereinafter) and the one including *L. etruscus*  
150 and *L. vallisarni* (gr. LEV hereinafter). A large number of *Leptobos* specimens reported by Masini (1989)  
151 from unknown localities of the Upper Valdarno and referred to *L. etruscus*, *L. cf. etruscus*, or *L. cf. vallisarni*  
152 were not considered due to their undefined stratigraphic provenience and uncertain taxonomy. The genera  
153 *Bos*, *Bison*, *Poephagus* (as well as other bovines) are often grouped within *Bos* based on molecular evidence  
154 (Groves and Grubb, 2011; Hassanin, 2014; Castelló, 2016). In order to keep consistency with decades of  
155 paleontological studies, we follow the nomenclature by Kostopoulos et al. (2018) which retain *Bison*  
156 (including both extinct and extant species) as a separate genus. In this context, *Bison* s.l. is used to indicate  
157 members of the three subgenera *Bison* (*Bison*), *Bison* (*Poephagus*), and *Bison* (*Eobison*).

158 The fossil material directly studied by us includes the following records (for the institutional  
159 abbreviations see below): *Leptobos* spp. from Villarroya (Spain), Upper Valdarno, Pantalla, and Umbria  
160 undetermined locality (Italy), housed in the ICP, IGF, NHMB, and SABAP\_UMB; *B. (E.) degiulii* from  
161 Pirro Nord (Italy), housed in the DST and PF; *B. (E.) degiulii* from Mygdonia Basin (Greece) housed in the  
162 LGPUT and NHCK; *B. (E.) cf. degiulii* from Capena (Italy) and *B. (B.) cf. schoetensacki* from Cava  
163 Redicicoli (Italy) housed in the MUST; *B. (E.)* sp. from Mugello (Italy) housed in the MGCB; *B. (E.) cf.*  
164 *georgicus* from Salita di Oriolo (Italy) housed in the MCSNF; *B. (B.) schoetensacki* from Isernia La Pineta  
165 (Italy), Cesi (Italy), and Vallparadís Section (Spain) housed in the MPPPL, MuPA, and ICP, respectively; *B.*  
166 (*B. priscus* from Siberia (Russia), housed in the Yakutia AS. The digital material studied by us includes the  
167 3D models of *B. (E.) georgicus* from Dmanisi (Georgia) housed in the MG; *B. (B.) bonasus* from Białowieża  
168 (Poland) housed in the MRI PAS; *B. (B.) latifrons* from several North American sites housed in the IMNH.  
169 Other comparative data were retrieved from the literature. Juvenile specimens are not included in the study.  
170 The measurements of the fossils were taken with digital and analogic callipers and recorded at the nearest 0.1  
171 mm. The measurements on the digital 3D models were taken with the aid of the software MeshLab (Cignoni  
172 et al., 2008). Statistical analyses were performed with PAST v.4 (Hammer et al., 2001). The anatomical  
173 terminology used in this paper are from Heintz (1970), Sala (1986), Masini (1989), Kostopoulos et al.

174 (2018), and Sorbelli et al. (2021a) with minor changes. Linear measurements follow Masini (1989) and  
175 Sorbelli et al. (2021a) with minor changes, and are explained in Table S2 and shown in Fig. S1.

176 The shape and size of horn-core bases were defined by bivariate plot comparing the anteroposterior and  
177 dorsoventral diameters (HDV vs HAP). The shape and size of the intertemporal bridge and frontal bones  
178 were analysed with bivariate plots featuring IBH vs IBW and FWmin vs FLmin, respectively. Log<sub>10</sub> diagrams  
179 of teeth (Simpson, 1941) were constructed based on the average values of eight selected variables of  
180 *Leptobos* and *Bison* samples from various Eurasian sites. The record of *B. (B.) priscus* from Krasny Yar, R–  
181 W was used as a standard of comparison ( $y=0$ ; data taken from Vasiliev, 2008). To assess the stoutness of  
182 limb bones, we used bivariate plots featuring Lmax vs DEW/Lmax % for the metapodials and astragali and  
183 Lmax vs PEW/Lmax % for the radii. Principal component analyses (PCAs) were performed based on  
184 metapodials and cranial remains to estimate the differences between *Leptobos* and *Bison* species. These  
185 analyses rely on seven Mosimann Log-Shape variables for the metapodials and six for the crania (Table 1),  
186 obtained by log-transforming the ratio between each measurement and the geometric mean of the  
187 measurements for each specimen (Jungers et al., 1995). Histograms were employed to determine the length  
188 of metatarsals and radii relatively to the metacarpals, featuring the average values of LmaxMT/LmaxMC-  
189 100 and LmaxRAD/LmaxMC%-100, respectively. The robusticity index of metacarpals was defined by the  
190 ratio DEW/Lmax %. Body mass estimations were obtained as average of three estimations based on  
191 metatarsal linear measurements (PEW, PET, and DEW) following Scott (1983), taken in consideration only  
192 complete specimens.

193 *Institutional abbreviations*—DGPU, Department of Geology, Panjab University, Chandigarh, (India);  
194 DST, Dipartimento di Scienze della Terra, Università di Firenze (Italy); GIN RAS, Geological Institute of  
195 Russian Academy of Sciences, Moscow (Russia); ICP, Institut Català de Paleontologia Miquel Crusafont,  
196 Sabadell (Spain); IGF, Museo di Storia Naturale, Sezione di Geologia e Paleontologia, Università di Firenze  
197 (Italy); IGPB, Istituto di Geologia di Bari; IMNH, Idaho Museum of Natural History, Pocatello (USA);  
198 LGPUT, Museum of Geology, Palaeontology and Palaeoanthropology of the Aristotle University of  
199 Thessaloniki (Greece); MAEGR, Museo Arqueológico y Etnológico de Granada (Spain); MCSNF, Museo  
200 Civico delle Scienze Naturali di Faenza (Italy); MG, S. Janashia Georgian National Museum, Tblisi  
201 (Georgia); MPLB, Museo Paleontologico “Luigi Boldrini” di Pietrafitta, Piegara (Italy); MPPPL, Museo di

202 Paleontologia e Preistoria Piero Leonardi, Università di Ferrara (Italy); MGCB, Museo di Geologia e  
203 Paleontologia “Giovanni Capellini”, Università di Bologna (Italy); MR PAR, Mammal Research Institute,  
204 Polish Academy of Sciences (Poland); MuPA, Museo Paleontologico Archeologico di Serravalle di Chienti  
205 (Italy); MUST, Museo Universitario di Scienze della Terra, Sapienza Università di Roma (Italy); NHCK,  
206 Natural History Collection of Kalamoto, Kalindoia (Greece); NHMB, Natural History Museum of Basel  
207 (Switzerland); PF, PaleoFactory Lab, Dipartimento di Scienze della Terra, Sapienza Università di Roma  
208 (Italy); PIN, Orlov Paleontological Museum, Moscow (Russia); SABAP UMB, Soprintendenza Archeologia  
209 Belle Arti e Paesaggio dell’Umbria, Perugia (Italy); Yakutia AS, Academy of Sciences Republik of Sakha,  
210 Yakutsk (Russia); ZIN RAS, Zoological Institute of Russian Academy of Sciences, St. Petersburg (Russia);

211

#### 212 **4. Results**

213 Class Mammalia Linnaeus, 1758

214 Order Artiodactyla Owen, 1848

215 Family Bovidae Gray, 1821

216 Subfamily Bovinae Gray, 1821

217 Genus *Bison* Hamilton Smith, 1827

218 Subgenus *Eobison* Flerov, 1972

219

220 *Emended diagnosis*—Small to mid-sized *Bison* with small horns (present in both males and females) and  
221 slender limbs. Cranium with relatively deep and high temporal fossae; lower than in *Leptobos* but higher  
222 than in *Bison* (*Bison*). The fossae extend far behind the horn-cores, along the lateral sides of the postcornual  
223 part of the cranium (not so much as in *Leptobos*). As a result, the occipital squama is separated from the  
224 frontal by an abrupt constriction, giving to it a trapezoidal to semi-circular or bell shape with a wide base and  
225 a narrow top. Nuchal crest outline, in dorsal view, undulated or semicircular, never squared or straight.  
226 Supraoccipital portion of the cranium quite reduced; intertemporal bridge almost completely fused with the  
227 nuchal crest. Orbits protruding anterolaterally, weakly tubular, without thickened edges. Postorbital notch  
228 variable depending on the sex but, generally, stronger than in *Leptobos* and weaker than in *Bison* (*Bison*).  
229 Frontals short; their length (distance from the edge of nuchal crest to nasion) is shorter than the width of  
230 frontals between postorbital notch. Frontals concave and pneumatized at the horn-core bases; elevated above  
231 the occipital area. Horn-core pedicles shorter than in *Leptobos*. Horn-cores short and stout, emerging

232 posterolaterally (the angle between the sagittal axis and the midline of the horn-core is comprised between  
233 50° and 75°); horn-cores bending posteriorly and ventrally at the base, then laterally and markedly dorsally  
234 at the tip; in some cases, there can be a small anticlockwise torsion; horn-cores dorsoventrally compressed at  
235 the base; section oval in the first two third, tending to get circular toward the tips; absence of keels;  
236 posteroventral surface strongly furrowed. Premaxillary bone ends before the nasals. Nasals elongated, ending  
237 at the level of the anterior margin of the orbit or slightly behind. Mandible with relatively straight corpus and  
238 short symphysis. Angle between ramus and corpus around 90° or slightly obtuse. Coronoid process projects  
239 backwards, often exceeding the level of processus condyloideus. Teeth are generally smaller than in *Bison*  
240 (*Bison*). Neural processes of the thoracic vertebrae well developed but less than in *Bison* (*Bison*). Limbs  
241 relatively slender and short. Metapodials stouter than in *Leptobos* but slenderer than in *Bison* (*Bison*).  
242 Metatarsal less than 20% longer than metacarpal. Radius generally shorter than in *Bison* (*Bison*) but longer  
243 than in *Leptobos*; more than 30% but less than 50% longer than the metacarpal.

244 *Bison (Eobison) degiulii* Masini, Palombo and Rozzi, 2013

245 (Figs 2–5; Figs S2–S6)

246 *Diagnosis*—Mid- to large-sized *B. (Eobison)*. Frontals concave between the horn-cores with a small crest  
247 running along the suture line in the posterior portion. Very short and stout horn-cores. Postcornual portion of  
248 the cranium moderately developed. Temporal fossae wide and stretched posteriorly. Nuchal crest outline in  
249 dorsal view, arched, feebly undulated. Occipital squama trapezoidal or bell shaped.

250

251 *Referred specimens*—The Pietrafitta record includes 93 cranial (Figs 2–4) and 337 postcranial remains  
252 (Fig. 5 and Figs S2–S6). See Table S2 for the complete list.

253

#### 254 4.1. *Description*

##### 255 4.1.1. *Cranial material*

256 Five incomplete crania were unearthed from Pietrafitta, all affected by taphonomic deformation. The most  
257 complete crania are SABAP UMB 19. 2.1178 and SABAP UMB 19. 2.1179, both flattened in dorsoventral  
258 direction. The former preserves most of the frontals, nasals, and the left horn-core whereas the latter is  
259 represented by the occipital and basioccipital areas and the horn-cores. In both specimens four of the upper

260 molars (left and right M2s and M3s) and three not identified upper premolars are still embedded in the  
261 cranium or matrix. SABAP UMB 19. 2.1178 is associated with its two hemimandibles. The ossified sutures  
262 and the stage of wear of teeth (were present) suggest that all the five crania belong to adult individuals. The  
263 small size of the horn-cores and the reduced tuberosities of the basioccipital indicate that SABAP UMB 19.  
264 2.1178 was probably a female whereas SABAP UMB 19. 2.1179 is characterized by larger horn-cores and a  
265 massive basioccipital, indicating it is a male. The horn-cores are located rather close to the orbits but far  
266 from each other and emerge posterolaterally (i.e., the angle between the sagittal axis and the midline of the  
267 horn-core is around  $65^\circ$ ). Compared to the overall cranium size, the horn-cores are small and stout whereas  
268 the pedicles are relatively long. The horn-cores are thicker at the pedicle contact and taper abruptly. Their  
269 section at the base is ovoidal with the major axis developing in antero-posterior direction, then it tends to  
270 become circular toward the tips. The horn-cores weakly curve ventrally in the first third, then bend laterally  
271 and dorsally, then posteriorly and somewhat medially (especially in SABAP UMB 20. 1.6803). The ventral  
272 surface is deeply grooved whereas the dorsal one is less furrowed. The grooves tend to disappear toward the  
273 apexes. In some specimens (i.e., SABAP UMB 21. 1.1323, SABAP UMB 20. 2.7345) the dorsal side of the  
274 tips is marked by deep irregular furrows that run parallel to the major axis of the horn-core. In all the  
275 specimens the tips are marked by numerous small nutrient foramina. In the broken specimen SABAP UMB  
276 21. 1.1179 it is possible to observe that the most basal part of the horn-core is strongly pneumatized. The  
277 horns and pedicles are elevated above the frontals, even if the compression does not permit to assess the  
278 degree of elevation. In SABAP UMB 19. 2.1178 the frontal area is quite wide and, although strongly  
279 deformed, it is possible to observe a slight concavity in the central part, between the pedicles. A small crest  
280 runs in the middle of this depression, following the interfrontal suture. This suture is very clear as well as the  
281 bregma. The supraorbital foramina are far from each other, positioned just behind the orbits, in narrow but  
282 elongated grooves. These grooves are anteriorly converging and end at the level of the posterior margin of  
283 the orbits. The postorbital constriction is strongly convex. One orbit is preserved in SABAP UMB 19.  
284 2.1178. Although the deformation altered the outline of the orbit, a certain tubular shape can be still  
285 recognized. The nasal bone wedges into the frontals with an acute angle, posteriorly to the anterior margin of  
286 the orbits. The nasals anterior end is bifurcated. The maxilla thins towards its anterior end expanding below  
287 the premaxilla, which prolongs anteriorly with a shovel-like knot in the anterior edge. The contact between

288 maxilla and premaxilla ends above the alveolar level. The posterior part of the premaxilla tapers rapidly and  
289 abruptly ends before the nasals. The postcornual portion of the neurocranium is relatively reduced in both  
290 SABAP\_UMB 19.2.1178 and SABAP\_UMB 19.2.1179. The intertemporal bridge (=chignon, sensu  
291 Duvernois, 1990) is fused with the nuchal crest which is characterized by a vaulted dorsal outline, and  
292 slightly undulated edges. The occipital squama is preserved in SABAP UMB 19. 2.1179 and SABAP UMB  
293 21. 2.1177; it is large and has a semicircular outline, marked by a narrow top and a wide base. The two  
294 temporal fossae indent behind the horns, nonetheless their posterior ends are well separated by the dorsal  
295 margin of the squama. The jugular processes are hooked, bending inward and projecting below the ventral  
296 margin of the condyles. The bulging and robust mastoid processes are separated from the jugular processes  
297 by a shallow groove. The condyles are large and sub-triangular with the pointed end located ventrally. The  
298 position and number of the foramina on the occipital squama are variable (e.g., a single foramen is visible  
299 dorsally to the right condyle of SABAP UMB 21. 2.1177). The basioccipital has a wedged shape formed by  
300 the two pairs of tuberosities. Overall, it is short and broad. The posterior tuberosities are triangular and wide  
301 and separated by a shallow depression. A faint, small crest starts between the anterior bases of these  
302 tuberosities, follows the median axis of the basioccipital, and reaches the anterior tuberosities. The latter are  
303 elongated but much smaller than the posterior ones, and are divided by a deep V-shaped furrow. Anteriorly  
304 to the anterior tuberosities there is a secondary crest, which is the prolongation of the one mentioned above.  
305 A small foramen is located on the posterior margin of this crest, just at the anterior limit of the tuberosities.  
306 SABAP UMB 19. 2.1179, SABAP UMB 21. 2.1177 and SABAP UMB 20. 1.6803 differs for having: larger  
307 posterior tuberosities divided by a narrower depression; wider and shorter anterior tuberosities; less  
308 developed crest in the valley between the anterior and posterior tuberosities; more developed crest anteriorly  
309 to the anterior tuberosities; and an overall larger size. In SABAP\_UMB 20.1.6802 the basioccipital still  
310 articulates with a portion of the basisphenoid and the pterygoid processes. The basisphenoid is strongly bent  
311 upward. A small canal is present on both sides of the basioccipital, just above the posterior tuberosities.

312 The most complete hemimandibles (SABAP UMB 130074 and SABAP UMB 130075) associated to  
313 SABAP UMB 19. 2.1178 are in good state of preservation except for some parts of the rami. The angle  
314 formed by the ramus and the corpus is slightly obtuse. The symphysis is relatively short with a bulge on the  
315 ventral edge and a rugose surface. In dorsal view, the symphyseal area is slightly inflated as well as the

316 molar portion of the mandibula. SABAP UMB 21. 1. 1298 is the only mandible preserving the processus  
317 condyloideus. The condyle is stout and depressed in the middle. The coronoid process is quite elongated and  
318 projects backwards, ending posteriorly to the processus condyloideus.

#### 319 4.1.2. *Dentognathic material*

320 The studied specimens include 28 isolated teeth, and 70 associated to mandibles.

321 The upper teeth are quadrangular in occlusal shape (especially the molars) with a wide base and high  
322 crown. In the P2, the parastyle and paracone are shifted mesially and the paracone fold is quite deep. The  
323 metastyle is poorly separated from the labial wall of the tooth. The enamel central cavity has a complex  
324 outline, being divided into three islets in SABAP UMB 20. 1.7361. The P3 is characterized by a well-  
325 developed metastyle, protruding distally. The parastyle is relatively reduced. The paracone is shifted toward  
326 the metastyle, oblique in labial view. The P4 is the largest and stoutest premolar and has a squared occlusal  
327 outline. It shows various degree of molarization. The metastyle is slightly more pronounced than the  
328 parastyle, the former protrudes distolabially whereas the latter mesiolabially. The paracone is not particularly  
329 protruding, its fold is slightly shifted toward the metastyle and it is divided from the styles by two  
330 pronounced furrows. The outline of the central cavity is U-shaped with a more complex distal wall, often  
331 featuring a small fold. SABAP UMB 20. 1.6804/8 and SABAP UMB 20. 1.6804/7, belonging to the same  
332 individual, show the most incipient molarization with a faint hypocone developing distolingually and the  
333 presence, in the latter specimen, of a small enamel islet distally to the central cavity. The upper molars show  
334 a high degree of hypsodonty with strong pillars and almost parallel margins, slightly diverging occlusally. In  
335 all the upper molars the distal lobe is slightly shifted labially. In M1 and M3 the mesial lobe is  
336 buccolingually more developed than the distal one whereas, in M2, the two lobes are generally of the same  
337 size. The styles are strong and labially projected, marked by a small constriction at the contact with the tooth  
338 which forms ribbon-like structures. In M1 and M2 the parastyle is slightly larger than the mesostyle and  
339 metastyle which are, more or less, of the same size. The metastyle is strongly projected distally in M3,  
340 creating, thus, a vertical, shallow furrow that delimitates its distal contact with the main body of the lobe. In  
341 all upper molars the paracone and metacone are less prominent than the styles. In occlusal view, the entostyle  
342 shows a subcircular (e.g., SABAP UMB 20. 1.6804/5) to relatively complex outline (e.g., SABAP UMB 20.  
343 1.7346/1). The cement is always present in upper molars, especially on the lingual side; in some specimens it

344 is scarce (e.g., SABAP UMB 20. 1.6806), but in few teeth is present in both the labial and lingual sides (e.g.,  
345 SABAP UMB 20. 1.7346/1). The cement, in the lingual side, penetrates deeply between the entostyle and the  
346 two lobes. The outline of the central cavities has an irregular crescent-moon shape, sometimes characterized  
347 by the so-called “bubaline fold” (sensu Merla, 1942). This fold is always present in the distal wall of the  
348 distal cavity, in some cases in both the cavities (e.g., SABAP UMB 20. 1.6804/3, SABAP\_UMB 20. 1.7348)  
349 although is always more complex in the distal one. No enamel islets were found between the two lobes.

350 The lower teeth are mesiodistally elongated with high crowns. The p3 is quite similar to the p4,  
351 differentiating in the less developed paraconid and the more joined entoconid/entostylid, a deeper distal  
352 valley developed mesially and a less globular metaconid. The p4 is narrow and, in most of the specimens,  
353 relatively shortened. The parastylid is sharp and projects mesiolingually, divided from the paraconid by a  
354 shallow and short groove that ends well before the neck (1<sup>st</sup> valley, sensu Heintz, 1970). In some specimens  
355 this groove is quite reduced (e.g., SABAP UMB 20. 1.7342). The paraconid is perpendicular to the  
356 mesiodistal axis of the tooth. The 2<sup>nd</sup> valley separating the paraconid from the metaconid, is deep and wide  
357 (but less in some specimens, e.g., SABAP UMB 20. 1.7340/3). The protoconid is quite developed, especially  
358 lingually, creating a small stylid-like bulge in the 2<sup>nd</sup> valley, mesially to the metaconid. The metaconid itself  
359 is a robust pillar and projects lingually (e.g., SABAP UMB 20. 1.7342) or distally with a thinner outline  
360 (e.g., SABAP UMB 21. 1.1236). Toward the cervix the metaconid is mesiodistally enlarged and shows a flat  
361 lingual surface. The 3<sup>rd</sup> valley, due to the constriction of the metaconid, is quite deep but ends well before the  
362 neck of the tooth. The entoconid and entostylid are almost completely fused, separated by a faint 4<sup>th</sup> valley  
363 which is more or less developed depending on the specimen but tends to disappear rapidly with the wear of  
364 the tooth. The labial valley separating the protoconid from the hypoconid is deep and short, ending before the  
365 base of the tooth. The molars have large and sharp metastyle projecting mesiolabially (especially m2 and  
366 m3). The protoconid is more developed mesiodistally than the hypoconid. The ectostylid is narrow and  
367 located in the mid-point between the two lobes, slightly shifted toward the mesial one. In m2 and m3 the  
368 ectostylid has an elliptical shape in occlusal view whereas it has a more complex outline in m1. Labially the  
369 molars are flattened, the protoconid and hypoconid are lingually separated by an open valley. The  
370 hypoconulid size varies depending on the specimen from mesiodistally elongated (e.g., SABAP UMB 20.

371 1.7340/5) to short and robust (e.g., SABAP UMB 20. 1.7341/6). All molars are characterized by a thick layer  
372 of cement on both sides.

#### 373 4.1.3. *Metapodials*

374 Eight metacarpals are known from the lignite of Pietrafitta. Four specimens (SABAP UMB 20. 1.7324,  
375 SABAP UMB 20. 1.7325, SABAP UMB 130038, IGF 1011) are almost complete and little deformed. There  
376 are no significant differences in the shape of the proximal facets neither in the trochlear surfaces of the  
377 analyzed specimens. Two metacarpals belong to the same individual (SABAP UMB 22. 4.247), being part of  
378 an almost complete skeleton. Generally, all specimens are characterized by a relatively slender built; the two  
379 epiphyses have roughly the same width and are much wider than the diaphysis. Seen from above, the  
380 proximal outline of the bone is D-shaped; the anterior margin has subrounded edge whereas the posterior one  
381 is markedly straight. The proximal articular surface is composed by two facets. The lateral one for the  
382 unciform is triangular and smaller than the medial for the trapezoid-magnum which is sub-quadrangular with  
383 rounded anterior margin. The former is separated from the latter by a small crest perpendicular to the  
384 mediolateral axis of the articulation. The lateral facet is anterolaterally marked by a small notch. The  
385 proximal foramen is visible only in SABAP UMB 20. 1.7324 and is located in a deep and narrow groove, at  
386 the posterior end of the crest that divides the two facets. A foramen is present distally to the proximal end, in  
387 the middle of the epiphysis. The anterior margin of the proximal epiphysis is marked by a large, protruding  
388 tuberosity. The diaphysis has hourglass-shaped medial and lateral margins, the minimum width of the shaft is  
389 located slightly above mid length. The anterior vascular groove is narrow and shallow, deeper in the distal  
390 portion. The ovoidal foramen is located just above the distal end of the vascular groove. The contact area  
391 between the diaphysis and the distal epiphysis is marked by an evident tubercle (in both medial and lateral  
392 sides), so that the maximum distal width between these tubercles is almost the same as that across the  
393 trochleae. A slight mediolateral constriction is present between these tubercles and the distal articulation  
394 surface of the corresponding trochlea. The intertrochlear margins converge distally. The distal intercondylar  
395 crests are generally sub-parallel. Two deep depressions are present proximally to either side of each trochlear  
396 crest, on the anterior side of the epiphysis. The two outer depressions are larger and deeper than the inner  
397 ones. The same four pits are present on the posterior side of the bone, dorsally to the crests. The lateral and

398 medial trochlear pits are deep and marked by radial rugosities. The lateral abaxial hemicondyle is slightly to  
399 strongly anteroposteriorly thinner than the medial one but is laterally more protruding.

400 Twelve metatarsals were recovered from Pietrafitta, of which six are well preserved. These bones are  
401 hourglass shaped with a slender and long shaft and a stout distal end. The proximal epiphysis hosts two  
402 major articulation surfaces (anterolateral and anteromedial facets for the articulation with the cubonavicular  
403 and cuneiform) and two secondary ones (posterolateral and posteromedial facets for the accommodation of  
404 the two cubonavicular posterior facets). The two primary facets converge anteriorly, separated by a well-  
405 developed ridge, proximally protruding in the anterior portion. The anteromedial articular facet is ovoidal,  
406 having the major axis oblique respect to the sagittal plane. The anteromedial margin is proximally concave  
407 and irregular whereas the posterolateral is flattened. The lateral facet is slightly smaller and less inclined with  
408 smoother outline. The angle between the two facets is variable between 40° and 45°. The posterolateral  
409 articular facet is small and leaf shaped with a very narrow anteroposterior diameter. The posteromedial facet  
410 has a subrounded outline and lies attached to the posteromedial margin of the medial primary facet. The  
411 proximal portion of the diaphysis is delimited by three crests that follow the main axis of the bone. The  
412 posterior surface of the diaphysis is slightly concave to flat in its proximal half, where two faint and shallow  
413 longitudinal grooves are present. Toward the distal portion, the shaft flattens and is marked by an ovoidal  
414 foramen. The epitrochlear tubercles at the contact between the diaphysis and the distal epiphysis are well  
415 developed, more than in the metacarpals; thus, the width at the level of the tubercles is about the same as that  
416 between the outer trochlear margins or even larger (e.g., SABAP UMB 20. 1.7328 and SABAP UMB 20.  
417 1.7330). The distal epiphysis slightly bends posteriorly. The two trochlear ridges are subparallel relative to  
418 the medial and lateral margins of the distal end and converge anteriorly. The lateral and medial trochlear pits  
419 are deep and proximally surrounded by radial rugosities.

420

#### 421 4.2. Comparisons and taxonomic assignment of the Pietrafitta sample

422 Although the Pietrafitta specimens are characterized by different taphonomic features, on the whole the  
423 record is homogeneous, both morphologically and metrically (Figs 2–9; Tables 2–4; Figs S2–S6; Tables S3,  
424 S5, S7–S9, S11, S13–S16, S18, S20, S22, S24, S28). The hypothesis that the bovid sample is roughly  
425 isochronous is supported by the overall uniformity of the faunal and floral assemblages along the Pietrafitta

426 stratigraphic succession and by the high sedimentation rate estimated for this kind of brown coal deposits  
427 (Cameron et al., 1989; Martinetto et al., 2014).

428 The general morphology of the cranial and postcranial remains from Pietrafitta is typical of bovines, and  
429 in particular of leptobovines and bison. Most of the systematics of these bovids has been historically based  
430 on skulls and metapodials but, as suggested by recent studies (e.g., Sorbelli et al., 2021a), other postcranial  
431 bones might provide useful taxonomic information.

432 The Pietrafitta specimens are different from the members of the earliest *Leptobos* species (gr. LSEM) in  
433 several cranial features including: shorter braincase (postcornual area), quite reduced intertemporal bridge,  
434 shorter and stouter horn-cores, nasal ending at the orbit level, premaxilla not in contact with the nasal, and  
435 abundant cement in the upper molars (Figs 6, 7; Table 5). Equally, the larger size and robusticity of all the  
436 postcranials clearly separate the Pietrafitta bovid from the LSEM leptobovines (Table 6; Tables S10, S12,  
437 S19, S23, S25, S26). At the same time, the increased stoutness and larger limb size of the late representatives  
438 of *Leptobos* (gr. LEV) make a comparative distinction from the Pietrafitta bovid more difficult. Indeed, the  
439 studied postcranial remains share several morphometric features with this group, and especially with *L.*  
440 *vallisarni* from Italy (see section 5.2; Figs 8, 9). In spite of that, taking in consideration the cranium, the  
441 differences from *Leptobos* gr. LEV are much more evident. The wider occipital, the anteroposterior  
442 compression of the intertemporal bridge, the longer nasals, and the shorter premaxillae clearly differentiate  
443 the Pietrafitta bovine from *L. etruscus*. Most interestingly, the double bending of the horn-cores, their wide  
444 angle of projection from the frontals, and their short and bulk built are characters that separate the Pietrafitta  
445 bovine from both *L. etruscus* and *L. vallisarni*.

446 It is clear that some of the aforementioned features can be interpreted as apomorphies, approaching the  
447 Pietrafitta bovine to *Bison* s.l. Nonetheless, the studied sample still retains several features that are generally  
448 considered primitive in the evolution of *Bison*, including: small and short horn-cores, weakly tubular orbits,  
449 concave frontals, trapezoidal occipital squama, strong indentation of the temporal fossae behind the horn-  
450 cores, and relatively short and slender limb bones.

451 On the whole, the Pietrafitta sample shows numerous similarities with the primitive bison forms  
452 attributed to the subgenus *Eobison*. In order to identify the samples at species level, within this subgenus, we

453 provide below detailed comparisons of its most complete and diagnostic elements with several *Leptobos* and  
454 *Bison* s.l. species (Figs 6–11; Tables 5, 6).

#### 455 4.2.1. *Horn-cores*

456 The best-preserved horn-core (right core of SABAP UMB 19. 2.1179) is plotted in the bivariate diagram  
457 that compares the anteroposterior (HAP) and dorsoventral (HDV) diameters of the core bases between  
458 various samples of the leptobovine and bisontine groups (Fig. 6a). All early *Bison*, on average, display  
459 smaller horns compared to the priscoid group of *B. schoetensacki* and *B. priscus*, with some specimens (e.g.,  
460 Kalamoto and Dmanisi samples) showing even smaller sizes than *Leptobos*. Most of *Eobison*, *Leptobos* gr.  
461 LSEM, and some *B. schoetensacki* specimens are characterized by HAP values considerably higher than  
462 HDV ones. These dorsoventrally compressed cores, are clearly different from the more cylindrical ones of  
463 *Leptobos* gr. LEV and Late Pleistocene-Holocene *Bison*. The Pietrafitta specimen falls within the variability  
464 of *Eobison*, in the area occupied by the Salita di Oriolo, Capena, and Dmanisi samples (Fig. 6a). The position  
465 of the horn-cores respect to the frontals can give some taxonomic hints (McDonald, 1981). The angle formed  
466 by the midline of the core at the base and the sagittal line is wider in *Bison* s.l. and *Leptobos* gr. LSEM than  
467 in *Leptobos* gr. LEV (Fig. 6b). In particular, the horns of *B. priscus*, *B. schoetensacki*, and *B. bison* are  
468 directed more laterally, forming wide angles, whereas *Leptobos* gr. LSEM, *B. menneri*, and *Eobison* show  
469 cores shifted backwards and, in *Leptobos* gr. LEV, even more posteriorly. The two specimens from  
470 Pietrafitta exhibit approximated angles of 65° and 70° fitting with the variability of *Bison*, and more  
471 specifically of *B. (Eobison)*.

#### 472 4.2.2. *Intertemporal bridge and frontals*

473 The anteroposterior compression of the intertemporal bridge represents an important trend in the  
474 evolution of *Leptobos* and *Bison* (Pilgrim, 1947; Masini, 1989; Duvernois, 1990; Cherin et al., 2019). The  
475 primitive forms of *Leptobos* are characterized by a well-developed postcornual portion of the neurocranium  
476 with a high and narrow intertemporal bridge, posteriorly elongated. Conversely, in later forms (i.e., *Leptobos*  
477 gr. LEV), the bridge tends to shorten anteroposteriorly and widen, due to the anterodorsal shifting of the  
478 occipital squama and the squeezing outward of the temporal fossae. In *Bison*, the occipital squama is even  
479 more advanced and the intertemporal bridge disappears, fusing with the nuchal crest. Following Cherin et al.  
480 (2019), in Fig. 7a we show the progressive change in the proportions of the intertemporal bridge from high

481 and narrow in *Leptobos* to short and wide in *Bison*. This latter condition, in which the bridge is fused with  
482 the nuchal crest, is present in the Pietrafitta bovid. Nonetheless, in this form, it is still present a strong  
483 indentation of the temporal fossae behind the horn pedicles and, consequently, a constriction of the upper  
484 occipital squama; a character which is present in later species of *Leptobos* (i.e., *L. vallisarni*) and in *Eobison*  
485 (Flerov, 1972; Toniato et al., 2017; Kostopoulos et al., 2018 among others).

486 Another evolutionary trend of *Leptobos* and *Bison* is the widening of the frontals and the reduction of  
487 their relative length. The biplot in Fig. 7b clearly shows that the earliest representatives of *Leptobos* and  
488 *Bison* (gr. LSEM and *Eobison*) have more elongated and narrower frontals compared with their respective  
489 descendants (gr. LEV and *B. priscus*). The Pietrafitta specimen SABAP\_UMB 19. 2.1178 falls close to the  
490 *Eobison* samples from Dmanisi, Capena, and Pirro Nord (Fig. 7b).

#### 491 4.2.3. Occipital and basioccipital

492 In order to assess the most diagnostic morphometric traits of the occipital and basioccipital in *Leptobos*  
493 and *Bison*, a PCA was performed on 41 crania, using six adjusted variables (Fig. 7d; Table S6). According to  
494 the results, the first two PCs explain the 79% of the total variance. PC1 (67.9% of variance) is influenced by  
495 the distance between the posterior ends of the temporal fossae (IBW positive), by the width of the foramen  
496 magnum and of the basioccipital posterior tuberosities (FMW and PPW both negative), and by the shape of  
497 the occipital squama (OSHmin negative, MW positive). This component divides the specimens having wide  
498 intertemporal bridge and dorsoventrally compressed occipital squama (x positive) from the individuals with  
499 narrow intertemporal bridge, high and narrow occipital squama, and wider foramen magnum and  
500 basioccipital (x negative). PC2 (11.8% of variance) is dominated by the width of the foramen magnum  
501 (FMW negative), the distance between the posterior ends of the temporal fossae (IBW negative), and by the  
502 size of the occipital squama and basioccipital tuberosities (MW, OSHmin and PPW positive). Thus, this  
503 component helps distinguishing the specimens with more developed occipital squama and basioccipital (y  
504 positive) from those with larger foramen magnum and wider intertemporal bridge (y negative). The  
505 Pietrafitta specimen SABAP\_UMB 19. 2.1179 is found in the top right area, close to the centre of the  
506 diagram, fitting with the variability of *Bison* s.l. (Fig. 7d). The bi-condylar width (CW) appears to be the less  
507 influencing variable in both PC1 and PC2.

#### 508 4.2.4. Teeth

509 The comparative study of teeth revealed that the Pietrafitta bovine was characterized by a mixture of  
510 plesiomorphic and apomorphic characters and a wide morphological variability. The degree of hypsodonty in  
511 bovines tends to increase from the middle Villafranchian on, reaching the highest levels in Late Pleistocene  
512 and Holocene forms of *Bison* (Asperen and Kahlke, 2017). The high and columnar teeth of the Pietrafitta  
513 bovine are different from those of early *Leptobos*, characterized by more brachydont dentition. On the  
514 contrary, the distinction with *Leptobos* gr. LEV is less marked due to their higher hypsodonty (Merla, 1949;  
515 Masini 1989; Duvernois, 1990). The amount of cement in the upper molars is another character that may  
516 bear diagnostic value, as it is quite scarce or absent in *Leptobos* and more commonly found in *Bison*. All  
517 upper teeth from Pietrafitta present abundant cement, especially within the folds of the lingual side. The  
518 mesiodistal constriction of the lingual cones of upper molars has been considered as plesiomorphic in  
519 *Leptobos* (Cherin et al., 2019). As a matter of fact, this character is recognizable – albeit variably – in all  
520 analysed *Leptobos* and some *Eobison* samples (e.g., Pirro, Mygdonia Basin), whereas is less common in  
521 more derived species such as *B. schoetensacki* and *B. priscus*. The Pietrafitta sample shows mesiodistal  
522 constriction of the lingual cones in all upper teeth, with various degree of development. The presence of a  
523 bubaline fold in the upper molars, sometimes considered as a primitive feature (e.g., Cherin et al., 2019), is  
524 not a reliable source of taxonomic information, being variably expressed within Bovinae and quite common  
525 in *Bison* s.l. (Merla, 1949; McDonald 1981; Asperen and Kahlke, 2017; Sorbelli et al., 2021a). The p4 has  
526 been considered one of the most diagnostic anatomical elements of *Eobison*. For instance, a large metaconid  
527 with cylindrical shape and lingual projection was considered a distinctive trait in the first description of the  
528 *B. (E.) degiulii* type material by Masini (1989). The latter author, however, suggests that the morphology of  
529 the p4 in this species may show a certain variability, confirmed by our personal analysis of the unpublished  
530 material from Pirro Nord (IGF collections). The Pietrafitta record shows a wide spectrum of morphologies of  
531 the p4 metaconid, from bulky and lingually projected (e.g., SABAP UMB 20. 1.7341) to thin and distally  
532 projected (e.g., SABAP UMB 130074), with a prevalence of the former condition (Fig. 4). Raw  
533 measurements (Table S4) and Log<sub>10</sub> ratio plot (Fig. 7c) show that the dimensions of the teeth from Pietrafitta  
534 are similar to those from Dmanisi and Capena, with short and relatively wide molars, amongst the largest  
535 within the *Eobison* group.

536 On the whole, the teeth from Pietrafitta are characterized by bison-like morphology, with some  
537 noteworthy exceptions. Although over the decades several authors have defined diagnostic characters in the  
538 teeth of *Leptobos* and *Bison* (e.g., Masini, 1989; Duvernois 1990; Cherin et al., 2019), we collided with the  
539 evidence that dental morphology is highly variable in these taxa. Skinner and Kaisen (1947: 140)  
540 exhaustively commented on this issue in their monography on extinct and extant bison “The pronounced  
541 similarity in *Bison* teeth has defied the present attempt to establish useful characters that would permit  
542 specific or even subgeneric separations based on tooth characters alone. Large population samples suggest  
543 tendencies but show no clear-cut differences. Isolated examples of various *Bison* species appear to have  
544 tooth characters that would aid in making specific determinations. These characters, however, soon  
545 intergrade when a large population sample is examined”. Other authors remarked furtherly the wide  
546 spectrum of variability in bovines’ dental traits when dealing with large samples: “Many students who dealt  
547 with very voluminous collections of bovines hold rather sceptical view on the subject, especially concerning  
548 isolated teeth” (Sheer, 1997: 114). The Pietrafitta sample reinforces this vision, allowing only trends to be  
549 recognized, not defined diagnostic characters. For these reasons, we advocate a cautious approach in using  
550 isolated teeth for taxonomic purposes in bovines, especially when small samples are available and/or the  
551 putative presence of more than one taxon is inferred.

#### 552 4.2.5. *Metapodials*

553 Metapodials (Table 6; Table S26) are the most frequently used postcranial bones in the taxonomy of  
554 Bovinae thanks to their abundance in the fossil record and distinct morphology (Schertz, 1936; Sher, 1997;  
555 Maniakas and Kostopoulos 2017; Sorbelli et al., 2021a). Recent studies, however, recommend a prudent  
556 approach due to the considerable degree of intraspecific variability of these elements, which might be  
557 partially caused by ecophenotypic changes (Maniakas and Kostopoulos 2017; Sorbelli et al., 2021a).  
558 Comparisons based on Lmax and DEW are widely used to assess the metapodial robusticity, which tends to  
559 increase on average from *Leptobos* to *Bison* (Sher, 1997; Kostopoulos et al., 2018; Bukhsianidze, 2020;  
560 Sorbelli et al., 2021a). The metacarpal biplot Lmax vs DEW/Lmax% (Fig. 8a) shows the presence of various  
561 morphotypes: the left part of the diagram is occupied by the slenderest forms such as *Leptobos* (especially  
562 the species of the gr. LSEM); the upper part is occupied by the large-sized and long-legged *B. menneri*; on  
563 the right, the more robust species such as *B. schoetensacki* and *B. priscus*, can be found. The Pietrafitta

564 sample is located in the middle of the diagram, within the variability range of late Villafranchian *Bison*. It is  
565 noticeable the partial overlap of the Pietrafitta sample with the most robust specimens of *L. etruscus* and *L.*  
566 *cf. vallisarni* from the Valdarno historical collection. Two specimens (SABAP UMB 130038 and SABAP  
567 UMB 20. 1.7325) are very short and slender, fitting well with *B. (Eobison)* sp. from Venta Micena (Spain),  
568 one metacarpal of *B. (E.) palaeosinensis*, and some of the metacarpals from the Upper Valdarno (Le Ville,  
569 La Viesca) referred to *L. cf. vallisarni*. Four specimens display larger and stouter built, fitting better with *B.*  
570 *(E.) degiulii* from the Mygdonia Basin and *B. (E.) georgicus* from Dmanisi, overlapping with the slenderest  
571 specimens of *B. schoetensacki* from various sites of Europe. The two associated metacarpals belonging to the  
572 only articulated skeleton from Pietrafitta (SABAP UMB 22. 4.247/7 and SABAP UMB 22. 4.247/8), fall at  
573 the left border of the *Eobison* group and overlap with the largest *L. etruscus*. The same plot built for the  
574 metatarsals reveals a similar pattern (Fig. 8b). The largest and stoutest sample, on the right, is mainly  
575 composed by the priscoid forms (*B. schoetensacki* and *B. priscus*). The upper part is occupied by the tallest  
576 specimens (i.e., *B. menneri* and some *B. schoetensacki* individuals), whereas the bottom left is occupied by  
577 *Leptobos*, characterized by slenderer and shorter metatarsals. Again, *Eobison* is found in the middle. The  
578 Pietrafitta record falls at the left border of the *Eobison* range, partially overlapping with *Leptobos*.

579 In order to increase taxonomic resolution, we performed a principal component analysis of seven adjusted  
580 variables, on both metacarpals and metatarsals (Fig. 9). The two principal components (PC1 and PC2)  
581 explain a significant amount of the variance, expressing 80% and 66% of the total variance for metacarpals  
582 and metatarsals, respectively. The metacarpal analysis shows that PC1 (71% variance) is negatively  
583 influenced by Lmax PET and DET, whereas positively by PEW, DW, and DEW. These results separate the  
584 longer and thicker metacarpals (x negative) from the shorter and wider ones (x positive). PC2 (9%) is mainly  
585 dominated by Lmax, DW, and DT, for negative values, whereas on the positive axis by PEW, PET, DEW,  
586 and DET, so that this axis separates longer metacarpals with larger diaphysis (y negative) from shorter ones  
587 with larger epiphyses (y positive). The PCA diagram, although not resolving the overlapping issues between  
588 groups, helps us in distinguishing the three major morphotypes without the influence of size effect: the  
589 *Leptobos* morphotype with slender diaphysis and deep and narrow epiphyses; the “true bison” morphotype  
590 (including *B. schoetensacki*, *B. priscus*, and *B. bonasus*) with robust built and wide diaphysis and epiphyses;

591 the *Eobison* morphotype (including also *B. menneri*) with intermediate features between the two former  
592 groups. The Pietrafitta specimens fall within this last group.

593 In the metatarsal analysis (Fig. 9b; Table S27), PC1 (48% variance) is mainly influenced by Lmax for the  
594 positive values and by DEW, PEW, DW, DT for the negative ones, thus dividing the long and slender  
595 metatarsals (x positive) from the stouter ones with wider shaft and distal end (x negative). PC2 (18%) is  
596 determined, positively, by the size of the distal epiphysis (DET and DEW) and the anteroposterior diameter  
597 of the proximal end (PET); whereas negatively by the width of the diaphysis (DW), which influences way  
598 more than the other variables (PEW, Lmax, and DT). PC2 allows to distinguish the metatarsals with thick  
599 ends (y positive) from the ones with wide shafts (y negative). The Pietrafitta specimens, characterized by  
600 very thick and wide diaphysis, lie along the negative PC2 axis, occupying an empty space, and partially  
601 overlapping the *Eobison* record.

602 Both the PCAs show that the Pietrafitta sample is characterized by metapodials with intermediate  
603 proportions between *Leptobos* and *Bison* s.s., fitting with the known *Eobison* variability. It is worth  
604 mentioning that some isolated bones from the late Villafranchian of Central Italy referred to *L. etruscus* and  
605 *L. cf. vallisarni* (Masini, 1989), display proportions similar to some *Eobison* specimens. It has to be  
606 considered, in fact, that the first *Bison* and last *Leptobos* likely coexisted in Europe during this period. Thus,  
607 it is possible that some of these isolated remains could have been erroneously attributed to *Leptobos* and,  
608 instead, pertain to *Bison*.

#### 609 4.2.6. Radii, astragali, and other postcranials

610 Among the other postcranial bones, the radii are the most helpful. As already shown by Sorbelli et al.  
611 (2021a), their size and built can provide interesting information for taxonomy. The biplot that compares the  
612 length and slenderness degree of this bone shows a clear increase in stoutness and size from *Leptobos* to  
613 *Bison* (Fig. 8d; Table S12). The Pietrafitta specimens plot in the space occupied by the early *Bison* samples  
614 from Pirro Nord, Capena, and the Mygdonia Basin at the upper border of the *Leptobos* cloud. Other  
615 postcranials such as calcanei and cubonavicular are not particularly significant due to their considerable  
616 dimensional homogeneity between the last *Leptobos* and early *Bison* (Tables S21, S25). Conversely, it is  
617 worth noting that the astragali allow a certain separation between “true” *Bison* and *Eobison-Leptobos* based  
618 on total length and, secondarily, stoutness (Fig. 8d; Table S23). The low number of complete humeri,

619 femurs, and tibiae does not allow us to perform dimensional comparisons. In spite of that, the average  
620 measurements of these elements in the Pietrafitta sample (and in other roughly coeval samples) are always  
621 intermediate between those available for *Leptobos* and *Bison* s.s. (Tables S10, S17, S19).

#### 622 4.2.7. Limb proportions and concluding remarks

623 Interesting results can also be obtained from the dimensional comparison between different bones. The  
624 analysis of the ratio between the metatarsal and metacarpal lengths shows that the latter bone is relatively  
625 longer in most primitive forms than in the priscoïd group (Fig. 9c). Indeed, in *Leptobos* and most *Eobison*,  
626 including the Pietrafitta sample, the metatarsals are less than 15% longer than the metacarpals, whereas in  
627 Late Pleistocene *Bison* from Eurasia and North America (including the living species *B. bonasus* and *B.*  
628 *bison*), it is more than 20% longer (Fig. 9c). On the contrary, the ratio between radius and metacarpal lengths  
629 shows that the Pietrafitta bovine and other records of *Eobison* share with *Bison* s.s. a relatively elongated  
630 radius compared to the metacarpal (Fig. 9c). This analysis reveals that in all late Villafranchian bison the  
631 radius length exceeds of more than 30% the metacarpal length, with the exception of the samples from  
632 Dmanisi and Venta Micena (Fig. 9c). Interestingly, the earliest members of *Bison* s.s. (i.e., *B. menneri*, *B.*  
633 *schoetensacki*, *B. p. priscus*) differ significantly from their later relatives (i.e., *B. p. mediator*, *B. antiquus*, *B.*  
634 *latifrons*, *B. bison*, *B. bonasus*) in having longer metacarpals compared to both metatarsals and radii.

635 Considering all the above pieces of evidence, the bovid sample from Pietrafitta can be confidently  
636 referred to the subgenus *Eobison*. More specifically, most crania and limb bone morphological and biometric  
637 characters are consistent with those described in *B. (E.) degiulii*, a well-known species from the late  
638 Villafranchian of Southern Europe (Italy and Greece) (Fig. 10). Another species with similar morphology is  
639 *B. (E.) georgicus* from Georgia. However, it is represented by a rather scarce sample, which needs to be  
640 enriched by new discoveries in order to clarify its taxonomic status (see Discussion).

641

## 642 5. Discussion

### 643 5.1. *Leptobos*: state of the art

644 During the twentieth century, several diagnoses of *Leptobos* have been proposed (Pilgrim, 1939; Merla,  
645 1949; Masini, 1989; Duvernois, 1990). Unfortunately, this multiplicity of interpretations has led to a rather  
646 vague definition of this genus. Nevertheless, some distinctive characters of *Leptobos* are accepted by the

647 different authors, including: horn-core bases located not far from the orbits, emerging posterolaterally,  
648 slightly to strongly compressed dorsoventrally; horn-cores furrowed in the first half, without evident keel;  
649 female hornless; premaxilla in contact with nasal; presence of ethmoidal fenestrae; temporal fossae well  
650 developed in the postcornual portion of the cranium; sub-triangular basioccipital; occipital squama from  
651 semi-circular to bell-shaped; presence of a more or less developed intertemporal bridge; slender limbs.  
652 Below we provide a review of *Leptobos* records in the European Villafranchian.

653 The first group of species (here called gr. LSEM), generally considered more primitive, includes *L.*  
654 *stenometopon*, *L. elatus*, *L. merlai*, and the poorly-known *L. furtivus*. *Leptobos stenometopon* (Fig. 12b) is a  
655 small- to mid-sized species described on the basis of a horned calvarium from the late Pliocene of Dusino  
656 (Italy; Rüttimeyer, 1867). The taxonomy of this species has been debated over the last decades, being  
657 considered either a younger synonym of *L. elatus* (Duvernois and Guérin, 1989; Duvernois, 1990, 1992) or a  
658 valid species (Masini, 1989; Geraads, 1992; Rodrigo, 2011; Masini et al., 2013; Cherin et al., 2019). It is also  
659 reported - as *L. cf. stenometopon* - from the Plio-Pleistocene transition in Montopoli (Italy) and from some  
660 sites of uncertain age in the Upper Valdarno (Merla, 1949; Masini, 1989). *Leptobos elatus* is a species  
661 described by Depéret (1884) on cranial material from the latest Pliocene of Perrier-Les Étouaires (France).  
662 Unfortunately, the type material is no longer available (Viret, 1954; Masini, 1989) and the neo-diagnosis  
663 given by Duvernois (1990, 1992) was based only on dental and postcranial remains and on the original  
664 descriptions by Depéret (1884). The synonymy between *L. elatus* and *L. stenometopon* proposed by  
665 Duvernois and Guérin (1989) and Duvernois (1990) has been questioned by other authors (e.g., Masini et al.,  
666 2013; Cherin et al., 2019) due to some flaws in the taxonomic status of *L. elatus* and to significant  
667 differences between the material from the type localities of the two species. Other remains attributed to *L.*  
668 *elatus* are from the sites of Rocaneyra (France; dated to the Plio-Pleistocene boundary), Villarroya (Spain;  
669 ca. 2.15 Ma), and Tegelen (The Netherlands; ca. 2.0 Ma) (Schreuder, 1946; Viret, 1954; Masini, 1989;  
670 Duvernois, 1990; Rodrigo, 2011). *Leptobos merlai* (Fig. 12c) is a mid- to large-sized species from the middle  
671 Villafranchian of France (type locality: Saint Vallier) and Italy. First referred to *L. elatus* (Viret 1949) then to  
672 *L. stenometopon* (Viret, 1954), De Giuli (1987) recognized diagnostic differences between the remains from  
673 Dusino and those from Saint Vallier and erected the species *L. merlai* to accommodate the latter. Later,  
674 Duvernois and Guérin (1989) and Duvernois (1990, 1992) similarly to what was done for *L. stenometopon*,

675 proposed *L. merlai* as a subspecies of *L. elatus*. Finally, in their description of the material from Pantalla (ca.  
676 2.0 Ma) and other Italian sites, Cherin et al. (2019) supported the validity of *L. merlai* as a distinct species  
677 (and questioned the validity of *L. elatus* itself, being based on scanty material). *Leptobos furtivus* is  
678 considered the smallest and actually rarest species of this genus in Europe. It was first described by  
679 Duvernois and Guérin (1989) on the basis of few teeth, some postcranials, and an isolated horn-core from the  
680 middle Villafranchian of Senèze (France), where also the larger *L. etruscus* is recorded. Isolated remains  
681 from the late Villafranchian of Italy are also referred to *L. furtivus* (Masini, 1989; Duvernois, 1990). An  
682 incomplete cranium and some postcranials from Piana del Cavaliere (Italy) may be attributed to the same or  
683 a similar species (*L. aff. furtivus* in Masini, 1989; Gentili and Masini, 2005; Fabbi et al., 2021). In spite of  
684 the above records, *L. furtivus* remains a little-known species and conclusions on its own validity can only be  
685 granted by future discoveries (Cherin et al., 2019).

686 The second group of species (gr. LEV in this manuscript) includes *L. etruscus*, *L. vallisarni*, and the  
687 poorly-known *L. bravardi*. *Leptobos etruscus* (Fig. 12e) - erected by Falconer (1868) based on the rich  
688 collection from the Upper Valdarno - is the best-known species of the genus, occurring in most Europe  
689 during the first stages of late Villafranchian. Fossils of *L. etruscus* have been found in several sites of Spain,  
690 France, Italy, Greece, Romania, and possibly the United Kingdom (Merla, 1949; Masini, 1989; Duvernois  
691 1990; Kostopoulos, 2006; Garrido, 2008, Breda et al., 2010; Croitor and Popescu 2011; Kostopoulos, 2022  
692 and references therein). This species has been considered closely related with *L. vallisarni* based on  
693 numerous craniodental similarities (Masini, 1989), to the point that Duvernois (1990) interpreted *L.*  
694 *vallisarni* as a junior synonym of *L. etruscus*. *Leptobos vallisarni* (Fig. 12f) - described by Merla (1949)  
695 based on a partial cranium from the late Villafranchian of the Upper Valdarno - is characterized by some  
696 apomorphic cranial features shared with the earliest bison. Unfortunately, very little is known on the  
697 postcranial anatomy of *L. vallisarni* due to absence of sites where associated cranial and postcranial remains  
698 were found. Several isolated metapodials from Central Italy were tentatively referred to this species by  
699 Masini (1989); however, due to the unclear stratigraphic context (e.g., Upper Valdarno indeterminate  
700 localities) and their affinities with the roughly coeval *L. etruscus* and *Eobison* spp., we advocate a prudent  
701 approach to their attribution. Two almost complete skulls from the Gonghe Basin (China) were attributed to  
702 *L. vallisarni* (Zheng et al., 1985), unfortunately not associated with postcranial bones, again. *Leptobos*

703 *bravardi* is a mid- to large-sized species described by Duvernois (1990) from Perrier-Les Étouaires, where  
704 also *L. elatus* is reported. The same author referred to this species also some of the remains from Villarroya  
705 (Spain), where, again, the co-occurrence of *L. elatus* is hypothesized. The validity of *L. bravardi* has been  
706 questioned by Rodrigo et al. (2010) and Rodrigo (2011) who attributed the variability recognized in the  
707 Spanish and French samples to sexual dimorphism. The fragmented neurocranium of *L. bravardi* from  
708 Perrier-Les Étouaires (Duvernois, 1990) was referred to *L. elatus* by Masini (1989).

709

## 710 5.2. *Leptobos*: comments on the European taxa

711 The crania of LSEM are easily distinguishable from those of LEV and *Bison* s.l. thanks to the peculiar  
712 morphology of the cornual and postcornual areas. The double curvature of the horns (anteroposterior and  
713 dorsoventral) differentiates these species from *L. etruscus* and *L. vallisarni*. The large size and slenderness of  
714 the horns distinguish them from primitive bison, whereas the dorsoventrally compression of the core bases  
715 (Fig. 6a) and the strong core curvature distinguish them from *Bison* s.s. The divergence angle between the  
716 horn-core axis and the sagittal axis is higher in LSEM than in LEV and fits with the variability of *Bison* s.l.  
717 (Fig 6b). The temporal fossae of LSEM are wide and deep, especially in the posterior portion, causing a  
718 constriction of the dorsal part in the occipital squama. In addition, the ventral portion of the latter is wider  
719 due to massive mastoid processes, resulting in a bell shape of the squama itself (Cherin et al., 2019). These  
720 *Leptobos* species, due to the posterior elongation of the temporal fossae, display a distinctly high and narrow  
721 intertemporal bridge, which clearly separates the nuchal crest from the frontals. This element, as mentioned  
722 above, tends to reduce as a consequence of the progressive shortening of the postcornual portion of the  
723 neurocranium in the evolution of *Leptobos* and *Bison*, and disappears in the most derived forms fusing with  
724 the nuchal crest (Fig. 7a). LSEM is also distinguished by markedly narrow and elongated frontals (Fig. 7b).  
725 Our analyses point out that, on the whole, the cranial morphology of *L. stenometopon*, *L. elatus*, and *L.*  
726 *merlai* is similar to a degree that makes it difficult their distinction at species rank (see PCA in Fig. 7d). The  
727 horn-core morphology is very useful also for the taxonomic recognition of LEV. In these species, the horn-  
728 core base section is almost perfectly circular (Fig. 6a), and the horn-cores emerge in a more posterior  
729 direction than the other bovines considered thus the angle between the middle line of the horn and the  
730 sagittal plane is quite reduced (Fig 6b). Moreover, the horn-cores are characterized by simple curvature

731 describing a half-moon shape and lie parallel to the frontal plane, clearly different from both LSEM and  
732 *Bison* s.l. In addition, in LEV the intertemporal bridge is much lower and wider than in LSEM; this character  
733 is quite evident in *L. vallisarni*, in which the shape of the intertemporal bridge is similar to what is observed  
734 in early *Bison* (Fig. 7a). The occipital squama of LEV is semicircular due to the increasing distance between  
735 the posterior edges of the temporal fossae and the shortening of the intertemporal bridge. Finally, there is a  
736 trend in the widening of the frontals and a general increase of stoutness and size of the cranium from LSEM  
737 to LEV as shown by our biometric data and analyses (Fig. 7; Table 5). Other cranial characters (apart from  
738 the teeth) are similar between the two *Leptobos* groups (e.g., presence of ethmoidal fenestrae, premaxillae in  
739 touch with the nasals, absence of horns in females). The teeth of *Leptobos* are characterized by strong  
740 polymorphism, and a relative homogeneity in size (Fig. 7c). Nonetheless, some characters can be useful to  
741 discriminate the two groups, provided that sufficiently large samples are available. These characters include  
742 the hypsodonty degree, molarization of P4, amount of enamel folds in the central cavities, and cement  
743 abundance (see Duvernois, 1990; Cherin et al., 2019).

744 The postcranials of LSEM show small size and overall light built compared with those of LEV and *Bison*  
745 s.l. Body mass estimations based on metacarpals give an average value slightly higher than 400 kg for  
746 LSEM, that is, the smallest in the studied sample (Fig. 11b). In our analyses, the metapodials and radii are  
747 characterized by narrow epiphyses and relatively elongated and slender diaphysis (Figs 8, 9; Table 6; Table  
748 S12). Comparing the lengths of different long bones, it emerges that the radius is relatively shorter than the  
749 metacarpal, while the latter is relatively longer than the metatarsal (Fig. 9c). The humeri, astragali, and  
750 calcanei are among the slenderest of all the studied records (Tables S10, S23, S25). Again, the samples  
751 referred to gr. LSEM do not present clear-cut biometric differences in the postcranials, except for the  
752 metacarpal sample from Saint Vallier (*L. merlai*), which appears slightly stouter than the others (Fig. 8a).  
753 The homogeneity in long bone size and proportions between *L. stenometopon*, *L. elatus*, and *L. merlai* might  
754 suggest that they actually represent a single species. From this perspective, the small differences in stoutness  
755 between the various records might be correlative with local ecomorphotypes rather than interspecific  
756 variation.

757 On the contrary, LEV features larger and stouter bodies compared with early *Leptobos*, and more similar  
758 to *Eobison* (Tables S10, S12, S17, S19, S21, S23, S25, S26). The average body mass of LEV is about 500 kg

759 (Fig. 11b). The radius is characterized by gracile structure and small size (Fig. 8c). The metacarpals are long  
760 and slender, but still stouter than those of LSEM (Figs 8a, 9a, 11a). Nonetheless, the low mean value for the  
761 length ratio RAD/MC% indicates that the radii are relatively shorter than the forelimb, i.e., the same as in  
762 LSEM (Fig. 9c). The low values of the ratio MT/MC% clearly separate *L. etruscus* from *Bison* s.l. (Fig. 9c).  
763 The astragali are small compared with those of *Bison* s.l. and overlap in size those of LSEM (Fig. 8d). The  
764 lack of postcranial material undoubtedly referable to *L. vallisarni* makes difficult the comparison between  
765 this species and its relatives. In fact, our biometric analyses of metapodials show a significant overlapping of  
766 the remains putatively attributed to *L. vallisarni* and those of earliest *Bison*. This issue suggests that isolated  
767 bones might have been misinterpreted and their attributions must be taken cautiously (e.g., sample from Le  
768 Ville in the Upper Valdarno, which probably belongs to a primitive form of bison).

769

### 770 5.3. *Eobison* and other “transitional” forms: state of the art

771 Several species have been included into *Eobison* since the establishment of this subgenus by Flerov  
772 (1979). Unfortunately, over the decades, *Eobison* has become a wastebasket taxon for all bison-like bovid  
773 remains found in the Eurasian Villafranchian, hampering its clear definition. The diagnosis provided by  
774 Flerov (1979) is rather insufficient, also due to the scarce fossil record available at that time. For this reason,  
775 we propose the emended diagnosis in Section 4.

776 Up to date, the oldest bison fossils belong to *B. (E.) sivalensis*. This species was erected by Lydekker  
777 (1878) on the basis of a partial skull from the Upper Siwaliks (Pakistan); later, Pilgrim (1939) published the  
778 diagnosis based on the same cranial material. Since then, few more cranial elements from the Indian Plio-  
779 Pleistocene were tentatively attributed to *B. sivalensis*, all chronologically constrained between 3.3 and 2.6  
780 Ma, making thus, this species the first representative of the *Bison* s.l. clade (Khan et al., 2010, 2011). In spite  
781 of that, since the holotype skull has apparently been lost (Olsen, 1990) and additional material is too  
782 fragmentary for clear attributions, there are no remains that can undoubtedly represent *B. (E.) sivalensis*.  
783 Olsen (1990) and Bukhsianidze (2020) even questioned referral of the type material to the genus *Bison* based  
784 on a number of morphological characters.

785 *Bison (E.) palaeosinensis* (Fig. 12g) is a small-sized primitive bison occurring in China shortly after *B.*  
786 *(E.) sivalensis*. The taxonomic status of this species has been debated for a long time. *Bison (E.)*

787 *palaeosinensis* was erected by Teilhard de Chardin and Piveteau (1930) on the basis of three partial skulls - a  
788 male (cranium A), a female (cranium B), and a third without sex attribution - and some limb bones from the  
789 Nihowan (=Nihewan) Basin. The authors described the characters of this species as “intermediate” between  
790 *Leptobos* and *Bison*, sharing more similarities with the latter. Two other incomplete skulls from Yushe and  
791 Jinyuan cave in Dalian were referred to this taxon (Teilhard de Chardin and Trassaert, 1938; Jin et al., 2021).  
792 Skinner and Kaisen (1947) designated the cranium A as lectotype, evidencing the similarities between  
793 cranium B and the holotype of *B. (E.) sivalensis*, thus not excluding a synonymy between the two species.  
794 McDonald (1981) attributed the cranium A to *B. priscus* and the cranium B to *B. (E.) sivalensis*, hence  
795 invalidating the species. Tong et al. (2016) redefined the taxon and described a large number of dentognathic  
796 and postcranial remains from the site of Shanshenmiaozui in the Nihowan Basin. These authors also  
797 remarked that the cranium B shared some important similarities with *Leptobos*. Up to date, this taxon  
798 represents the best-known species of *Eobison* in Asia.

799 *Bison (E.) georgicus* (Fig. 12h) is a mid-sized, primitive species, originally defined by Burchak-  
800 Abramovich and Vekua (1994) as *Dmanisibos georgicus*, based on material (including a fragmented and  
801 reconstructed cranium) from the renowned late Villafranchian site of Dmanisi. Bukhsianidze (2005) revised  
802 the material and included it into *Bison (Eobison)*. The limb bone (in particular metapodials) sample is  
803 heterogeneous, showing the presence of two different morphotypes, possibly referable to another taxon,  
804 suggesting that *Eobison* from Dmanisi could have been sympatric with a more primitive bovid, possibly  
805 *Leptobos*, as hypothesized in other regions of Eurasia (Tong et al., 2016; Kostopoulos et al., 2018; Sorbelli et  
806 al., 2021a).

807 *Bison (E.) degiulii* (Fig. 12i, j, k) is, up to date, the best-known species of early *Bison*. It was described  
808 informally by Masini (1989) in his PhD dissertation and, later, validated by Masini et al. (2013), based on  
809 material from Pirro Nord (latest Villafranchian). The holotype is a partial cranium, associated with several  
810 postcranials. An almost complete skeleton from Capena was also referred to this species (Masini, 1989;  
811 Masini et al., 2013). Masini (1989) and Sorbelli et al. (2021a) attributed several large bovid remains from  
812 Sainzelles (France) to *Eobison cf. degiulii*. Kostopoulos et al. (2018) described the rich record from  
813 Mygdonia as *B. (Bison) cf. degiulii*, choosing the open nomenclature due to the lack of clear-cut diagnostic  
814 characters in this Greek sample, which seems to display a large number of derived features. The authors,

815 indeed, stressed that the morphological features of this species can be considered closer to *B. (Bison)* than to  
816 *B. (Eobison)*, in agreement with Bukhsianidze (2005), referring this species to the former subgenus. The  
817 timespan covered by the aforementioned sites is comprised between 1.7 and 1.2 Ma, thus suggesting that  
818 primitive bison, at least in the Mediterranean area, survived until the beginning of the Epivillafranchian and,  
819 potentially, co-existed with the larger and more derivate “prisoid” forms of *Bison* s.s., such as *B. (B.)*  
820 *schoetensacki*, whose first occurrence is at ca. 1.2 Ma (Kostopoulos et al., 2018; Sorbelli et al., 2021a).

821 *Bison (E.) tamanensis* was first mentioned, without a diagnosis, by Verestchagin (1959) on scanty cranial  
822 material from Tzimbal in the Taman Peninsula (Russia). Later on, the species was validated as member of  
823 *Eobison* by Flerov (1979) who gave a diagnosis and designated a lectotype (a partial calvarium). In more  
824 recent years, a better-preserved skull was found in Semibalki (Taman Peninsula) and attributed to *B. cf.*  
825 *tamanensis* (Baigusheva, 2000). Isolated metapodials and a fragment of horn-core from the three Ukrainian  
826 sites of Kairy, Chortkiv, and Cherevychnyy were also referred to this species (David and Svistun, 1981). At  
827 present, our knowledge of *B. (E.) tamanensis* is poor and the taxonomic validity of this species is still  
828 discussed (*species inquirenda* in Kostopoulos et al., 2018).

829 The species *Eobison suchovi* was erected by Alekseeva (1967) based on a single horn-core and few  
830 postcranial bones from Dolinskoye (Ukraine). Due to the quite scanty record, the taxonomic validity of this  
831 species has been questioned and it has been often synonymized with *B. tamanensis* (Flerov, 1975; Gromov  
832 and Baranova, 1981; Croitor, 2016).

833 *Adjiderebos cantabilis* (Fig. 12d) was described by Dubrovo and Burchak-Abramovich (1986) on the  
834 basis of a nearly complete skull from the middle-late Villafranchian of Kushkuna (Azerbaijan; Bukhsianidze  
835 and Koiava, 2018; Krijgsman et al., 2019). Bukhsianidze (2005) suggested that this species might belong to  
836 an early form of *Eobison*. No postcranials are known to date. From older layers of the same locality, scanty  
837 remains of another bovid were described by Burchak-Abramovich et al. (1980) and referred to *Protobison*  
838 *kushkunensis*. This latter taxon, dated to the late early Villafranchian, could be the first “bisontine” form  
839 outside Asia (Bukhsianidze and Koiava, 2018). This type material of this species although poorly preserved,  
840 share some plesiomorphic characters with LSEM (Bukhsianidze, 2005).

841 *Probison dehmi* is a small bovid from the Tatrots of the Upper Siwaliks (Pakistan) described by Shani  
842 and Khan (1968). The holotype is a badly preserved partial cranium with one horn-core, showing common  
843 traits with the Asian *Leptobos falconeri* and *B. (E.) sivalensis* (Shani and Khan, 1968; Kahn et al., 2010).

844 Over the decades, several other samples of bison-like bovids were described from late Villafranchian sites  
845 of Europe and referred to as “early *Bison*” or “*Eobison*”, but without a specific attribution. The largest of  
846 these samples is from Venta Micena. This site, dated between 1.6 and 1.5 Ma (Palmqvist et al., 2022),  
847 yielded a large number of bovid taxa (Moyà Solà, 1987). This last author recognized, in the larger-sized  
848 form, some bison-like characters and attributed the remains to *Bison* sp. Later on, the study of a partial  
849 juvenile calvarium confirmed that a relatively small and primitive bison was present in Venta Micena during  
850 the late Villafranchian (Martinez-Navarro et al., 2011). The postcranial bones from the site are slightly  
851 stouter compared with those of *Leptobos* but shorter and more gracile than those of *Bison* s.s., resembling  
852 those of *Eobison* although characterized by relatively smaller size (Masini, 1989; Sorbelli et al., 2021a).

853 A well-preserved neurocranium with both horn-cores from Salita di Oriolo (Italy) was referred to *Bison*  
854 (*Eobison*) sp. by Toniato et al. (2017) due to primitive characters in the occipital and temporal. To *Bison*  
855 (*Eobison*) sp. were attributed other scanty records from the Italian peninsula, including an isolated horn-core  
856 from Mugello and few postcranial bones from Cava Redicicoli (Italy) (Masini, 1989; Caloi and Palombo,  
857 1995). Bonnoure and Dimitrijevic (2006) referred to an undefined primitive bison several fragmentary  
858 remains from the lower layers of Trlica cave (Montenegro). A skull fragment from the Early Pleistocene of  
859 Maar du Riege (France) may belong to the same group (Ambert et al., 1996). Finally, from the late  
860 Villafranchian site of Taurida cave (Crimea, Ukraine), the putative co-occurrence of *Leptobos* and *Eobison* is  
861 reported (Loopatin et al., 2019), but no further information of the fossils is available. Due to the scarcity of  
862 data, the last three records are not further discussed and were not included in our comparisons.

863

#### 864 5.4. *Eobison* and other “transitional” forms: comments

##### 865 5.4.1. *Dmanisi*

866 The Dmanisi bovid has been considered the first bison reaching Europe in the first part of the late  
867 Villafranchian. The few remains attributed to *B. (E.) georgicus* include a neurocranium and some  
868 postcranials, although more elements are pending to be published (Bukhsianidze, 2020). The neurocranium

869 (MG D354) is broken at the level of the orbits, posteriorly to the nasals. The occipital and basioccipital  
870 regions are well preserved on the whole. The right horn-core and the base of the left are present, although the  
871 reconstruction of their original orientation is not trustworthy (Bukhsianidze, pers. comm.). In agreement with  
872 Bukhsianidze (2005), the neurocranium from Dmanisi shows many features supporting its attribution to  
873 *Eobison*, including: trapezoidal occipital squama, high temporal fossae strongly indenting within the  
874 postcornual area, weakly tubular orbits, arched nuchal crest in dorsal view, long pedicles and posterolaterally  
875 directed horns (angle of 55° with the sagittal axis). Our biometric comparisons of cranial districts reveal that  
876 the Dmanisi bovid is well-fitting with the variability of *Bison* s.l., and in particular with the most primitive  
877 forms (Figs 7, 8). The postcranial elements evidence a certain heterogeneity, which might suggest the  
878 presence of two different taxa in Dmanisi (Kostopoulos et al., 2018; Sorbelli et al., 2021a), as advocated in  
879 Taurida cave (Lopatin et al., 2019) and Tsiotra Vryssi (Kostopoulos et al., 2018). The three most complete  
880 metacarpals from Dmanisi have been re-measured and the analyses show that two of them (MG D2812 and  
881 MG D3426) have proportions compatible with *Eobison* and one (MG D2288) features slenderer proportions,  
882 fitting with the variability of the largest *Leptobos* (Fig. 8a). The two complete radii from Dmanisi (MG  
883 D2165 and MG D2962) have short and stout appearance, recalling robust forms of *Eobison* and differing  
884 from the smaller radii of *Leptobos* spp. and Venta Micena *Bison* (Fig. 8c). In our opinion, at least the two  
885 stouter metacarpals and the radii can be confidently referred to *Eobison* and not to *Leptobos*. The same goes  
886 for the cranium MG D354, which shows *Eobison* features. The more pronounced concavity of the forehead  
887 and the stronger constriction of the upper occipital squama, are traits that distinguish the Dmanisi specimen  
888 from the younger remains of *B. (E.) degiulii* from Pirro Nord and Kalamoto (Figs 6, 7, 11; Table 5).  
889 However, as far as our current knowledge is concerned, it is not possible to attest if these differences can be  
890 due by inter- on intraspecific variability (e.g., sexual dimorphism, which we know to play an important role  
891 in determining the cranial morphology in bison). Pending new discoveries, we consider *B. (E.) georgicus* as  
892 a distinct taxon. Although we cannot exclude the presence of both *Leptobos* and *Eobison* in Dmanisi, new  
893 fossils and more evidence are needed to attest the coexistence of these two genera in the Georgian site. It is  
894 worth remembering that, in Southern Europe, in the interval corresponding to the age of the Dmanisi  
895 assemblage (ca. 1.8 Ma; early late Villafranchian), *Leptobos* was the only bovine present. It will be replaced  
896 by *Eobison* only after ca. 1.6 Ma (latest Villafranchian).

#### 897 5.4.2. *Venta Micena*

898 The postcranials *Venta Micena* represents one of the richest records of late Villafranchian *Bison*. Our  
899 analyses of metapodials show that this bovine was relatively small and had quite short metacarpals (few  
900 specimens exceed 230 mm in length) and elongated metatarsals (most of the specimens exceed 260 mm in  
901 length), which is a common feature of *Bison* species (Fig. 8a). The robusticity biplots and PCAs show that  
902 the metapodial record from *Venta Micena* falls within the variability of *Eobison*, mostly in the area occupied  
903 by the smallest and slenderest specimens. As stated by other authors (Moyà-Solà, 1987; Palmqvist et al.,  
904 2022), the wide biometric variability of the sample is the result of marked sexual dimorphism (Sorbelli et al.,  
905 2021a). The overall small size of the specimens from *Venta Micena* (estimated average body mass of 440 kg;  
906 Fig. 11b) might be related to the fact that most fossils are part of a biogenic concentration interpreted as a  
907 *Pachycrocuta brevirostris* den (Arribas and Palmqvist, 1998; Palmqvist et al., 2022). It has been  
908 hypothesized that this fossil assemblage is strongly biased in favour of females and young individuals,  
909 characterized by small sizes and, thus, more commonly hunted by predators like hyaenas (Arribas and  
910 Palmqvist, 1998). On the other hand, the three complete radii from the site display an oddly primitive  
911 proportions with reduced lengths (Figs 8c, 9c); indeed, the value for RAD/MC% is similar to those found in  
912 *Leptobos*. Nonetheless, the lack of an abundant sample of complete radii might affect these results, as does  
913 the aforementioned sex ratio bias in metacarpals. The juvenile cranium of *Bison* sp. MAEGR VM 8000  
914 described by Martínez-Navarro et al. (2011) is strongly damaged and quite uninformative. Even so, the  
915 estimated diameters of the horn-core bases are within the variability of *Eobison* (Fig. 6a). Its reduced size is  
916 most probably due to the early ontogenetic stage. Waiting for a comprehensive reappraisal of the *Venta*  
917 *Micena* sample, we agree with previous scholars in referring it to the subgenus *Eobison* without further  
918 conclusions on its specific attribution. The possible co-occurrence between *Leptobos* and *Bison* in *Venta*  
919 *Micena* (Sorbelli et al., 2021a) might be suggested by the presence of two metacarpals and one metatarsal  
920 with a slightly elongated morphology, fitting with the variability of gr. LEV. However, in both biplots and  
921 PCAs, the larger *Leptobos* and *Eobison* are similar each other, making it difficult to differentiate them in the  
922 overlapping areas (Figs 8a-b, 9a-b).

#### 923 5.4.3. *Pirro Nord and Mygdonia Basin*

924 The vertebrate assemblage from the karstic site of Pirro Nord is dated, on a biochronological basis, to  
925 1.6–1.3 Ma, although some authors lean towards the upper part of this interval (ca. 1.4–1.3 Ma; Arzarello et  
926 al., 2015 and references within). The holotype of *B. (E.) degiulii* (IGPB s.n., temporarily hosted in DST), a  
927 partial cranium of an elder individual, is broken just anteriorly to the orbits. Only horn-core pedicles are  
928 partially preserved. It clearly shows bison-like features such as enlarged occipital, laterally placed horn-  
929 cores, and protruding orbits. In the first descriptions by De Giuli et al. (1986) and Masini (1989), the authors  
930 emphasized the primitive aspect of the cranium, fitting with the original definition of *Eobison* by Flerov  
931 (1979). On the contrary, Bukhsianidze (2005) and Kostopoulos et al. (2018) suggested that the Pirro Nord  
932 bovine already shows a derived morphology, referable to “true” *Bison*. In fact, the deep and narrow  
933 temporal fossae and doomed frontals are apomorphic character observed in the Pirro Nord cranium.  
934 Nonetheless, the section of the horn-core base and shape of the occipital squama, the size of the teeth, and  
935 the proportions of the limb bones are still primitive (Figs 10, 12; Table 5). The intertemporal bridge biplot  
936 (Fig. 7a) shows that the Pirro Nord specimen exhibits a marked dorsal constriction of the occipital squama,  
937 resembling the *Eobison* remains from Dmanisi and Pietrafitta and *L. vallisarni* from the Upper Valdarno.  
938 The biplots on frontal and horn-core dimensions highlight the same similarities (Figs 7, 8). Teeth from Pirro  
939 Nord are characterized by quite small size (Fig. 7c). Their morphology is variable, as already mentioned by  
940 Masini (1989). The Pirro Nord bovine displays, on average, more robust metacarpals than the slightly older  
941 primitive bison from the Iberian and Italian Peninsulas (Figs 8a, 11a). The metatarsals, on the contrary, are  
942 less stout and fit better with the smaller specimens of *Eobison* (Fig. 8b). These discrepancies in metapodial  
943 proportions, however, might be caused by the strong sexual dimorphism of these bovids, thus to a possible  
944 sex bias in the Pirro Nord record. The length ratio between metatarsal and metacarpal points out that the  
945 former is less than 15% longer than the latter, fitting with the variability of *Eobison* and some *Bison* s.s.  
946 species (Fig. 9c). Taking in consideration that the younger age estimations for the Pirro Nord assemblage are  
947 around 1.3 Ma, it is possible that this record represents one of the last and most derived forms of *Eobison*,  
948 roughly coeval with the stout *B. (E.) cf. degiulii* from Apollonia. The latter locality lies in the Mygdonia  
949 Basin as well as the sites of Tsiotra Vryssi, Krimni, and Kalamoto-2, with an overall age spanning from 1.78  
950 to 1.2 Ma (Kostopoulos et al., 2018; Konidaris et al., 2021). Among the most interesting remains of *B. (E.)*  
951 *cf. degiulii* from this area, a fragmentary cranium and several complete limb bones including metapodials

952 and radii stand out. The cranium (NHCK KLT-638; Fig. 10f) from Kalamoto-2, lacks the horn-cores and  
953 most of the splanchnocranium, being broken at the level of the P3. Most of its features, including the  
954 morphology of the orbits, occipital, and frontals, are similar to those of the holotype of *B. (E.) degiulii* from  
955 Pirro Nord (Kostopoulos et al., 2018). The horn-core bases from Mygdonia sites are among the smallest and  
956 dorsoventrally compressed among those examined, fitting with other *Eobison* samples and with *B. menneri*  
957 (Fig. 6a). The intertemporal bridge is not preserved; however, the occipital squama shows a constriction in  
958 its upper section due to the posterior elongation of the temporal fossae, as in all *Eobison*. The robusticity  
959 biplot performed on the metapodials shows that, compared with the other records of *Eobison*, the Mygdonia  
960 sample (especially the Apollonia specimens), is characterized by marked stoutness, which causes a partial  
961 overlap with the records of *B. schoetensacki* (Fig. 8a). In spite of that, PCAs confirm that the Mygdonia  
962 bovine clearly falls within the variability of early bison (Fig. 9a). In the metrical analysis, the four complete  
963 radii from Apollonia fall at the upper limit of the *Eobison* range, reinforcing the idea that this record can be  
964 referred to a short-legged, but still relatively slender form of *Bison*, different from the larger and stouter *B.*  
965 *schoetensacki* and *B. priscus* from the Epivillafranchian onwards. This is also supported by length ratios  
966 between metacarpals, metatarsals, and radii (Fig. 9c; Table 6; Table S12, S26).

967 In sum, the records of *B. (E.) degiulii* from Pirro Nord and the Mygdonia Basin clearly represent a  
968 derived stage of this group, marked by a greater expression of bisontine characters and larger body size (i.e.,  
969 estimated body mass of ca. 600 kg, on average; Fig. 11b). In particular, the metacarpals from Apollonia (ca.  
970 1.2 Ma) are the most robust among all primitive bison, setting a turning point, at the late Villafranchian-  
971 Epivillafranchian transition, towards increased forelimb stoutness. This might be related to an aridification of  
972 the habitat, already inferred for the locality (Maniakas and Kostopoulos, 2017; Kostopoulos et al., 2018).  
973 Based on our review, we agree with Bukhsianidze (2005) and Kostopoulos et al. (2018) that *B. (E.) degiulii*  
974 is not completely fitting with the original diagnosis of *Eobison* by Flerov (1979). However, we have  
975 provided compelling evidence that this species is markedly different from *Bison* s.s. and conforms to our  
976 emended diagnosis of the subgenus.

#### 977 5.4.4. Capena, Salita di Oriolo, and other Italian sites

978 The bovid skeleton (MUST (ex MPUR) s.n.) recovered from Capena (late Early Pleistocene) in 1970, has  
979 been recently restored allowing its study in this paper (Fig. 10b1-3). Masini (1989), in his preliminary

980 description of the material, recognized clear bisontine features and attributes this specimen to the newly  
981 erected species *B. (E.) degiulii* due to some primitive traits of the appendicular skeleton. The skeleton is  
982 almost complete and in good state of preservation, although the skull is severely cracked and deformed. In  
983 spite of that, the horn-cores are almost intact; they are large-sized and emerge laterally, dorsoventrally  
984 flattened at the base, and with single curvature (directed downward at the base then upward at the tips). The  
985 biplot in Fig. 6a shows that this specimen has the most compressed and large horns in the *Eobison* group  
986 (although dorsoventral compression might be slightly exaggerated by taphonomic deformation). The  
987 postcornual portion is shortened and the occipital region is wide and flattened (Fig. 10b2). The left  
988 premaxilla is partially intact, and its posterior margin does not reach the nasal; the latter has a bifid anterior  
989 edge (Fig. 10b1). The rest of the cranium is severely damaged thus providing poor information. Teeth do not  
990 differ significantly from those of other bovines. The bubaline fold is present in all upper molars, which  
991 present faint constriction of the lingual lobes. The cement is abundant, especially on the labial side,  
992 concentrated in the recess between the protocone and metacone. The entostyle of the upper molars has a  
993 relatively simple outline. The p4 is characterized by a large and cylindrical metaconid. In the Log<sub>10</sub> diagrams,  
994 dental dimensions follow the general trend of *Eobison*. Based on the size of the appendicular skeleton, the  
995 Capena bovine stands out for being particularly large (estimated body mass of about 600 kg; Fig. 11b). The  
996 metapodials and radii are among the stoutest and largest in *Eobison*, always placed in the upper range of  
997 variability of this group in both biplots and PCAs, very close to *Bison* s.s. (Figs 8, 9). Moreover, the  
998 metatarsal and radius are elongated with respect to the metacarpal, recalling the typical condition of more  
999 derived species of *Bison*. Consequently, the size and proportions of this skeleton may lead to the question of  
1000 whether it belongs to a relatively large and derived species *Eobison*, such as the specimens from Pirro Nord  
1001 and Apollonia, or a to small-sized form of true *Bison* (e.g., female of *B. schoetensacki*). The lack of any other  
1002 fossil from Capena does not help dispel this doubt. Even if we are more inclined to agree with Masini (1989)  
1003 and Masini et al. (2013) in referring this specimen to *B. (E.) degiulii*, here we prefer cautiously to use the  
1004 open nomenclature *B. (E.) cf. degiulii*.

1005 The calvarium of *Eobison* (MCSNF s.n.) from the “Sabbie Gialle” formation of Salita di Oriolo (Fig.  
1006 10c), dated to the late Early Pleistocene (Toniato et al., 2017), was found associated with a fragment of  
1007 mandible, some upper teeth, and a cervical vertebra. The fossil is strongly weathered but not fractured or

1008 deformed except for the left jugular process and some portions of the occipital which are missing (Fig. 10c1-  
1009 4). The dorsal edge of the right orbit is still preserved as well as the basioccipital region. The horn-cores are  
1010 virtually intact. The occipital is characterized by a large but narrow squama with a marked bell-shape. The  
1011 intertemporal bridge is wide and dorsoventrally compressed. The frontals are wide and slightly inflated  
1012 between the horn-core bases; as a result, the forehead is concave except for a small crest running along the  
1013 interfrontal suture. The orbits are weakly tubular. The pedicles are relatively long. The horn-cores are  
1014 massive and relatively short, without an evident keel; they emerge in posterolateral direction (ca. 72° with  
1015 the sagittal axis; Fig. 6b), bend laterally after the first third, and point posteriorly toward the tips; in posterior  
1016 view posterior, they are directed downward and then curve markedly upward. The horn-core base is wide,  
1017 with strong dorsoventral compression (Fig. 6a). The teeth are heavily worn, ma. This, affected also their  
1018 morphology, making their description difficult (e.g., the metaconid, entoconid, and entostylid of p4 are  
1019 fused). However, it is possible to recognize constricted lingual lobes in the upper molars and a bubaline fold  
1020 in the P4. To sum up, the general morphology of the material from Salita di Oriolo is primitive (see, in  
1021 particular, the strong constriction of the upper occipital squama, the high intertemporal fossae, the concave  
1022 frontals, and the overall small size). Similarities with *B. (E.) degiulii*, *B. (E.) palaeosinensis*, and *B. (E.)*  
1023 *georgicus*, especially in the occipital region, are striking (Figs 10, 12; Table 5). In our PCA of neurocranial  
1024 variables, the Salita di Oriolo specimen falls within the range of variation of *Bison* s.l. (Fig. 7d), but it stands  
1025 out as one of the smallest known crania of *Eobison* (Fig. 10). Based on this evidence, we support the  
1026 attribution of the Salita di Oriolo sample to a primitive form of *Bison*, with strong affinities with *B. (E.)*  
1027 *degiulii*. The age of the fossils is debated (Toniato et al., 2017). Considering the similarities between the  
1028 bovine material and *B. (E.) degiulii*, we tend to consider the “Sabbie Gialle” formation of late Villafranchian  
1029 age, that is, slightly older than previously estimated (ca. 1.0 Ma; Toniato et al., 2017).

1030 An isolated right horn-core (MGCB s.n.) from Mugello, although damaged and badly restored, appears to  
1031 preserve its original morphology. It is relatively short and has a wide base. In dorsal view, it slightly curves  
1032 posteriorly, then anteriorly, then again posteriorly, toward the tip. In anterior view, it bends downward and  
1033 then markedly upward. The base is dorsoventrally compressed. Furrows are mainly concentrated on the  
1034 posteroventral portion. The shape and size (Fig. 6a) are well fitting with the variability of primitive bison, as  
1035 already suggested by Masini (1989). For these reasons, we confirm referral of this specimen to *B. (Eobison)*

1036 sp. This specimen might be the same coming from the lignite mine of Lumena (Mugello) and referred to  
1037 *Leptobos* sp. by Merla (1949).

1038 From Cava Redicicoli, dated around 0.8 Ma, the co-occurrence of the large *B. schoetensacki* and the  
1039 smaller *B. (Eobison) cf. degiulii* is reported, due to the presence of few metapodials of different sizes (Caloi  
1040 and Palombo, 1995; Masini et al., 2013). The two complete metacarpals, however, show a feeble discrepancy  
1041 in size and proportions that might be explainable by sexual dimorphism. In the PCA (Fig. 8a), they fall  
1042 within the variability of larger *Eobison* and smaller *B. schoetensacki*. Even if the Cava Redicicoli sample is  
1043 surely referable to *Bison* s.l., and most probably to a single species, further studies are needed to assess their  
1044 specific taxonomy.

1045 Several metapodials from the Upper Valdarno area were ascribed to *Leptobos cf. vallisarni* and *L. cf.*  
1046 *etruscus* by Masini (1989) due to their large size and relatively stout appearance. Our comprehensive  
1047 comparative analysis reveals that some of these fossils might be attributed to *Eobison*. Of the four  
1048 metacarpals from Le Ville previously referred to *L. cf. vallisarni*, two (IGF 3279, IGF 2374) show different  
1049 preservation and are characterized by shorter diaphysis and stouter structure, fitting better with the variability  
1050 shown by the *Eobison* group. The same goes for the metacarpal IGF VA661 from La Viesca (Figs 8a, 9a).  
1051 For these three specimens we propose a new identification under *B. (Eobison) sp.*

#### 1052 5.4.5. Eastern Europe and Asia

1053 The (lost) holotype of *B. (E.) sivalensis* is a partial skull missing of the horn-cores and part of the  
1054 splanchnocranium. In our revision of the species original definition, we found an incongruence in the horn-  
1055 core description by Lydekker (1878). The author states: “The horn-cores are compressed antero-posteriorly;  
1056 their anterior surface is flat from above downwards and concave from within outwards; their posterior  
1057 surface is convex in both directions” (Lydekker, 1878: 23). However, in the drawing of the specimen  
1058 (Lydekker, 1878: pl. XV) the figured section of the horn is clearly dorsoventrally compressed and, in the  
1059 measurement given by the same author (Lydekker, 1878: 125–138), the anteroposterior diameter is 3.4  
1060 inches (=86.4 mm) whereas the transverse (dorsoventral?) diameter is 2.6 inches (=66 mm), suggesting,  
1061 again, a dorsoventral compression of the horn-core. This issue led to a series of misleading conceptions of  
1062 this species. For example, Khan et al. (2010) attributed to *B. (E.) cf. sivalensis* two complete horn-cores from  
1063 late Pliocene deposits of the Tatrot formation, distinguishing from *L. etruscus* for their anteroposterior

1064 compression, absence of torsion, and shorter length. Taking in consideration that the latter two characters are  
1065 not diagnostic in differencing this sample from *L. etruscus* (the length of the horn-cores actually fit with the  
1066 variability of *L. etruscus*, which in addition, shows only a feeble torsion), and, if it is true that the holotype of  
1067 *B. (E.) sivalensis* had dorsoventrally compressed horn-core bases, then these horn-cores from the Tatrot more  
1068 resemble the sub-cylindrical sections typical of *L. etruscus* and *Bison* s.s. (Fig. 6a). Based on literature data  
1069 and available figures, the shape of the occipital area of *B. (E.) sivalensis*, with its markedly indenting  
1070 temporal fossae, recalls the typical condition of primitive bison. Olsen (1990) and Bukhsianidze (2020),  
1071 following the idea of Lydekker (1878), suggested that this species could belong to a different *Bison* subgenus  
1072 and attributed it to cf. *Poepagus*, in light of similarities with the wild yak. Although we agree that several  
1073 aspects (e.g., morphology, chronological and geographical ranges) may lead to the same conclusion, the lack  
1074 of available material does not allow us to infer further on the taxonomy of this form.

1075 *Bison (E.) palaeosinensis* is considered one of the most primitive forms of bison, unfortunately suffering  
1076 from a confused taxonomic definition, only partially solved by the emended diagnosis recently given by  
1077 Tong et al. (2016). One of the main concerns lies in the differences between the two most complete syntypic  
1078 specimens, cranium A and cranium B, originally attributed to a male and female, respectively (Teilhard de  
1079 Chardin and Piveteau, 1930). We were not able to study directly the type material, but the measurements and  
1080 figures available in the literature, clearly show that cranium B is actually different from cranium A,  
1081 somewhat resembling *L. vallisarni* in various traits of the horn-cores and occipital. However, we still know  
1082 too little about the intraspecific variability of primitive bison to rule out that the differences between the two  
1083 crania may actually be due to sexual dimorphism (Tong et al., 2016). The idea by McDonald (1981) that  
1084 cranium A belongs to a small form of *B. priscus* is debatable due to the clearly primitive features of the  
1085 occipital squama and temporal fossae, as well as its age. Some similarities especially in the shape and  
1086 orientation of the horn-cores, can be observed between cranium B and the holotype *B. (E.) sivalensis* figured  
1087 by Lydekker (1878). However, the unavailability of the type material of the latter species imposes caution in  
1088 taxonomic conclusions. Biometric comparisons support affinities of both cranium A and cranium B with  
1089 other records of *Eobison* (Table 5). The same goes for cranial morphology, which shows an array of  
1090 primitive characters including flattened to concave frontals marked by a sagittal crest along the suture, long  
1091 pedicles, short postcornual section, high temporal fossae. Moreover, cranium A shares with *L. vallisarni* and

1092 *Eobison* the intertemporal bridge proportions (Fig. 7a), again disproving its possible inclusion in *B. priscus*.  
1093 The postcranials referred to *B. (E.) palaeosinensis* are characterized by a relatively slender built and small  
1094 size, fitting with the largest *Leptobos* specimens and most *Eobison* (Figs 6, 7; Table 6; Tables S12, S19, S23,  
1095 S25, S26).

1096 The holotype of *B. (E.) tamanensis* is a portion of left horn-core still attached to fragments of occipital,  
1097 temporal, and frontal bones (ZIN RAS 26010). Although the postorbital constriction is visible, it seems to be  
1098 less developed than in most primitive forms of *Bison*. The temporal fossa is slightly indenting the occipital  
1099 area, ending not far from the horn-core base, which has a relatively short pedicle. The horn-core is  
1100 dorsoventrally flattened and has a faint keel on the posterior side. As shown in the biplot in Fig. 6a, the  
1101 specimen is quite large compared to other records of *Eobison*, falling not far from the area occupied by *B.*  
1102 *schoetensacki* from Süssenborn and Mauer. David and Svistun (1981) assigned to this species three isolated  
1103 metacarpals from different Ukrainian localities (Cherevichny, Chertkov, and Kairy) but, as already suggested  
1104 by Sher (1997), there is strong heterogeneity in the sample and their attribution to *B. (E.) tamanensis* is  
1105 mainly based on the pre-Tiraspolian age of the deposits in which they were found. Croitor (2010) referred  
1106 the metacarpals from Kairy and Chertkov to *B. cf. menneri* and the one from Cherevichny to *Eobison* sp. The  
1107 robusticity biplot (Fig. 8a) shows that the Chertkov specimen is, in fact, similar to the largest *B. menneri*  
1108 individuals, whereas the Kairy one fits better with the variability of *B. schoetensacki*, and the slenderest  
1109 specimen from Cherevichny falls in the overlapping range between *Eobison* and *B. menneri*. Therefore, as  
1110 for the first two specimens, we agree with the identification by Croitor (2010), while we prefer to attribute  
1111 the metacarpal from Kairy to *B. cf. schoetensacki* due to its robustness. The poor preservation of the  
1112 holotype of *B. (E.) tamanensis* prevents to define clear-cut diagnostic characters for this species which,  
1113 however, seems to be overall similar to *B. schoetensacki*. The presence of a more complete cranium from  
1114 Semibalki attributed to *B. (E.) cf. tamanensis* (Baigusheva et al., 2000) could help in assessing the validity of  
1115 this taxon. Kostopoulos et al. (2018) noticed some similarities between the cranium from Semibalki and that  
1116 of *B. schoetensacki* suggesting, again, a possible synonymy between the two species. If confirmed, this may  
1117 reinforce the idea that large and stout priscaid *Bison* had already arrived at the gates of Europe at the end of  
1118 the Villafranchian and reached the Mediterranean area during the Epivillafranchian (e.g., Le Vallonnet and  
1119 Vallparadís Section; Sorbelli et al., 2021a).

1120 *Bison (E.) suchovi* is the less known species of *Eobison*. The holotype, described by Alekseeva (1967), is  
1121 a distal fragment of juvenile female horn-core from Dolinskoye (GIN RAS 391-2) characterized by stout  
1122 built and subtriangular section with flattened base. This scanty material does not allow to assess the validity  
1123 of this species. Croitor (2010) suggested that the alleged differences between this horn-core and the holotype  
1124 of *B. (E.) tamanensis* are due to sexual dimorphism, hence considering *B. (E.) suchovi* as a junior synonym  
1125 of *B. (E.) tamanensis*. Sher (1997) stated that the metapodials from Dolinskoye are similar to those from  
1126 Süssenborn, Mauer, and Mosbach, which, at least for the first two localities, belong to *B. schoetensacki*  
1127 (Sorbelli et al., 2021a). Based on these pieces of evidence, *B. (E.) suchovi* is an invalid species in our  
1128 opinion.

1129 The almost complete cranium of *Adjiderebos cantabilis* (PIN 3723-1) from Kushkuna is one of the best  
1130 preserved and yet under-considered specimens of Villafranchian large bovid in Eurasia. The cranium misses  
1131 a portion of the occipital area and fragments of the splanchnocranium, but is in a very good state of  
1132 preservation on the whole. The first descriptions of this outstanding fossil were given by Dubrovo and  
1133 Burchak-Abramovich (1983, 1986), who erected this new genus of bovine evidencing its similarities with  
1134 *Bison* and *Leptobos*. In fact, this cranium shows several features shared with these two genera, but also some  
1135 autapomorphies including: zygomatic arc elevated over the mandibular articulation and unciform rims at the  
1136 external borders of the occipital condyles (Dubrovo and Burchak-Abramovich 1986; Bukhsianidze, 2005).  
1137 Both these characters, however, can be found in other genera of bovines. The zygomatic elevation at the  
1138 mandibular articulation is clearly visible in some specimens of *Leptobos*, *Bos*, and *Bison* s.l. (e.g., the female  
1139 *L. etruscus* NHMB VA605 from the Upper Valdarno and the holotype of *B. (E.) degiulii*) and the hooked  
1140 process of the occipital condyles is well known in *Bison* s.s. (McDonald, 1981: 187, pl. 26). Our  
1141 morphometric analyses show that the horn-core bases share the same proportions with the smallest *Eobison*  
1142 specimens whereas the postcornual constriction is quite developed with a marked posterior indentation of the  
1143 temporal fossae, which resembles the condition found in the most primitive *Leptobos* forms (Fig. 7a). In  
1144 addition, the frontals are extremely narrow (Fig. 7b). The well-developed intertemporal bridge, the  
1145 premaxilla in contact with the nasal, and the presence of ethmoidal fenestrae are all leptobovine features. On  
1146 the other side, the nasals ending posteriorly to the anterior border of the orbits and the wide angle of horn-  
1147 core insertion are bisontine features. On the whole, the characters that would distinguish *Adjiderebos* from

1148 *Leptobos* and *Eobison* seem to be not diagnostic, thus, most probably, this species actually belongs to one of  
1149 these two genera. A direct revision of the Kushkuna material, paired with a phylogenetic study of these  
1150 clades, are requested in order to clarify the taxonomic position of this extremely interesting “transitional”  
1151 form between *Leptobos* and *Eobison*.

1152 The holotype of *Probison dehmi* is a fragmented cranium from the Tatrots stage of the Upper Siwaliks  
1153 (DGPU B/7), dated to the late Pliocene (Shani and Khan, 1968). The cranium is severely damaged and badly  
1154 restored. Nevertheless, it shows interesting characters in common with gr. LSEM including: long frontals,  
1155 shallow and high temporal fossae, strong constriction of the upper occipital squama, developed mastoid  
1156 processes, premaxilla in contact with nasal, presence of ethmoidal vacuity. Shani and Khan (1968) stated that  
1157 the only complete horn-core is anteroposteriorly compressed; in fact, observing the holotype pictures, it is  
1158 clear that the horn-core is markedly compressed in dorsoventral direction. It shows a peculiar orientation,  
1159 bending downward and forward. According to Bukhsianidze (2005), it is most probable that the horn-core  
1160 was mounted on reverse (the base is completely hidden by plaster). This hypothesis is supported by the fact  
1161 that the well-marked ribs and furrows visible on the convex surface of the horn-core, are normally present on  
1162 the ventral side and not on the dorsal. If the horn-core was actually mounted incorrectly, it would originally  
1163 emerge posterolaterally, then bend upward and forward with a strong posterior twist as in gr. LSEM lineage.  
1164 A revision and new restoration of the material is required to assess the validity of this taxon or the possible  
1165 incorporation of *P. dehmi* into *Leptobos*.

#### 1166

#### 1167 5.5. Comments on the evolutive trends in limb proportions and body mass

1168 Several studies on the morpho-functional adaptations of bovid limb bones in response to  
1169 climate/environmental changes over the Pleistocene, have been published in the last decades (Scott, 1985;  
1170 Plummer and Bishop, 1994; Klein et al., 2013; Scott and Barr, 2014; Maniakas and Kostopoulos, 2018;  
1171 Etienne et al., 2021; among others). The comprehensive study by Scott (1979) on the proportions of bovid  
1172 limbs showed that long-legged small-sized taxa typical of open habitats rely on their cursorial skills to outrun  
1173 the predators, whereas in other taxa, passive defence strategies based on large body size and/or gregarious  
1174 behaviour are preferred. Consequently, small-sized cursorial forms underwent to an elongation of the distal  
1175 limb elements, while similar evolutionary changes are not expected in large bovids living in the same open

1176 environments (Scott, 1979, 1985). On the contrary, in these forms, it is observed a marked increase in limbs  
1177 width against length, most probably as an allometric response to heavier bodies (Biewener, 2005; Etienne et  
1178 al., 2020). This is indirectly confirmed by regression equations for body mass prediction that, if performed  
1179 on the length of distal bones, overestimate the mass of small bovids of open/arid habitats (i.e., Antilopini,  
1180 Alcelaphini) and largely underestimate the mass of larger bovines, whereas the epiphysis measurements give  
1181 much more reliable estimations (Scott, 1983). In addition, it has been proved that bovids exceeding 300 kg in  
1182 body mass are characterized by an allometric pattern of the anterior zygopodium (Scott, 1985; Etienne et al.,  
1183 2020), in the fact that this section underwent to a stronger increase of size and robusticity, compared to the  
1184 other forelimb sections, as a response to heavier mass and subsequent change in posture. In their analysis of  
1185 several samples of *Bison* from the Pleistocene of Europe, Sorbelli et al. (2021a) suggested that the robusticity  
1186 of metacarpals in bison is inversely related to the amount of tree cover in the environment.

1187 The present study, including data from numerous species of *Leptobos* and *Bison* covering a wide  
1188 chronological range, supports the above observations and gives further hints on the morphological response  
1189 of these bovines to environmental changes over the last 2 Ma in Western Palearctic. *Leptobos* (especially  
1190 LSEM) is considered as typical of relatively closed environments (Rook and Martinez-Navarro, 2010). In  
1191 fact, it is inferred that one of the main causes of the extinction of these forms was the progressive contraction  
1192 of wooded areas started at the end of the Pliocene and specially during the late Villafranchian (Head and  
1193 Gibbard, 2005, 2015; Maslin and Brierley, 2015). *Bison*, due to its capability to adapt to a wider spectrum of  
1194 environments, occupied the new available niches, in particular after the onset of new asymmetric  
1195 glacial/interglacial cycles and the increase of seasonality during the EMPT (Maslin and Brierley, 2015). Our  
1196 results show a clear trend towards an increased robusticity of metapodials from the Pliocene onwards (Fig.  
1197 11a). The violin boxplot of body mass estimations (Fig. 11b) highlights a constant increase of average  
1198 weight starting from the late Pliocene and reaching a peak during the late Middle Pleistocene, before  
1199 recording an opposite trend in the Late Pleistocene and Holocene (see also Maniakas & Kostopoulos 2017).  
1200 Similarly, the biplot of radius robusticity (Fig. 8d) and the RAD/MC% histogram (Fig. 9c) show that  
1201 *Leptobos* and *Bison* underwent to a clear increase in size and stoutness of the anterior zygopodium correlated  
1202 with larger body masses (Fig. 11b).

1203 It is even possible to observe the same parallel trends within *Leptobos* and *Bison* at smaller scale. In  
1204 *Leptobos*, there is a transition from the slender and short metapodials of LSEM to the stouter and longer ones  
1205 of LEV, paired with an increase of radius size, body mass, and hypsodonty, following the first contraction of  
1206 forested environments at the middle-late Villafranchian transition (Rook and Martinez-Navarro, 2010;  
1207 Madurell-Malapeira et al., 2014). In *Bison*, due to their long history, wide geographic distribution, and  
1208 abundant fossil record, the situation is much more complex. The primitive *Eobison*, roughly coeval to the last  
1209 *Leptobos*, features increased body mass and robusticity, characterized by larger and heavier heads and limbs,  
1210 although not yet displaying the massive appearance of later forms. Their limb structure shows a wide  
1211 spectrum of variability (e.g., slender morphology from Pietrafitta, stouter from Apollonia), which could  
1212 reflect either separate taxa or ecophenotypic variations. It is clear, however, that close to the beginning of the  
1213 Epivillafranchian, their size and stoutness started to increase, especially in the most arid environments such  
1214 as those of Pirro Nord and Apollonia. An even wider variability is present in *Bison* s.s. due to the extremely  
1215 widespread geographic distribution of these derived forms. It is true indeed, that if both *Leptobos* and  
1216 *Eobison* are mainly (if not exclusively) found between the 45th and 35th parallels, later forms of *Bison* (i.e.,  
1217 priscoid group) populated all the Palearctic, from Siberia to the Mediterranean area (Fig. 1a) (Kahlke, 1999).  
1218 This wide distribution led to the evolution of several morphotypes adapted to the dominant habitats and  
1219 climatic conditions, in a timespan that covers more than 1 Ma. The very slender *B. menneri* and the stouter *B.*  
1220 *schoetensacki* were the most common forms in Europe during the Early-Middle Pleistocene boundary (Sala,  
1221 1986; Moullé, 1992; Brugal, 1995; Sher, 1997; Sorbelli et al., 2021a). Although characterized by large size  
1222 and priscoid characters, they still featured slender built compared with the extremely stout *B. priscus*. As  
1223 already pointed out by Sher (1997) and Sorbelli et al. (2021a), the common misconception that *B.*  
1224 *schoetensacki* was a small-sized forest-adapted form, is disproved by the records from Le Vallonnet and  
1225 Vallparadís Section, which prove that these animals were adapted to a wider spectrum of habitats and very  
1226 large body mass, even reaching 1,000 kg (Fig. 11b). The first occurrence of *B. priscus* is still debated.  
1227 Between 0.9 and 0.5 Ma, bison with priscoid morphologies started to appear in fossil records such as those  
1228 from the Kuznetsk Basin, Tiraspol, and Mosbach, which are, however, still pending a clear taxonomic status  
1229 (Freudenberg; 1914; Sher, 1997; Foronova, 2001; Sorbelli et al., 2021a). In the late Middle Pleistocene, giant  
1230 and massive forms such as *B. p. gigas* and *B. p. priscus* from Riverenert, Châtillon-Saint-Jean, Romain-La-

1231 Roche, and Taubach, are recorded (body mass reaching 1,600 kg; Fig. 11b) (Mourer-Chauvire, 1972; Flerov,  
1232 1977a, 1977b; Vercoutère and Guerin, 2010). Some of these animals, however, were still characterized by  
1233 relatively elongated metacarpals (Fig. 11a). From the Late Pleistocene on, the body mass and metapodial  
1234 length of *B. priscus* (often referred to the subspecies *B. p. mediator*; Kahlke, 1999) started to decrease. In  
1235 agreement with the aforementioned model by Sorbelli et al. (2021a), the slenderer forms of *Bison* s.s. (e.g.,  
1236 the long-legged *B. menneri* from Untermassfeld or the samples of *B. schoetensacki* from Le Vallonnet and  
1237 Mauer) are found in more humid, heterogeneous habitats (i.e., mosaic of forested and open patches), while  
1238 the samples of *B. schoetensacki* from more arid sites such as Süssenborn, Durfort, and Vallparadís Section  
1239 exhibit stouter proportions (Fig. 11a). Similarly, the tallest and slenderest specimens of *B. priscus* are those  
1240 from Taubach, a site characterized by a fully developed interglacial conditions with warm and forested  
1241 environment (Kahlke, 1977), whereas the stouter members of this species inhabited open, steppe/prairie-like  
1242 habitats during the last glacial both in Eurasia and North America (Fig. 11a). The common misconception  
1243 that the smallest and stoutest forms of *B. priscus* living during MIS 3 and MIS 2 (*B. p. mediator* sensu  
1244 Flerov, 1979) were typical of forested areas, was already questioned with reason by Sala (1986), Guérin and  
1245 Valli (2000), and Castañós (2012). These authors emphasized how *B. p. mediator* is always associated with  
1246 horses and reindeer (e.g., in Kiputz IX), suggesting that the typical habitat of this subspecies was the cold  
1247 steppe of the last glaciation tundra. Our robusticity analysis confirms this idea, evidencing that *B. p.*  
1248 *mediator* was well adapted to open environments thanks to its robust and stout limbs (Fig. 11a). This  
1249 correlation is maintained also in Pleistocene *Eobison* from the Mediterranean area and in Holocene species.  
1250 *Bison* (*E.*) *degiulii* from Pietrafitta lived in the surroundings of a humid and wooded habitat (see Section 2)  
1251 and featured slender metapodials (slightly stouter than those of *L. etruscus*); conversely, members of the  
1252 same species from Pirro Nord and Apollonia, adapted to drier or rocky environments, bear more robust  
1253 limbs, similar to those shown by the stouter *B. schoetensacki* (Maniakas and Kostopoulos, 2017; Blain et al.,  
1254 2019). The extant subspecies *B. bison bison* populates the open grasslands of North America and is  
1255 characterized by the shortest metacarpals among living bison, whereas the taller and slenderer woodland  
1256 bison *B. bison athabascae* and European wisent *B. bonasus* prefer sub-hygic and mixed habitats  
1257 (McDonald, 1982; Stephenson et al., 2001).

1258 These multiple pieces of evidence are furtherly supported by the dietary habits inferred for extinct and  
1259 extant taxa by Asperen and Kahlke (2017). Based on tooth wear analysis, these authors showed that the  
1260 extremely tall *Bison* from Taubach and Untermassfeld, as well as *B. b. athabascae* are better defined as  
1261 browser-mixed feeders, whereas the stouter *Bison* from Süßenborn and extant *B. bonasus* feature more  
1262 grazer signatures. The North American plain bison *B. b. bison*, namely a species with extremely short  
1263 metacarpals, shows typical dental wear features of an open habitat grazer.

1264 It has been already remarked by several scholars that the mean size of *Bison*, except for few exceptions,  
1265 decreased after the last interglacial in the Holarctic region (Vasiliev, 2008; Saarinen et al., 2016; Maniakas  
1266 and Kostopoulos, 2017). Our body mass estimations computed on a large sample of *Leptobos* and *Bison*  
1267 confirm this statement, showing a clear trend toward gigantism from *Leptobos* to *Eobison* and to *Bison* s.s.  
1268 and an opposite trend from the Late Pleistocene on (Fig. 11b). In particular, during the latest Middle  
1269 Pleistocene (e.g., MIS 6–MIS 5e), the largest forms are recorded (e.g., samples from Châtillon-Saint-Jean,  
1270 Romain-La-Roche, and Taubach). The main driving factors of increasing body size throughout the Early-  
1271 Middle Pleistocene, are still matter of debate. Maniakas and Kostopoulos (2017) observed that bison  
1272 specimens from Joint Mitnor Cave mammal assemblage were significantly smaller than the roughly  
1273 isochronous ones from Taubach (ca. 0.12 Ma). This apparently contrasts with the Bergmann’s rule, which  
1274 expects a positive correlation between body size and latitude. Meiri et al. (2007) pointed out that in taxa with  
1275 wide geographical ranges (such as *Bison*), body size can be influenced by a complex array of factors,  
1276 including not only ecogeographical rules but also - often predominantly - food availability and interspecific  
1277 competition. During the MIS 5e interglacial, the presence of numerous geothermal springs at Taubach would  
1278 have favoured the development of a humid dense forest with a large amount of food (Asperen and Kahlke,  
1279 2017), while the paleoenvironmental conditions of the British islands would have been harsher. Therefore,  
1280 the Taubach bison may have reached gigantic size thanks to extremely favourable trophic conditions, while  
1281 the Joint Mitnor Cave bison may have kept smaller dimensions due to the combination of severe  
1282 environmental context and geographic isolation. A recent study on *B. priscus* inhabiting the Yukon territory  
1283 in the last 0.05 Ma demonstrated that they experienced body mass increase and population growth during  
1284 warmer periods and opposite trends in harsher phases (Abigail et al., 2021). In their comprehensive analysis  
1285 of *Bison* s.s. records from the Middle-Late Pleistocene of Germany and UK, Saarinen et al. (2016) showed

1286 that populations living in open arid landscapes had smaller body size than those living in wooded humid  
1287 landscapes. The same pattern was found in other large herbivores such as caballine horses, in which body  
1288 mass increased in relatively closed, highly productive environments as a result of the combined effect of  
1289 resource availability, habitat partitioning, inter- and intraspecific competition, and herd density (Asperen et  
1290 al., 2010; Saarinen et al., 2021).

1291 Our study provides further evidence to support the above inferences. The progressive harshening of  
1292 climate during the Last Glacial Maximum did not allow the development of very large body size in *Bison*,  
1293 which in most Late Pleistocene sites (e.g., Roter Berg, Habarra, Gral), are characterized by mean body mass  
1294 lower than 1,000 kg, in contrast with the considerably heavier weights reached during MIS 5e and previous  
1295 interglacial stages (e.g., Romain-La-Roche, Taubach) (Fig. 11b). Going further back in time, the Pietrafitta  
1296 sample fits perfectly into this model. The primitive bison from this locality was characterized by relatively  
1297 slender limbs but large body on the whole (i.e., slightly heavier than 600 kg; Fig. 11b). Slender limbs and  
1298 large body size can be related with a warm climate environment rich in forested patches, as suggested by  
1299 palaeobotanical and paleoherpetological evidence (Martinetto et al., 2014; Sorbelli et al., 2021b). The  
1300 excellent ecological suitability would have allowed the Pietrafitta bison to develop larger dimensions than,  
1301 for example, the almost contemporary Venta Micena bison, which lived in a more arid environment and with  
1302 greater competition in the herbivore guild.

## 1303 **6. Conclusions**

1304 Pietrafitta (ca. 1.5 Ma) is one of the key sites of the European late late Villafranchian thanks to the  
1305 extremely rich paleozoological and paleobotanical evidence (Martinetto et al., 2014; Sorbelli et al., 2021b).  
1306 The abundant large-sized bovid record from Pietrafitta, which was preliminarily believed to represent an  
1307 advanced form of *Leptobos* (Masini, 1989; Gentili and Masini, 2005), is here referred to *B. (E.) degiulii*  
1308 based on a combination of characters found in several skeletal districts (e.g., cranium and metapodials).  
1309 *Bison (E.) degiulii* represents one of the earliest representatives of its genus that appeared in Western  
1310 Palearctic.

1311 The study of the Pietrafitta material triggered the reappraisal of all available records of Villafranchian  
1312 *Leptobos* and *Bison*, with a focus on the subgenus *Eobison*. Our revision of the taxonomy and variability of  
1313 this group of early bison has resulted in the emended diagnosis provided in this paper, which has served as a

1314 basis for the summary of comparative characters in Table 7. As far as the available data are concerned, we  
1315 think that among all the species of *Eobison* described so far, only three can be considered valid, namely *B.*  
1316 (*E.*) *palaeosinensis*, *B. (E.) georgicus*, and *B. (E.) degiulii*. Among these, the best-known species - also  
1317 thanks to the very rich sample from Pietrafitta - is *B. (E.) degiulii*, which occurred in Mediterranean Europe  
1318 during the latest Villafranchian. Additional fossils would be desirable to enrich our knowledge of *B. (E.)*  
1319 *palaeosinensis* and *B. (E.) georgicus* which, thanks to their older age and geographical distribution, could  
1320 clarify further aspects on the evolutionary transition between *Leptobos* and *Bison*.

1321 In agreement with previous works, our study also confirms the validity of limb bone dimensions and  
1322 proportions as (1) reliable taxonomic tools, (2) proxies for the estimation of body mass, and (3) predictors of  
1323 paleoenvironmental preferences in bovines. Thanks to these data, we suggest that European *Leptobos* and  
1324 *Bison* responded to the progressive Pleistocene climatic deterioration through changes in limb proportions  
1325 and average body mass. In particular, the contraction of forested habitats and the spread of open  
1326 environments under colder and more arid climate, led to an increase in metapodial stoutness and in body  
1327 size, and to higher hypsodonty associated with grazing-based diet. This trend can be along the transition  
1328 from *Leptobos* to *Bison*, but also on a smaller scale, within the two groups.

1329 Despite the average tendency to increase in size, a more detailed analysis of the *Bison* s.l. record in the  
1330 interval 1.8–0.01 Ma (Fig. 11) shows that there have been fluctuations. In our opinion, it is probable that  
1331 these may be attributable to punctual habitat changes and/or shifts between glacial and interglacial stages,  
1332 but often the paleoenvironmental data and dating available for the various paleontological sites do not have  
1333 sufficient resolution to verify this hypothesis. It should not be forgotten that these phenomena of variation in  
1334 body size and proportions were determined by a complex combination of factors, including ecogeographical  
1335 patterns (e.g., Cope-Depéret and Bergmann's rules), resource availability, intraspecific competition, and  
1336 predatory pressure.

1337

### 1338 **Acknowledgments**

1339 This work is funded by the Agencia Estatal de Investigación European Regional Development Fund of  
1340 the European Union (CGL2016-76431-P and CGL2017-82654-P, AEI/FEDER-UE) and the Generalitat de  
1341 Catalunya (CERCA Programme). LS is supported by the FI AGAUR fellowship (ref. 2020 FI\_B1 00131)

1342 funded by the Secretaria d'Universitats i Recerca de la Generalitat de Catalunya and the European Social  
1343 Fund. The present manuscript has been prepared during a Visiting Professor fellowship at the Earth Sciences  
1344 Department of the University of Florence awarded to JM-M in May–June 2021. The present manuscript is  
1345 part of the PhD thesis of LS within the Geology PhD program of the Universitat Autònoma de Barcelona.  
1346 We would like to thank Paola Romi (Soprintendenza Archeologia Belle Arti e Paesaggio dell'Umbria) for  
1347 the possibility to study the Pietrafitta bovid collection at the MPLB and for her kindness and availability;  
1348 Elisabetta Cioppi and Luca Bellucci for the possibility to study the Valdarno bovid collections at the IGF;  
1349 Lorenzo Rook for the possibility to study the Pirro Nord bovid collection at the DST; Maia Bukshianidze for  
1350 having sent to us the 3D model and pictures of *B. (E.) georgicus* holotype and for precious information about  
1351 the Dmanisi bovid; Marco Sami for having sent to us measurements and pictures of the Salita di Oriolo  
1352 bovid; Michela Contessi for having sent to us pictures of the Mugello bovid; Roman Croitor for having sent  
1353 to us measurements of some East European *B. (B.) priscus*; Omar Cirilli for the help in the bibliographic and  
1354 data collection at the DST; Dawid Adam Iurino for the useful discussions about the methodological  
1355 approach; Grant Zazula and Elizabeth Hall for having sent to us some comparative *B. (B.) bison* pictures; and  
1356 Elpiniki Maria Parparousi, for the digital data elaboration.

1357

## 1358 **References**

- 1359 Agadzhanian, A.K., Vislobokova, I.A., Shunkov, M.V. and Ulyanov, V.A., 2017. Pleistocene mammal  
1360 fauna of the Trlica locality, Montenegro. *Fossil Imprint*, 73(1–2), 93–114.
- 1361 Alekseeva, L.I., 1967. To the history of subfamily Bovinae during the Pleistocene in the European USSR. In:  
1362 *Paleontologiya, geologiya i poleznye iskopaemye Moldavii Vyp. 2. Shtiintsa, Kishinev*, pp. 123–127  
1363 (in Russian).
- 1364 Ambrosetti, P., Cattuto, C. and Gregori, L., 1989. Geomorfologia e neotettonica nel bacino di  
1365 Tavernelle/Pietrafitta (Umbria). *Il Quaternario*, 2, 57–64.
- 1366 Ambrosetti, P., Argenti, P., Basilici, G., Gentili, S. and Ikome, F.E., 1992. The pleistocenic fossil vertebrata  
1367 of the Pietrafitta basin (Umbria, Italy): preliminary taphonomic analyses. In *Taphonomy: processes  
1368 and products. European Paleontological Association Workshop, Strasbourg* (pp. 20-21).

1369 Ambrosetti, P., Basilici, G., Capasso Barbato, L., Carboni, M.G., Di Stefano, G., Esu, D., Gliozzi, E.,  
1370 Petronio, C., Sardella, R. and Squazzini, E., 1995. Il Pleistocene inferiore nel ramo Sud-Occidentale  
1371 del Bacino Tiberino (Umbria): aspetti litostratigrafici e biostratigrafici. *Il Quaternario*, 8(1), 19–36.

1372 Arribas, A. and Palmqvist, P., 1998. Taphonomy and palaeoecology of an assemblage of large mammals:  
1373 hyaenid activity in the lower Pleistocene site at Venta Micena (Orce, Guadix-Baza Basin, Granada,  
1374 Spain). *Geobios*, 31, 3–47.

1375 Arzarello, M., Peretto, C. and Moncel, M.H., 2015. The Pirro Nord site (Apricena, Fg, Southern Italy) in the  
1376 context of the first European peopling: Convergences and divergences. *Quat. Int.*, 389, 255–263.

1377 Asperen E.N. van, 2010. Ecomorphological adaptations to climate and substrate in late Middle Pleistocene  
1378 caballoid horses. *Palaeogeogr. Palaeoclimatol. Palaeoecol.* 297: 584–596.

1379 Asperen van, E.N., Kahlke, R.D., 2017. Dietary traits of the late early Pleistocene *Bison menneri* (Bovidae,  
1380 Mammalia) from its type site Untermassfeld (Central Germany) and the problem of Pleistocene  
1381 ‘wood bison’. *Quat. Sci. Rev.* 177, 299–313.

1382 Baigusheva, V.S., 2000. New data on Tasmanian faunal assemblage from the excavation site near Semibalki  
1383 village (Azov Area). *Historical-Archaeological in Azov and Lower Don in 1998*, 16, 27–57.

1384 Biewener, A.A., 2005. Biomechanical consequences of scaling. *J. Exp. Biol.*, 208(9), 1665–1676.

1385 Blain, H.A., Fagoaga, A., Ruiz-Sánchez, F.J., Bisbal-Chinesta, J.F. and Delfino, M., 2019. Latest  
1386 Villafranchian climate and landscape reconstructions at Pirro Nord (southern Italy). *Geology*, 47(9),  
1387 829–832.

1388 Breda, M., Collinge, S.E., Parfitt, S.A., Lister, A.M., 2010. Metric analysis of ungulate mammals in the early  
1389 Middle Pleistocene of Britain, in relation to taxonomy and biostratigraphy: I: Rhinocerotidae and  
1390 Bovidae. *Quat. Int.* 228, 136–156.

1391 Brugal, J.-P., 1995. Le bison (Bovidae, Artiodactyla) du Pléistocène moyen ancien de Durfort (Gard,  
1392 France). *Bull. Mus. Natl. Hist. Nat.* 16C 2–4, 349–381.

1393 Bukhsianidze, M., 2005. The fossil Bovidae of Dmanisi. Ph.D. Dissertation, Università degli Studi di  
1394 Ferrara.

- 1395 Bukhsianidze, M. and Koiava, K., 2018. Synopsis of the terrestrial vertebrate faunas from the Middle Kura  
1396 Basin (Eastern Georgia and Western Azerbaijan, South Caucasus). *Acta Palaeontol. Pol.* 63(3), 441–  
1397 461.
- 1398 Bukhsianidze, M., 2020. New Results on Bovids from the Early Pleistocene site of Untermassfeld. In:  
1399 Kahlke, R.D. (Ed.), *Das Pleistozän von Untermaßfeld bei Meiningen (Thüringen). Teil 4.* Habelt-  
1400 Verlag, Bonn, pp. 1169–1195.
- 1401 Burchak-Abramovich, N.I., Gadzhiev, D.D., Vekua, A.K., 1980. Concerning ancestral form of bisons from  
1402 the Akchagyl of Southern Caucasus. *Bull. Georgian Natl. Acad. Sci.* 97(2), 48–488 (in Russian).
- 1403 Burchak-Abramovich, N.I., Gadzhiev, D.V., Vekua, A.K., 1994. On a new Pleistocene Bovine from Eastern  
1404 Georgia. *Teriologii Paleoteriologija*, 253–261 (in Russian).
- 1405 Caloi, L. and Palombo, M.R., 1995. The early Galerian fauna of Redicicoli. XIV INQUA Congress, Berlin.
- 1406 Cameron, C.C., Esterle, J.S. and Palmer, C.A., 1989. The geology, botany and chemistry of selected peat-  
1407 forming environments from temperate and tropical latitudes. *Int. J. Coal. Geol.* 12(1–4), 105–156.
- 1408 Castaños, J., Castaños, P., Murelaga, X. and Alonso-Olazabal, A., 2012. Kiputz IX: un conjunto singular de  
1409 bisonte estepario (*Bison priscus* Bojanus, 1827) del Pleistoceno Superior de la Península Ibérica.  
1410 *Ameghiniana*, 49(2), 247–261.
- 1411 Castelló, J. 2016. *Bovids of the World: Antelopes, Gazelles, Cattle, Goats, Sheep, and Relatives.* Princeton:  
1412 Princeton University Press, 664 p.
- 1413 Cherin, M., Bizzarri, R., Buratti, N., Caponi, T., Grossi, F., Kotsakis, T., Pandolfi, L., Pazzaglia, F., Barchi,  
1414 M.R., 2012. Multidisciplinary study of a new Quaternary mammal-bearing site from Ellera di  
1415 Corciano (central Umbria, Italy): preliminary data. *Rend. Onl. Soc. Geol. Ital.*, 21, 1075–1077.
- 1416 Cherin, M., D’Allestro, V., Masini, F., 2019. New bovid remains from the Early Pleistocene of Umbria  
1417 (Italy) and a reappraisal of *Leptobos merlai*. *J. Mamm. Evol.* 26, 201–224.
- 1418 Cignoni, P., Callieri, M., Corsini, M., Dellepiane, M., Ganovelli, F. and Ranzuglia, G., 2008. Meshlab: an  
1419 open-source mesh processing tool. In *Eurographics Italian chapter conference, 2008*, pp. 129–136.
- 1420 Clark, P.U., Archer, D., Pollard, D., Blum, J.D., Rial, J.A., Brovkin, V., Mix, A.C., Pisias, N.G., Roy, M.,  
1421 2006. The middle Pleistocene transition: characteristics, mechanisms, and implications for long-term  
1422 changes in atmospheric pCO<sub>2</sub>. *Quat. Sci. Rev.* 25, 3150–3184.

- 1423 Coltorti, M., Feraud, G., Marzoli, A., Peretto, C., Ton-That, T., Voinchet, P., Bahain, J.J., Minelli, A.,  
1424 Hohenstein, U.T., 2005. New  $^{40}\text{Ar}/^{39}\text{Ar}$ , stratigraphic and palaeoclimatic data on the Isernia la  
1425 Pineta Lower Palaeolithic site, Molise, Italy. *Quat. Int.* 131, 11–22.
- 1426 Conti, M.A. and Esu, D., 1981. Considerazioni sul significato paleoclimatico e geodinamico di una serie  
1427 lacustre pleistocenica inferiore presso Tavernelle (Perugia, Umbria). *Geogr. Fis. Din. Quat.*, 4, 3–10.
- 1428 Conti, M.A. and Girotti, O., 1977. Il Villafranchiano nel “Lago Tiberino”, ramo sud-occidentale: schema  
1429 stratigrafico e tettonico. *Geol. Romana*, 16, 67–80.
- 1430 Croitor, R., 2010. Critical remarks on genus *Bison* (Bovidae, Mammalia) from Pleistocene of Moldova.  
1431 *Revista Archeologica* 5, 172–188. (in Russian)
- 1432 Croitor R., 2016. Genus *Bison* (Bovidae, Mammalia) in Early Pleistocene of Moldova. In: E. Coropceanu  
1433 (Ed.), *Materialele Conferinței științifice naționale cu participare internațională “Mediul și dezvoltare  
1434 durabilă”*, Ediția a III-a, Chișinău, pp. 14–20.
- 1435 Croitor, R. and Popescu, A., 2011. Large-sized ruminants from the Early Pleistocene of Leu (Oltenia,  
1436 Romania) with remarks on biogeographical aspects of the " *Pachycrocuta* event". *Neues Jahrb. fur  
1437 Geol. Palaontol. - Abh.*, 261(3), 353–371.
- 1438 David, A.I. and Svistun, V.I., 1981. Bison remains from Upper Pliocene and Lower Pleistocene deposits of  
1439 Moldova and South Ukraine. In: *Biostratigraphy of Anthropogene and Neogene of South-West of  
1440 USSR*. Shtiintsa; Kishinev pp. 3–15 (in Russian).
- 1441 De Giuli, C., 1987. Late Villafranchian faunas in Italy: The Selvella Local Fauna in southern Chiana Valley -  
1442 Umbria. *Palaeontogr. Ital.* 74, 11–50.
- 1443 De Giuli, C., Masini, F. and Torre, D., 1987. The latest villafranchian faunas in Italy: the Pirro Nord fauna  
1444 (Apricena, Gargano). *Palaeontogr. Ital.* 74, 51–62.
- 1445 Depéret, C., 1884. Nouvelles études sur les ruminants pliocenes et quaternaires d’Auvergne. *Bull. Soc. Geol.*  
1446 *Fr.* 22, 247–284
- 1447 Dong, W., 2008. Nouveau matériel de *Leptobos* (*Smertiobos*) *crassus* (Artiodactyla, Mammalia) du  
1448 Pléistocène inférieur à Renzidong (Chine de l’Est). *Geobios* 41: 355–364.
- 1449 Dubrovo, I.A., and Burchak-Abramovich, N.I., 1983: A New Pliocene Bovid Genus, *Adjiderebos* gen. nov.,  
1450 from Transcaucasia. *Dokl. Akad. Nauk SSSR*, 276(3), 717–720.

- 1451 Dubrovo, I.A. and Burchak-Abramovich, N.I., 1986. New data on the evolution of Bovine of the tribe  
1452 Bovini. *Quartärpaläontologie*, 6,13–21.
- 1453 Duvernois, M.P., 1990. Les *Leptobos* (Mammalia, Artiodactyla) du Villafranchien d'Europe Occidentale.  
1454 *Doc. lab. géol. Fac. sci. Lyon* 113, 1–213.
- 1455 Duvernois, M.P., 1992. Mise au point sur le genre *Leptobos* (Mammalia, Artiodactyla, Bovidae);  
1456 implications biostratigraphiques et phylogénétiques. *Geobios* 25, 155–166.
- 1457 Duvernois, M.P., Guérin, C., 1989. Les Bovidae (Mammalia, Artiodactyla) du Villafranchien Supérieur  
1458 d'Europe Occidentale. *Geobios* 22: 339–379.
- 1459 Etienne, C., Filippo, A., Cornette, R. and Houssaye, A., 2021. Effect of mass and habitat on the shape of  
1460 limb long bones: A morpho- functional investigation on Bovidae (Mammalia: Cetartiodactyla). *J.*  
1461 *Anat.* 238(4), 886–904.
- 1462 Fabbi S., Romano M., Strani F., Sardella R., Bellucci L. (2021) The Pleistocene vertebrate fauna of the  
1463 Oricola-Carsoli intermontane Basin (Latium-Abruzzi, Italy): state of the art and historical review.  
1464 *Boll. Soc. Pal. Ital.* 60: 255–268.
- 1465 Falconer, H., 1868. *Paleontological memoirs and notes*, vol. II, l. R. J. Murchison (Ed.), London, 590 p.
- 1466 Flerov, C.C., 1972. The Most Ancient Bisons and the History of Genus *Bison*. In: *Teriologiya*, vol 1. Nauka,  
1467 Novosibirsk, pp. 81–86 (in Russian).
- 1468 Flerow, K.K., 1975. Die Bison-reste aus den travertinen von Weimar-ehringdorf. *Abh. Zentr. Geol. Inst.* 23,  
1469 171–199.
- 1470 Flerov, K.K., 1977a. Gigantic Bisons of Asia. *J. Palaeontol. Soc. India.* 20, 77–80.
- 1471 Flerov, K.K., 1977b. Die fossilen Bisonreste von Taubach und ihre Stellung in der Entwicklungsgeschichte  
1472 der Gattung *Bison* in Europa, In: Kahlke, R.-D. (Ed.), *Das Pleistozan von Untermaßfeld bei*  
1473 *Meiningen (Thuringen)*, Teil 1. *Quartarpalaontologie* 2, pp. 179–208.
- 1474 Flerov, K.K., 1979. Morphology, systematic, evolution, ecology. In: Sokolov, E.V. (Ed.), *European Bison*.  
1475 Nauka, Moscow, pp. 9–127 (in Russian).
- 1476 Foronova, I.V., 2001. Quaternary Mammals of the South-East of Western Siberia (Kuznetsk Basin).  
1477 *Phylogeny, Biostratigraphy, and Paleoecology*. A.V. Kanygin (Ed.), Publishing house of SB RAS.  
1478 Branch “GEO”, 234 p. (in Russian).

- 1479 Freudenberg, W., 1914. Die saugtiere des alteren Quartas von Mitteleuropa. Geol. Paleont, Abh, 12,  
1480 pp.455–670.
- 1481 Garrido, G., 2008. Primera cita de *Leptobos etruscus* (Falconer, 1868) (Bovidae, Artiodactyla, Mammalia)  
1482 en la Península Ibérica (Fonelas P-1, Cuenca de Guadix, Granada). Cuadernos del Museo Geominero  
1483 10: 489–516.
- 1484 Gentili, S., Abbazzi, L., Masini, F., Ambrosetti, P., Argenti, P. and Torre, D. 1996. Voles from the Early  
1485 Pleistocene of Pietrafitta (central Italy, Perugia). Acta Zool. Cracov., 3: 185–199.
- 1486 Gentili, S., Barili, A., Ambrosetti, P., 2000. Un museo per i fossili di Pietrafitta. Nuova Museologia, 2, 16–  
1487 17.
- 1488 Gentili, S., Masini F., 2005. An outline of Italian *Leptobos* and a first sight on *Leptobos* aff. *vallisarni* from  
1489 Pietrafitta (early pleistocene, Perugia) Quaternarie Hors Series, 2, 81–89.
- 1490 Gliozzi, E., Abbazzi, L., Argenti, P., Azzaroli, A., Caloi, L., Barbato, L.C., Di Stefano, G., Esu, D.,  
1491 Ficarelli, G., Girotti, O., Kotsakis, T., 1997. Biochronology of selected mammals, molluscs and  
1492 ostracods from the Middle Pliocene to the Late Pleistocene in Italy. The state of the art. Riv. Ital.  
1493 Paleontol. Stratigr, 103, 369–388.
- 1494 Geraads, D., 1992. Phylogenic analysis of the tribe Bovini (Mammalia: Artiodactyla). Zoological Journal of  
1495 the Linnean Society 104, 193–207.
- 1496 Gromov, I.M., Baranova, G.I. (Eds.), 1981. Catalogue of Mammals of the USSR (Pliocene-Recent). Nauka,  
1497 Leningrad, 456 p. (in Russian).
- 1498 Groves, C. and Grubb, P., 2011. Ungulate Taxonomy. The John Hopkins University Press, Baltimore, 309 p.
- 1499 Guérin, C., 1982. Première biozonation du Pléistocène européen, principal résultat biostratigraphique de  
1500 l'étude des Rhinocerotidae (Mammalia, Perissodactyla) du Miocène terminal au Pléistocène  
1501 supérieur d'Europe occidentale. Geobios, 15(4), pp.593–598.
- 1502 Guérin, C. and Valli, A.M., 2000. Le gisement pléistocène supérieur de la grotte de Jaurens à Nespouls,  
1503 Corrèze, France: les Bovidae (Mammalia, Artiodactyla). Publications du musée des Confluences,  
1504 1(1), 7–39.
- 1505 Hammer, Ø., Harper, D.A., Ryan, P.D., 2001. PAST: Paleontological statistics software package for  
1506 education and data analysis. Paleontol. Electron. 4, 1–9.

- 1507 Hassanin, A., 2014. Systematic and Evolution of Bovini, in Melletti, M. and Burton, J. (eds), Ecology,  
1508 Evolution and Behavior of Wild Cattle: Implications for Conservation. Cambridge University Press,  
1509 Cambridge: 7-20.
- 1510 Head, M.J., Gibbard, P.L., 2005. Early-Middle Pleistocene transitions: an overview and recommendation for  
1511 the defining boundary. Geol. Soc. London Special Publications 247, 1–18.
- 1512 Head, M.J. and Gibbard, P.L., 2015. Early–Middle Pleistocene transitions: linking terrestrial and marine  
1513 realms. Quat. Int. 389, 7–46.
- 1514 Jin, C., Wang, Y., Liu, J., Ge, J., Zhao, B., Liu, J., Zhang, H., Shao, Q., Gao, C., Zhao, K. and Sun, B., 2021.  
1515 Late Cenozoic mammalian faunal evolution at the Jinyuan Cave site of Luotuo Hill, Dalian,  
1516 Northeast China. Quat. Int. 577, 15–28.
- 1517 Jungers, W.L., Falsetti, A.B., Wall, C.E., 1995. Shape, relative size, and size-adjustments in morphometrics.  
1518 Am. J. Phys. Anthropol. 38, 137–161.
- 1519 Kahlke, H.-D. (Ed.), 1977. Das Pleistozän von Taubach bei Weimar. Quartärpaläontologie, vol. 2, 1–509 p.
- 1520 Kahlke, R.D., 1999. The History of the Origin, Evolution and Dispersal of the Late Pleistocene *Mammuthus-*  
1521 *Coelodonta* Faunal Complex in Eurasia (large mammals). Fenske Companies, Rapid City (1999),  
1522 218 p.
- 1523 Kelly, A., K., Miller, Joshua, H., DeSantis, L., Wooller, M., Zazula, G., 2021. Niche Stability Through  
1524 Environmental Change: A Late Pleistocene Chronology of Bison *Priscus* Paleoecology from Yukon  
1525 Territory, Canada. SVP 2021 Annual Meeting. 2021 Program and Abstract Book, pp. 157.
- 1526 Khan, M.A., Kostopoulos, D.S., Akhtar, D.S., Nazir, M., 2010. Bison remains from the upper Siwaliks of  
1527 Pakistan. N. Jahrb. Geol. Palaontol. Abh. 258, 121–128.
- 1528 Khan, M.A., Nasim, S., Ikram, T., Ghafoor, A. and Akhtar, M., 2011. Dental remains of early bison from the  
1529 Tatrot formation of the upper Siwaliks, Pakistan. J Anim Plant Sci, 21(4), 862–867.
- 1530 Klein, R.G., Franciscus, R.G., Steele, T.E., 2010. Morphometric identification of bovid metapodials to genus  
1531 and implications for taxon-free habitat reconstruction. J. Archaeol. Sci. 37, 389–401.
- 1532 Konidaris, G.E., Kostopoulos, D.S., Maron, M., Schaller, M., Ehlers, T.A., Aidona, E., Marini, M.,  
1533 Tourloukis, V., Muttoni, G., Koufos, G.D. and Harvati, K., 2021. Dating of the Lower Pleistocene

1534 vertebrate site of Tsiotra Vryssi (Mygdonia Basin, Greece): Biochronology, magnetostratigraphy,  
1535 and cosmogenic radionuclides. *Quat.*, 4(1), 1.

1536 Kostopoulos, D.S., Maniakas, I., Tsoukala, E., 2018. Early bison remains from Mygdonia Basin (Northern  
1537 Greece). *Geodiversitas* 40, 283–319.

1538 Kottek, M., Grieser, J., Beck, C., Rudolf, B., Rubel, F., 2006. World map of the Köppen-Geiger climate  
1539 classification updated. *Meteorol. Z.*, 15, 259–263.

1540 Krijgsman, W., Tesakov, A., Yanina, T., Lazarev, S., Danukalova, G., Van Baak, C.G., Agustí, J., Alçiçek,  
1541 M.C., Aliyeva, E., Bista, D. and Bruch, A., 2019. Quaternary time scales for the Pontocaspian  
1542 domain: Interbasinal connectivity and faunal evolution. *Earth-Sci. Rev.* 188, 1–40.

1543 Lydekker, R., 1878. Crania of ruminants from the Indian Tertiaires. *Pal. Indica* 10 (1), 88–171.

1544 Lopatin, A.V., Vislobokova, I.A., Lavrov, A.V., Startsev, D.B., Gimranov, D.O., Zelenkov, N.V.,  
1545 Maschenko, E.N., Sotnikova, M.V., Tarasenko, K.K. and Titov, V.V., 2019. The Taurida Cave, a  
1546 new locality of Early Pleistocene vertebrates in Crimea. *Dokl. Biol. Sci.* 485(1), 40–43.

1547 Lona, F. and Bertoldi, R. 1972. La storia del Plio-Pleistocene Italiano in alcune sequenze vegetation-ali  
1548 lacustri e marine. *Atti Accad. Naz. Lincei Cl. Sci. Fis. Mat. Natur. Memorie Serie 8* (11), 1–47.

1549 Madurell-Malapeira, J., Ros-Montoya, S., Espigares, M.P., Alba, D.M., Aurell-Garrido, J., 2014.  
1550 Villafranchian large mammals from the Iberian Peninsula: paleobiogeography, paleoecology and  
1551 dispersal events. *J. Iber. Geol.* 40, 167–178.

1552 Maniakas, I., Kostopoulos, D.S., 2017. Morphometric-palaeoecological discrimination between *Bison*  
1553 populations of the western Palaearctic. *Geobios* 50, 155–171.

1554 Martinetto, E., Bertini, A., Basilici, G., Baldanza, A., Bizzarri, R., Cherin, M., Gentili, S., Pontini, M.R.,  
1555 2014. The plant record of the Dunarobba and Pietrafitta sites in the Plio-Pleistocene  
1556 palaeoenvironmental context of Central Italy. *Alp. Mediterr. Quat.* 27, 29–72.

1557 Martinetto, E., Momohara, A., Bizzarri, R., Baldanza, A., Delfino, M., Esu, D., Sardella, R., 2017. Late  
1558 persistence and deterministic extinction of “humid thermophilous plant taxa of East Asian affinity”  
1559 (HUTEA) in southern Europe. *Palaeogeogr. Palaeoclimatol. Palaeoecol.*, 467, 211–231.

1560 Martínez-Navarro, B., Pérez-Claros, J.A., Palombo, M.R., Rook, L., Palmqvist, P., 2007. The Olduvai  
1561 buffalo *Pelorovis* and the origin of *Bos*. *Quat. Res.* 68, 220–226.

- 1562 Martínez-Navarro, B., Ros-Montoya, S., Espigares, M.P., Palmqvist, P., 2011. Presence of the Asian origin  
1563 Bovini, *Hemibos* sp. aff. *Hemibos gracilis* and *Bison* sp., at the early Pleistocene site of Venta  
1564 Micena (Orce, Spain). *Quat. Int.* 243, 54–60.
- 1565 Masini, F., 1989. I bovini villafranchiani dell'Italia. Ph.D. Dissertation, Università di Modena-Bologna-  
1566 Firenze-Roma.
- 1567 Masini, F., Palombo, M.R., Rozzi, R., 2013. A reappraisal of the Early to Middle Pleistocene Italian  
1568 Bovidae. *Quat. Int.* 288, 45–62.
- 1569 Maslin, M.A. and Brierley, C.M., 2015. The role of orbital forcing in the Early Middle Pleistocene  
1570 Transition. *Quat. Int.* 389, 47–55.
- 1571 McDonald, J. N., 1981. North American Bison: their Classification and Evolution. University of California  
1572 Press, Berkeley, 316 p.
- 1573 Mead, J.I., Jin, C., Wei, G., Sun, C., Wang, Y., Swift, S.L. and Zheng, L., 2014. New data on *Leptobos*  
1574 *crassus* (Artiodactyla, Bovidae) from Renzidong Cave, Early Pleistocene (Nihewanian) of Anhui,  
1575 China, and an overview of the genus. *Quat. Int.* 354, 139–146.
- 1576 Meiri, S., Yom- Tov, Y. and Geffen, E., 2007. What determines conformity to Bergmann's rule?. *Glob. Ecol.*  
1577 *Biogeogr.*, 16(6), 788–794.
- 1578 Merla, G., 1949. I *Leptobos* Rütim. italiani. *Palaeontogr. Ital.* 46, 41–155.
- 1579 Moullé, P. É., 1992. Les grands mammifères du Pléistocène inférieur de la grotte du Vallonet (Roquebrune-  
1580 Cap-Martin, Alpes-Maritimes). Étude paléontologique des Carnivores, Équidés, Suidés et Bovidés.  
1581 Ph.D. Dissertation, Muséum national d'Histoire naturelle, Paris.
- 1582 Mourer-Chauviré, C., 1972. Etude de nouveaux restes de vertébrés provenant de la carrière Fournier à  
1583 Châtillon-Saint-Jean. III. Artiodactyles, chevaux, oiseaux. *Quaternaire* 9, 271–305.
- 1584 Moyà-Solà, S., 1987. Los bóvidos (Artiodactyla, Mammalia) del yacimiento del Pleistoceno inferior de  
1585 Venta Micena (Orce, Granada, Espana). *Paleontol. Evol. Mem. Esp.* 1, 181–236.
- 1586 Olsen, S.J., 1990. Fossil ancestry of the yak, its cultural significance and domestication in Tibet. *Proc. Acad.*  
1587 *Nat. Sci. Philadelphia*, pp.73–100.
- 1588 Palmqvist, P., Espigares, M.P., Pérez-Claros, J.A., Figueirido, B., Guerra-Merchán, A., Ros-Montoya, S.,  
1589 Rodríguez-Gómez, G., García-Aguilar, J.M., Granados, A. and Martínez-Navarro, B., 2022. Déjà vu:

- 1590 a reappraisal of the taphonomy of quarry VM4 of the Early Pleistocene site of Venta Micena (Baza  
1591 Basin, SE Spain). *Sci. Rep.* 12(1), 1–16.
- 1592 Pazzaglia, F., Barchi, M.R., Buratti, N., Cherin, M., Pandolfi, L. and Ricci, M., 2013. Pleistocene calcareous  
1593 tufa from the Ellera basin (Umbria, central Italy) as a key for an integrated paleoenvironmental and  
1594 tectonic reconstruction. *Quatern. Int.*, 292, 59–70.
- 1595 Pilgrim, G.E., 1937. Siwalik antelopes and oxen in the American Museum of Natural History. *Bull. Am.*  
1596 *Mus. Nat. Hist.* 72, 7.
- 1597 Pilgrim, G.E., 1939. The fossil Bovidae of India. *Mem. Geol. Surv. India, Pal. Indica*, n.s. 26 (1), 356 p.
- 1598 Plummer, T.W., Bishop, L.C., 1994. Hominid paleoecology at Olduvai Gorge, Tanzania as indicated by  
1599 antelope remains. *J. Hum. Evol.* 27, 47–75.
- 1600 Rodrigo, M.A., 2011. Los Bóvidos Villafranchienses de La Puebla de Valverde y Villarroya: sistemática,  
1601 filogenia y paleobiología. Ph.D. Dissertation, Universidad de Zaragoza.
- 1602 Rook, L. and Martínez-Navarro, B., 2010. Villafranchian: the long story of a Plio-Pleistocene European large  
1603 mammal biochronologic unit. *Quat. Int.* 219(1–2), 134–144.
- 1604 Rüttimeyer, L., 1867. Versuch einer natürlichen Geschichte des Rindes. *Neuen Denkschr. Algem. Schweiz.*  
1605 *Ges. Naturwiss.* II (Vol. 22).
- 1606 Saarinen, J., Eronen, J., Fortelius, M., Seppä, H. and Lister, A.M., 2016. Patterns of diet and body mass of  
1607 large ungulates from the Pleistocene of Western Europe, and their relation to vegetation. *Palaeontol.*  
1608 *Electron.* 19(3), 1–58.
- 1609 Saarinen, J., Cirilli, O., Strani, F., Meshida, K. and Bernor, R.L., 2021. Testing Equid Body Mass Estimate  
1610 Equations on Modern Zebras—With Implications to Understanding the Relationship of Body Size,  
1611 Diet, and Habitats of *Equus* in the Pleistocene of Europe. *Front. Ecol. Evol.* 9, 90.
- 1612 Sala, B., 1986. *Bison schoetensacki* Freud. from Isernia la Pineta (early Mid-Pleistocene – Italy) and revision  
1613 of the European species of bison. *Palaeontogr. ital.* 74, 113–170.
- 1614 Schertz, E., 1936. Zur Unterscheidung von *Bison priscus* Boj. und *Bos primigenius* Boj. an Metapodien und  
1615 Astragalus, nebst Bemerkungen über einige diluviale Fundstellen. *Seckenbergiana* 18, 37–71.
- 1616 Schreuder, A., 1946. The Tegelen fauna, with a description of new remains of its rare components (*Leptobos*,  
1617 *Archidiskodon meridionalis*, *Macaca*, *Sus strozzii*). *Arch. Neerl. Zool.*, 7(1), 153–204.

- 1618 Scott, K.M., 1979. Adaptation and allometry in bovid postcranial proportions. Ph.D. Dissertation. Yale  
1619 University, New Haven.
- 1620 Scott, K.M., 1983. Prediction of body weight of fossil Artiodactyla. Zool. J. Linn. Soc. 77(3), 199–215.
- 1621 Scott, K.M., 1985. Allometric trends and locomotor adaptations in the Bovidae. Bull. Am. Mus. Nat. Hist.  
1622 197, 197–288.
- 1623 Scott, R.S., Barr, W.A., 2014. Ecomorphology and phylogenetic risk: implications for habitat reconstruction  
1624 using fossil bovids. J. Hum. Evol. 73, 47–57.
- 1625 Shani, M. R. and Khan, E., 1968. *Probison dehmi* n. g. n. sp., a recent find of the Upper Siwalik bovid. Mitt.  
1626 Bayer. Staatssamml. Paläontol. Hist. Geol. 8, 247–251.
- 1627 Sher, A.V., 1997. An Early Quaternary *Bison* population from Untermaßfeld: *Bison menneri* sp. nov. In:  
1628 Kahlke, R.D. (Ed.), Das Pleistozän von Untermaßfeld bei Meiningen (Thüringen). Teil 2. Habelt-  
1629 Verlag, Bonn, pp. 101–180.
- 1630 Simpson, G.G., 1941. Large Pleistocene felines of North America. Am. Mus. Novit. 1136, 1–27.
- 1631 Skinner, M.R., Kaisen, O.C., 1946. The fossil *Bison* of Alaska and preliminary revision of the genus. Bull.  
1632 Am. Mus. Nat. Hist. 89, 123–256.
- 1633 Sorbelli, L., Alba, D.M., Cherin, M., Moullé, P.É., Brugal, J.P. and Madurell-Malapeira, J., 2021a. A review  
1634 on *Bison schoetensacki* and its closest relatives through the early-Middle Pleistocene transition:  
1635 Insights from the Vallparadís Section (NE Iberian Peninsula) and other European localities. Quat.  
1636 Sci. Rev. 261, 106933.
- 1637 Sorbelli, L., Villa A., Gentili S., Cherin M., Carnevale, G., Tschopp E., and Delfino, M., 2021b. The Early  
1638 Pleistocene ectothermic vertebrates of Pietrafitta (Italy) and the last Western European occurrence of  
1639 *Latonia* Meyer, 1843. C. R. Palevol. 20(26), 555–583.
- 1640 Soubrier, J., Gower, G., Chen, K., Richards, S.M., Llamas, B., Mitchell, K.J., Ho, S.Y., Kosintsev, P., Lee,  
1641 M.S., Baryshnikov, G., Bollongino, R., 2016. Early cave art and ancient DNA record the origin of  
1642 European bison. Nat. Commun. 7, 13158.
- 1643 Stephenson, R.O., 2001. Wood bison in late Holocene Alaska and adjacent Canada: paleontological,  
1644 archaeological and historical records. In: SC Gerlach and MS Murray (Eds), Wildlife and People in

1645 Northern North America. Essays in Honor of R. Dale Guthrie. British Archaeological Reports,  
1646 International Series, pp. 944.

1647 Teilhard de Chardin, P., Piveteau, J., 1930. Les mammifères fossiles de Nihowan (Chine). Ann. Paleontol.  
1648 19, 1–134.

1649 Teilhard de Chardin, P. and Trassaert, M., 1938. Caviconia of South-Eastern Shansi. Pal. Sin. New Ser. C, 6,  
1650 1–106.

1651 Tong, H.W., Chen, X., Zhang, B., 2016. New fossils of *Bison palaeosinensis* (Artiodactyla, Mammalia) from  
1652 the steppe mammoth site of early Pleistocene in Nihewan Basin, China. Quat. Int. 445, 450–468.

1653 Toniato, G., Marabini, S., Sala, B. and Vai, G.B., 2017. Revised bison skull from the Salita di Oriolo quarry  
1654 near Faenza, “sabbie gialle”, Pleistocene, Northern Apennines. Alp. Mediterr. Quat. 30(1), 11–23.

1655 Torre, D., Ficarelli, G., Masini, F., Rook, L. and Sala, B., 1992. Mammal dispersal events in the early  
1656 Pleistocene of Western Europe. Cour. Forsch. Inst. Senckenberg, 153, 51–58.

1657 Vasiliev, S.K., 2008. Late Pleistocene bison (*Bison p. priscus* Bojanus, 1827) from the southeastern part of  
1658 western Siberia. Archaeol. Ethnol. Anthropol. Eurasia 34, 34–56.

1659 Vercoutère, C. and Guérin, C., 2010. les Bovidae (Mammalia, Artiodactyla) du Pléistocène moyen final de  
1660 l’Aven de Romain-La-Roche (Doubs, France). Rev. de Paleobiologie 9, 655–696.

1661 Verestchagin, N.K., 1959. The Mammals of the Caucasus. A History of the Fauna. Academia Nauk,  
1662 Leningrad (translated from Russian; Israel Program for Scientific Translations, Jerusalem 1967).

1663 Viret, J., 1949. La vie dans la moyenne vallée du Rhône au début des temps quaternaires (Essai d’écologie de  
1664 la faune des mammifères fossiles de Saint-Vallier). Bull. Mens. Soc. Linn. Lyon 18, 20–24

1665 Viret, J., 1954. Le loess à banc durcis de Saint-Vallier (Drôme) et sa faune de mammifères villafranchiens.  
1666 Avec une analyse granulométrique. Nouv. Arch. Mus. Hist. Nat. Lyon 4, 1–200.

1667 Zheng, S.H., Wu, W.Y., Li, Y., Wang, G.D., 1985. Late Cenozoic mammalian faunas of Guide and Gonghe  
1668 Basins, Qinghai Province. Vert. PalAs. 23, 89–134 (in Chinese).

1669 Zucchetta, G., Gentili, S. and Pavia, M., 2003.: A new early Pleistocene bird association from Pietrafitta  
1670 (Perugia, Central Italy). Riv. Ital. Paleontol. S., 109, 527–538.

1671

1672

1673 **Figures Captions**

1674 **Fig. 1** a, Major Eurasian sites with the occurrence of primitive *Bison*: 1, Venta Micena (Spain); 2, Sainzelles  
1675 (France); 3, Maar du Riege (France); 4, Salita di Oriolo (Italy); 5, Central Italy (Upper Valdarno, Pietrafitta,  
1676 Capena); 6, Pirro Nord (Italy); 7, Trlica (Montenegro); 8, Mygdonia Basin (Greece); 9, Taurida Cave  
1677 (Ukraine); 10, Dmanisi (Georgia); 11, Kushkuna (Azerbaijan); 12, Upper Siwaliks (Pakistan); 13, North-  
1678 West China (Bajjazui, Huiduipo, Zhangjiapo, Yangguo, Genjiagou, Gonghe); 14, South Shanxi area (Yushe,  
1679 Linyi, Pinglu, Mianchi; China); 15, Nihewan Basin area (Jiajiashan, Chifeng, Xiashagou, Shanshenmiaozui,  
1680 Cenjiawan; China). Shaded blue area: putative geographic distribution of *B. priscus* during the Pleistocene  
1681 (taken from Kahlke, 1999). b, Time scale showing the chronological/biochronological distribution of the  
1682 studied taxa in Europe and the most important Plio-Pleistocene sites of Europe and Western Asia mentioned  
1683 in the text. Abbreviations: Ma, mega annum; GC, geochronology; ELMA, European Land Mammal Age; C,  
1684 glacial/interglacial cyclicality; MS, magnetostratigraphy.

1685 **Fig. 2** Cranial remains of *Bison (Eobison) degiulii* from Pietrafitta. a, Cranium SABAP UMB 19. 2.1178 in  
1686 dorsal view; b Cranium SABAP UMB 19. 2.1179 in posterior (b1) and ventral (b2) views; c, Atlas SABAP  
1687 UMB 21. 4.1314 in anterior (c1), posterior (c2), dorsal (c3), and ventral (c4) views; d, Occipital and  
1688 basioccipital fragments SABAP UMB 21. 4.1177 in posterior (d1) and ventral (d2) views; e, Occipital and  
1689 basioccipital fragments SABAP UMB 19. 2.1179 in posterior (e1) and ventral (e2) views. Scale bar: 100  
1690 mm.

1691 **Fig. 3** Cranial remains of *Bison (Eobison) degiulii* from Pietrafitta. a, Basioccipital SABAP UMB 20.  
1692 1.6802/4 in ventral (a1) and lateral (a2) views; b, Basioccipital SABAP UMB 20. 1.6803/2 in ventral (b1)  
1693 and lateral (b2) views; c, Horn-core tip SABAP UMB 21. 4.1323; d, Left horn-core SABAP UMB 19.  
1694 2.1179/1 in posterior (d1), dorsal (d2), and ventral (d3) views; e, Left horn-core SABAP UMB 21. 4.1179 in  
1695 posteroventral view; f, Nasal fragments SABAP UMB 20. 1.6802/3 in dorsal view; g, Right horn-core  
1696 SABAP UMB 19. 2.1179/2 in anterior view; h, Right premaxilla SABAP\_UMB 20.1.6802/2 in lateral view;  
1697 i, Left horn-core SABAP UMB 20. 1.6803/1 in posterior view; j, Left horn-core SABAP UMB 20. 1.6802/5  
1698 in anterior view; k, Right (?) horn-core tip SABAP\_UMB 20.1.7345 in posterior (k1) and dorsal (k2) views.  
1699 Scale bar: 50 mm.

1700 **Fig. 4** Dentognathic remains of *Bison (Eobison) degiulii* from Pietrafitta. a, Right M3 SABAP UMB 20.  
1701 1.7346/1 in occlusal (a1), labial (a2), and lingual (a3) views; b, Right M3 SABAP UMB 20. 1.6804/2 in  
1702 occlusal view; c, Right M2 SABAP UMB 20. 1.7348 in occlusal (c1) and labial (c2) views; d, Right M2  
1703 SABAP UMB 21. 4.1325 in occlusal view; e, Left M1 SABAP UMB 20. 1.6804/5 in occlusal (e1) and  
1704 lingual (e2) views; f, Right M1 SABAP UMB 20. 1.7350 in occlusal view; g, Right P4 SABAP UMB 21  
1705 .4.1230 in occlusal (g1) and labial (g2) views; h, Right P4 SABAP UMB 20. 1.6804/7 in occlusal view; i,  
1706 Left P3 SABAP UMB 20. 1.6802/14 in occlusal (i1) and labial (i2) views; j, Left P2 SABAP UMB 20.  
1707 1.7361 in occlusal (j1) and labial (j2) views; k, Right m3 SABAP UMB 20. 1.1227 in occlusal (k1), labial  
1708 (k2), and lingual (k3) views; l, Left m2 SABAP UMB 20. 1.7365 in occlusal (l1) and labial (l2) views; m,  
1709 Left m1 SABAP UMB 20. 1.1228 in occlusal (m1) and labial (m2) views; n, Right hemimandible with p4-  
1710 m1 SABAP UMB 21. 4.1235 in occlusal (n1) and lingual (n2) views; o, Right hemimandible with p2-p4  
1711 SABAP UMB 21. 4.1236 in labial (o1) and occlusal (o2) views; p, Fragment of right mandibular ramus  
1712 SABAP UMB 21. 4.1298 in lateral view; q, Right hemimandible with p2-m3 SABAP UMB 130074 in  
1713 occlusal (q1) and lingual (q2) views; r, Right hemimandible with p2-m3 SABAP UMB 20. 1.7341 in  
1714 occlusal (r1) and labial (r2) views; s, Left hemimandible with p2-m3 SABAP UMB 20. 1.7342 in occlusal  
1715 (s1) and labial (s2) views. Scale bars: 30 mm (top) and 50 mm (bottom).

1716 **Fig. 5** Metapodials of *Bison (Eobison) degiulii* from Pietrafitta. a, Left metacarpal SABAP UMB 20. 1.7325  
1717 in anterior (a1) and proximal (a2) views; b, Right metacarpal SABAP UMB 130038 in anterior (b1) and  
1718 proximal (b2) views; c, Right metacarpal SABAP UMB 22. 4.247/8 in anterior (c1) and proximal (c2) views;  
1719 d, Left metacarpal SABAP UMB 20. 1.7324 in anterior (d1) and proximal (d2) views; e, Right metatarsal  
1720 SABAP UMB 20. 1.7330 in anterior (e1) and proximal (e2) views; f, Right metatarsal SABAP UMB 20.  
1721 1.7331 in anterior (f1) and proximal (f2) views; g, Left metatarsal in anterior (g1) and proximal (g2) views;  
1722 h, Right metatarsal SABAP\_UMB 20.1.7332 in anterior (h1) and proximal (h2) views. Scale bars: 50 mm.

1723 **Fig. 6** Horn-core morphology in several species of *Leptobos* and *Bison* s.l. a, Biplot of the horn-core section  
1724 at the base. Measurement abbreviations are explained in Table S2 and shown in Fig. S1. Measurements are  
1725 in mm. b, Diagram of the angle between the sagittal axis of the cranium and the horn-core midline.

1726 **Fig. 7** Biometric comparisons of cranial and dental variables in several species of *Leptobos* and *Bison* s.l. a,  
1727 Biplot of the intertemporal bridge shape. b, Biplot of the frontal shape. Measurement abbreviations are

1728 explained in Table S2 and shown in Fig. S1. Measurements are in mm. c, Log<sub>10</sub> ratio of upper molars and  
1729 third lower molar; d, Principal component analysis of six selected variables of the occipital and basioccipital  
1730 areas.

1731 **Fig. 8** Bivariate plots of postcranial stoutness in several species of *Leptobos* and *Bison* s.l. a, Metacarpals; b,  
1732 Metatarsals; c, Radii; d, Astragali. Measurement abbreviations are explained in Table S2 and shown in Fig.  
1733 S1. Measurements of Lmax are in mm.

1734 **Fig. 9** Comparisons of limb bone dimensions and proportions in several species of *Leptobos* and *Bison* s.l. a–  
1735 b, Bivariate plots of the first two principal component (PC) scores resulting from PCA of metacarpals (a) and  
1736 metatarsals (b), based on seven selected variables presented in Table 1; c, Histogram representing the ratio  
1737 between the lengths of metacarpal (MC), radius (RAD), and metatarsal (MT).

1738 **Fig. 10** Morphological comparison between *Bison* (*Eobison*) specimens from the Mediterranean area, studied  
1739 in this work. a, *B. (E.) degiulii* from Pietrafitta, SABAP UMB 19. 2.1178 in dorsal view; b, *B. (E.) degiulii*  
1740 from Pietrafitta, SABAP UMB 21. 4.1177 in posterior view; *B. (E.) degiulii* from Pietrafitta, SABAP UMB  
1741 19. 2.1179 in posterior (c1) and ventral (c2) views; d, *B. (E.) cf. degiulii* from Capena, MUST (ex MPUR)  
1742 s.n., in dorsal (d1), posterior (d2), and ventral (d3) views; e, *B. (E.) cf. degiulii* from Salita di Oriolo,  
1743 MCSNF s.n., in dorsal (e1), posterior (e2), right lateral (e3), and ventral (e4) views; f, *B. (E.) degiulii*  
1744 holotype from Pirro Nord, IGPB s.n., in dorsal (f1), posterior (f2), right lateral (f3), and ventral (f4) views;  
1745 g, *B. (E.) degiulii* from Kalamoto-2, NHCK KLT-638, in dorsal (g1), posterior (g2), left lateral (g3), and  
1746 ventral (g4) views; f, *B. (Eobison)* sp. from Mugello, MGCB s.n., in anterior (h1) and dorsal (h2) views.  
1747 Scale bar: 100 mm.

1748 **Fig. 11** Evolutionary trends in *Leptobos* and *Bison* over the last 3 Ma. a, Stoutness violinplot of metacarpals;  
1749 b, Body mass estimation violinplot. The numbers at the base of the plots indicate the number of specimens  
1750 for each sample. At the bottom, the age of the samples is shown on a chronological and paleoclimatological  
1751 diagram of LR04 Benthic Stack (from Lisiecki and Raymo, 2005); each paleontological locality is associated  
1752 with the predominant paleoenvironment in Western Palearctic in the considered time span. European Land  
1753 Mammal Ages are indicated at the bottom. Abbreviations: Ta, Taubach; CSJ, Chatillon St. Jean; RLR,  
1754 Romain-la-Roche; JMCMA, Joint Mitnor Cave Mammal Assemblage sites; Ha, Habarra; Gr, Gral; CF, Cava  
1755 Filo; Si, Sirejol; RB, Roter Berg.

1756 **Fig. 12** Comparisons of Bovinae considered in this review. a, Life reconstructions of *Leptobos* and *Bison*  
1757 species from the late Pliocene (left) to Late Pleistocene (right). From left to right: *Leptobos* gr. LSEM,  
1758 *Leptobos* gr. LEV, *Bison (Eobison) degiulii*., *Bison (Poephagus) menneri*, *Bison (Bison) schoetensacki*,  
1759 *Bison (Bison) priscus*. b–o, Morphological comparison between craniums of several species of Bovinae  
1760 mentioned in this work: b, *L. stenometopon* from Dusino in dorsal (b1) and posterior (b2) views; c, *L. merlai*  
1761 from St. Vallier in dorsal (c1) and posterior (c2) views; d, *A. cantabilis* from Kushkuna in dorsal (d1) and  
1762 posterior (d2) views; e, *L. etruscus* from Upper Valdarno in dorsal (e1) and posterior (e2) views; f, *L.*  
1763 *vallisarni* Le Fratte (Upper Valdarno) in dorsal (f1) and posterior (f2) views; g, *B. (E.) palaeosinensis* from  
1764 Nihowan in dorsal (g1) and posterior (g2) views; h, *B. (E.) georgicus* from Dmanisi in dorsal (h1) and  
1765 posterior (h2) views; i, *B. (E.) degiulii* from Pietrafitta in dorsal view; j, *B. (E.) degiulii* from Pirro Nord in  
1766 dorsal (j1) and posterior (j2) views; k, *B. (E.) degiulii* from Kalamoto-2 in dorsal (k1) and posterior (k2)  
1767 views; l, *B. (P.) menneri* from Untermassfeld in dorsal (l1) and posterior (l2) views; m, *B. (B.) schoetensacki*  
1768 from Mauer in dorsal (m1) and posterior (m2) views; n, *B. (B.) priscus* from Po River in dorsal (n1) and  
1769 posterior (n2) views; o, *B. (B.) bonasus* from Białowieża in dorsal (o1) and posterior (o2) views. Specimens  
1770 not in scale. Artworks by LS.

Figure 1

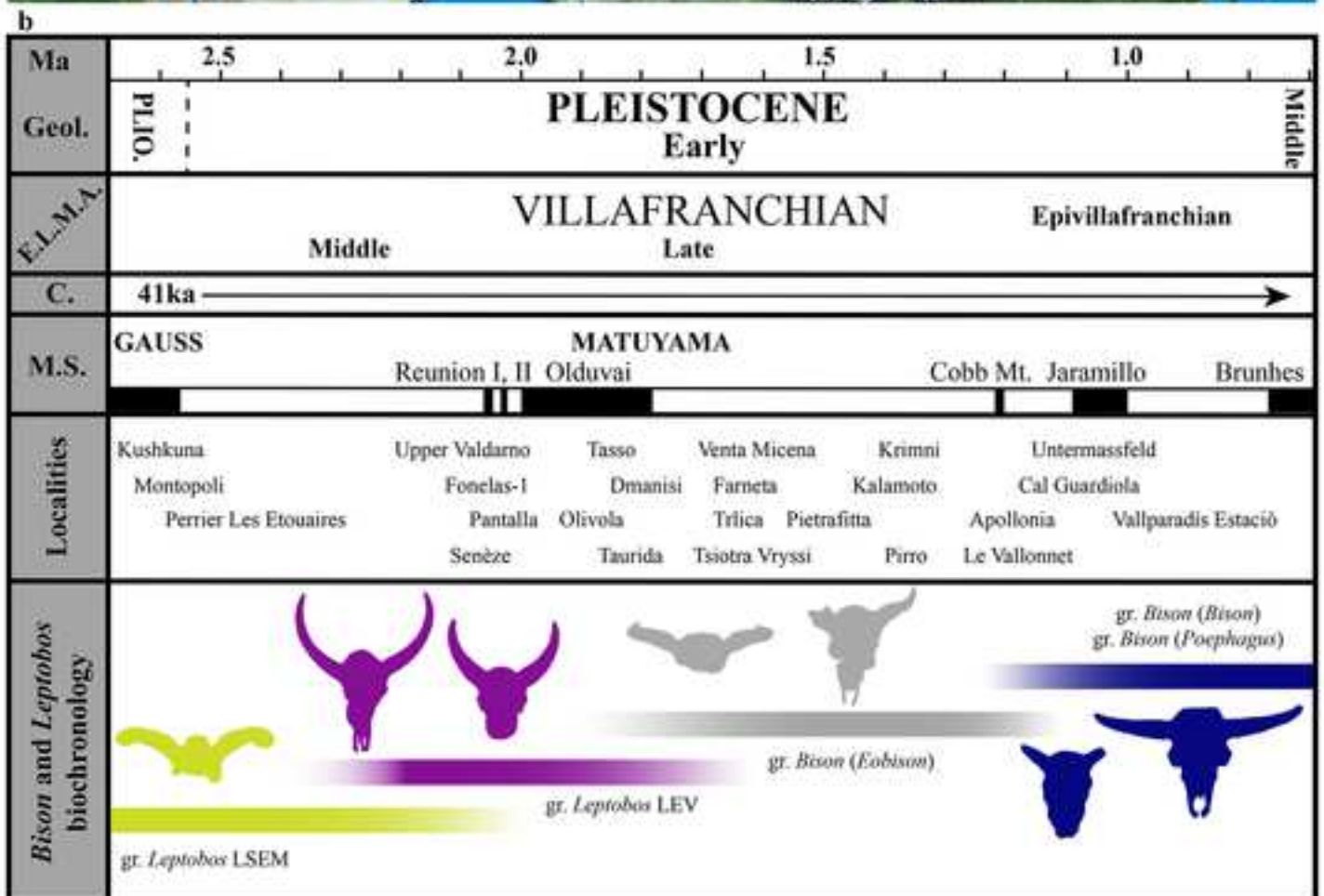


Figure 2



Figure 4



Figure 3

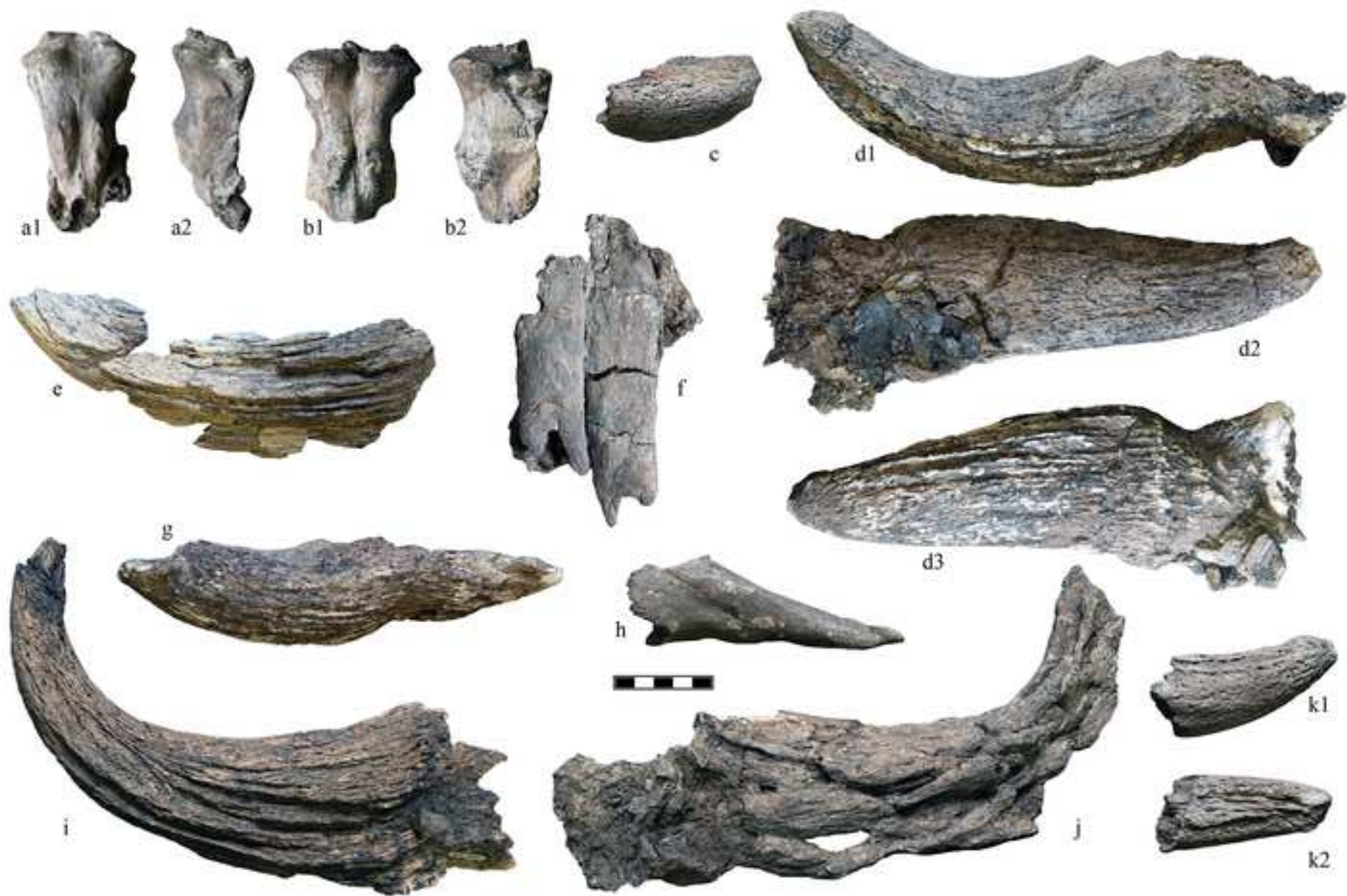


Figure 5



Figure 6

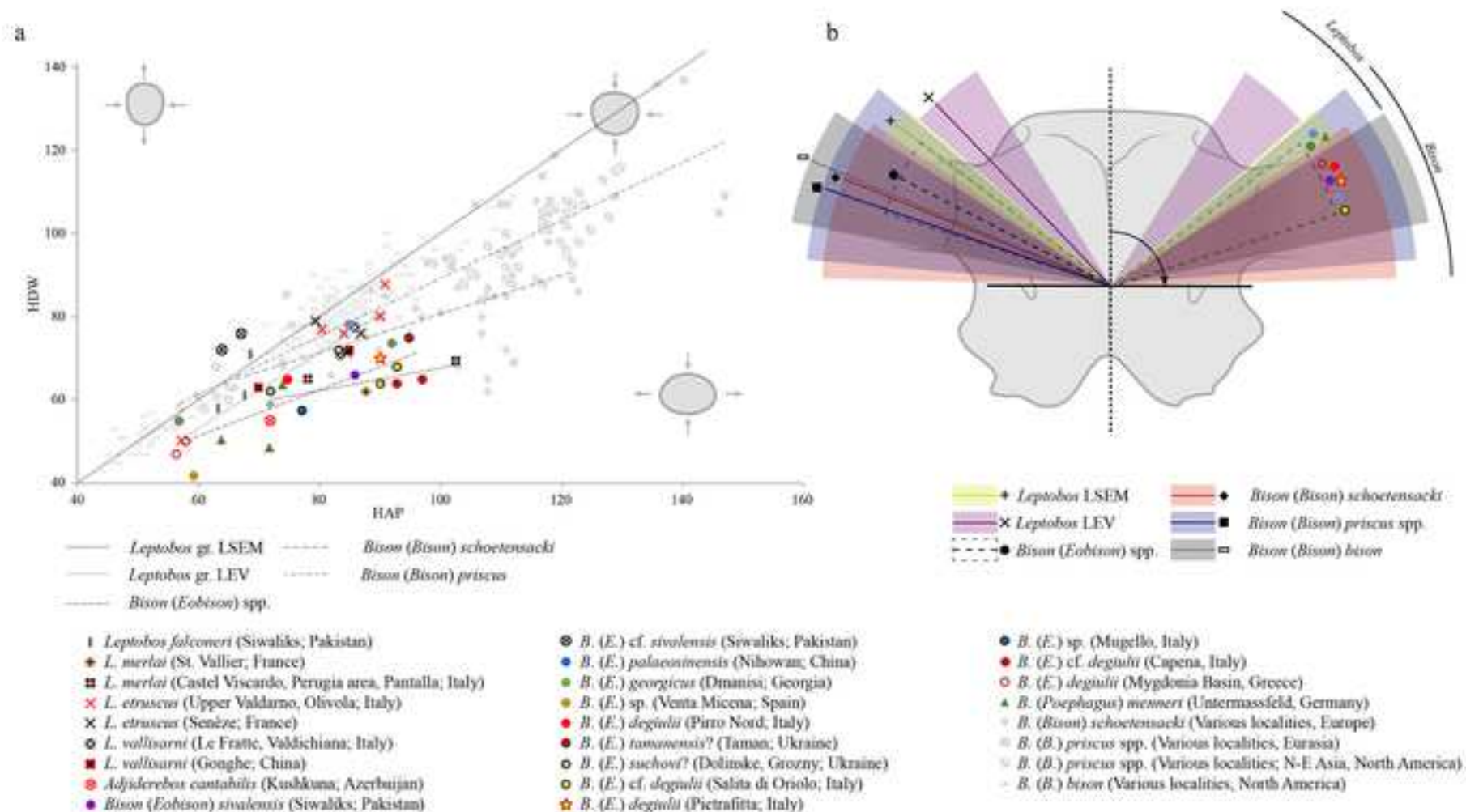


Figure 7

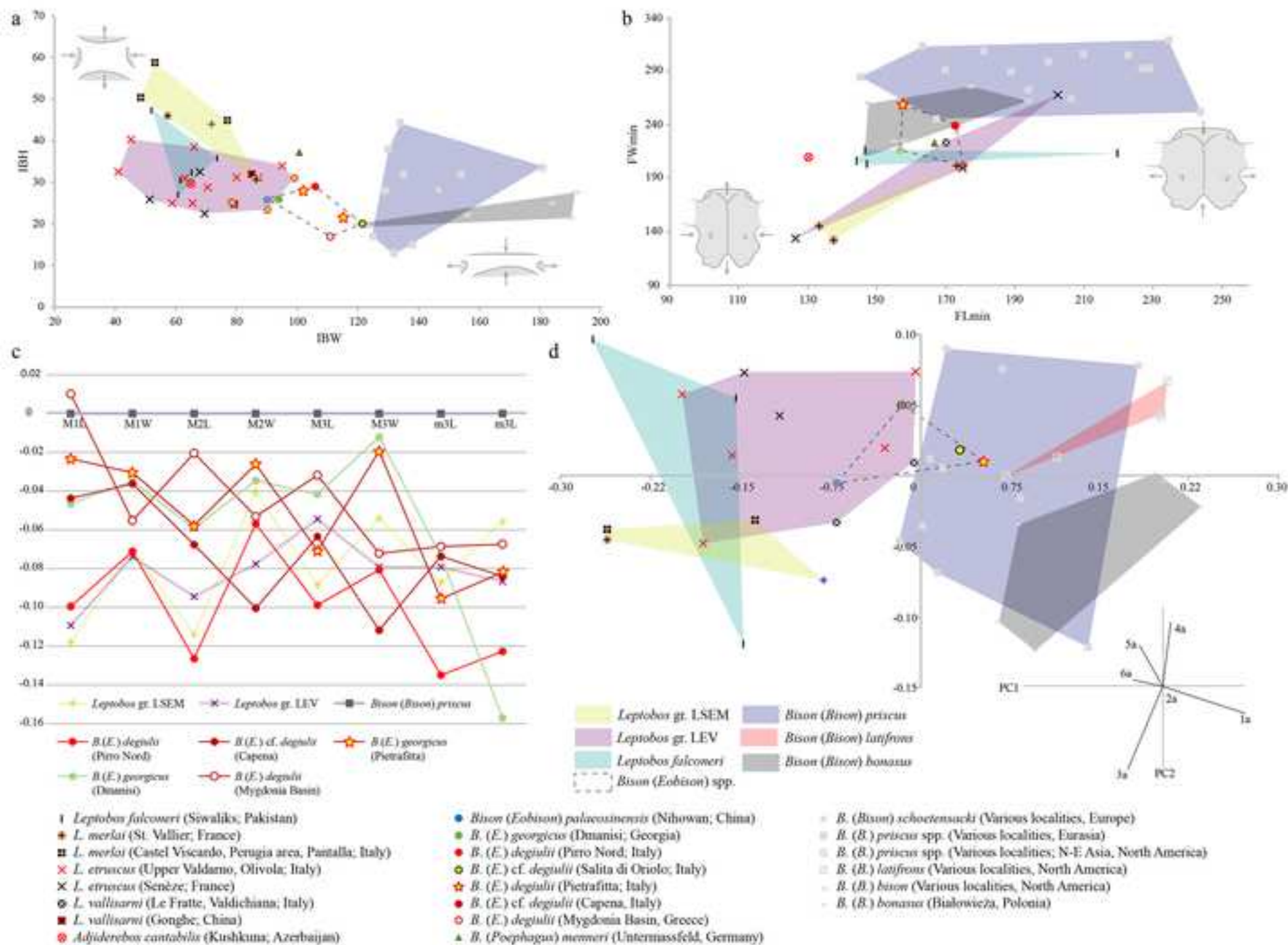


Figure 8

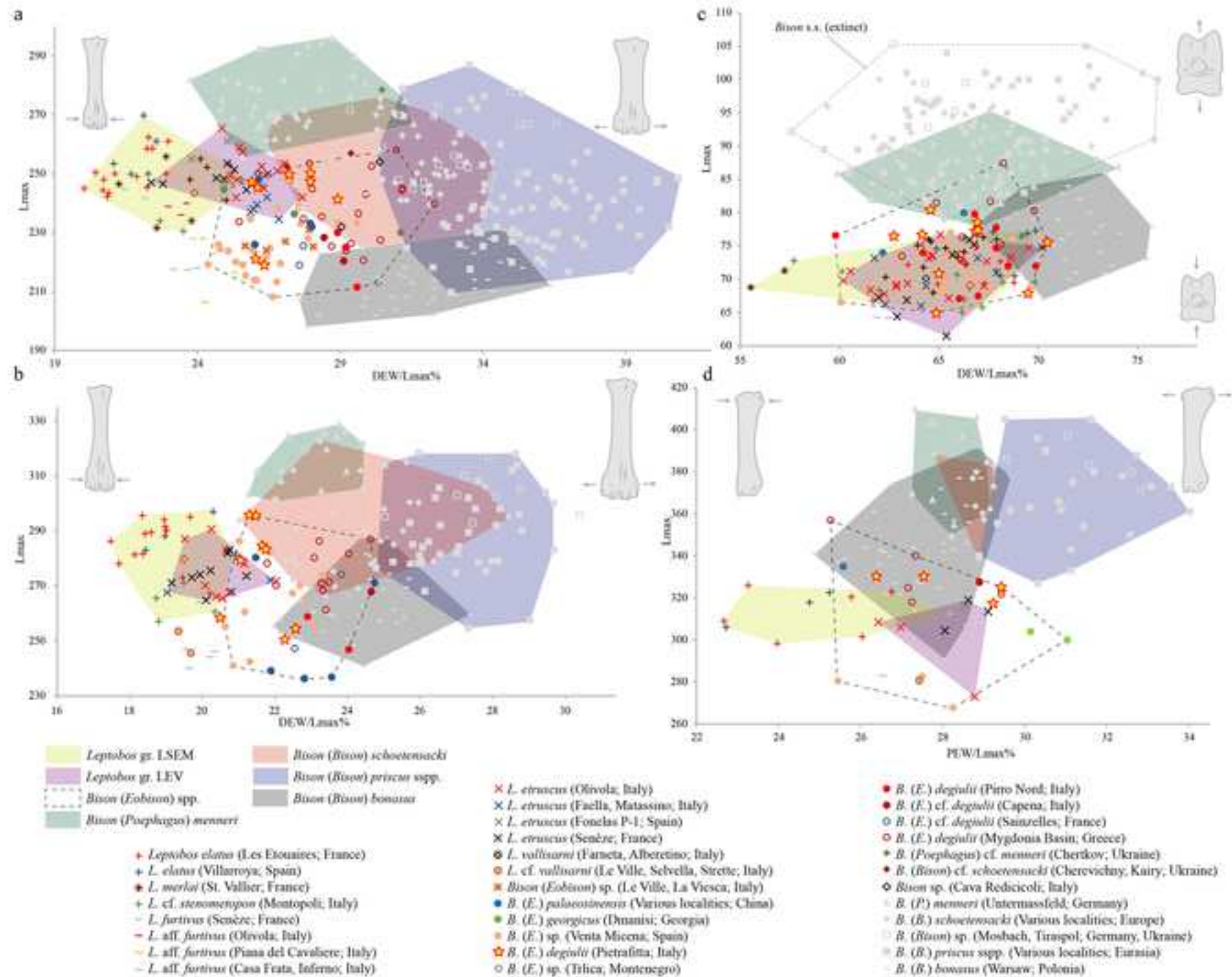


Figure 9

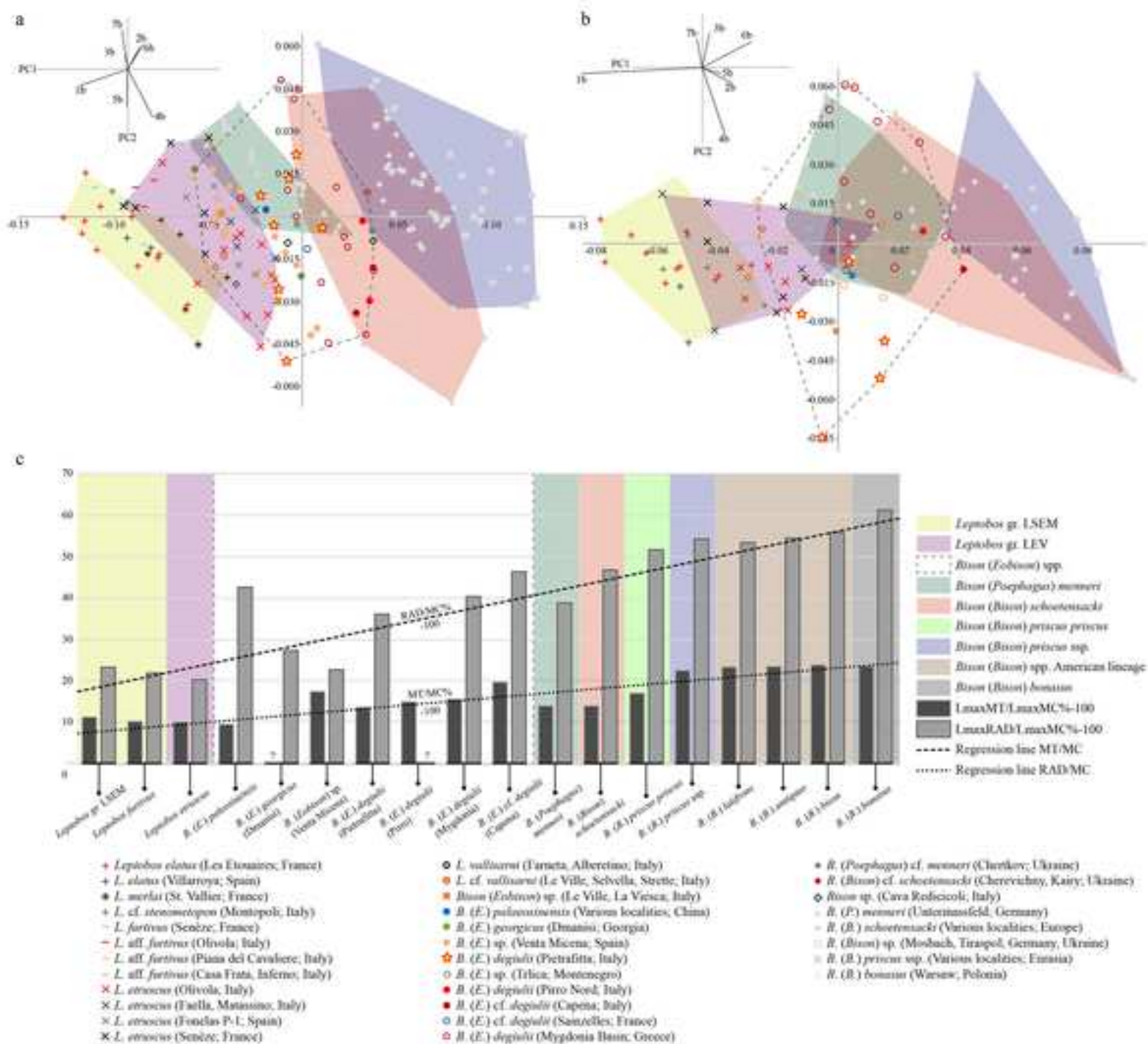
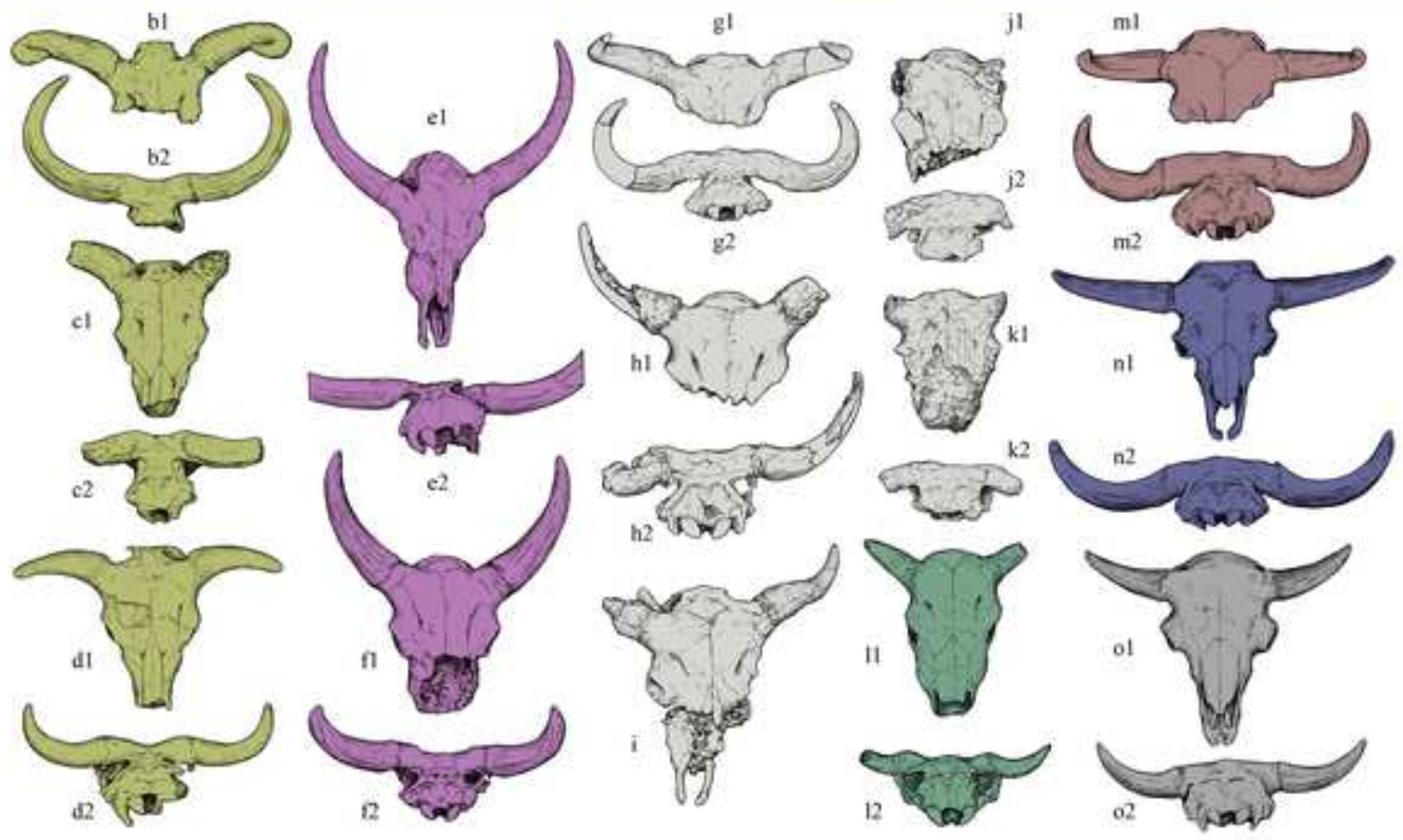


Figure 10





Figure 12



**Table 1.** Shape variables used for the principal component analysis. Abbreviations: ms, Mosimann shape variable; GM, Geometric Mean. Abbreviations of measurements as in Fig. S1 and Table S2.

<b>Selected Mosimann shape variables for cranium</b>	<b>Fig. 7</b>	<b>Selected Mosimann shape variables for metapodials</b>	<b>Fig. 9</b>
msIBW = Log (IBW/ GM)	1a	msLmax = Log (Lmax/GM)	1b
msCW = Log (CW/ GM)	2a	msPEW = Log (PEW/GM)	2b
msFMW = Log (FMW/ GM)	3a	msPET = Log (PET/GM)	3b
msMW = Log (MW/ GM)	4a	msDW = Log (DW/GM)	4b
msOSHmax = Log (OSHmax/ GM)	5a	msDT = Log (DT/GM)	5b
msPPW = Log (PPW/ GM)	6a	msDEW = Log (DEW/GM)	6b
		msDETm = Log (DETm/GM)	7b

**Table 2.** Measurements (mm) of the crania of *Bison (Eobison) degiulii* from Pietrafitta. Estimated measurements are in italics. Abbreviations as in Fig. S1 and Table S2.

<b>Measurements</b>	SABAP UMB 19. 2.1178	SABAP UMB 19. 2.1179	SABAP UMB 20. 1.6803	SABAP UMB 20. 1.6802	SABAP UMB 21. 4.1177
<b>HS°</b>	65.0	75	-	-	-
<b>IBW</b>	<i>102.0</i>	<i>115.0</i>	-	-	-
<b>IBH</b>	<i>28.0</i>	<i>22.0</i>	-	-	-
<b>CW</b>	-	116.0	-	-	101.0
<b>FMH</b>	-	36.0	-	-	35.0
<b>FMW</b>	-	34.0	-	-	39.9
<b>MW</b>	-	202.0	-	-	217.4
<b>OSHmin</b>	-	<i>90.0</i>	-	-	-
<b>OSHmax</b>	-	<i>130.0</i>	-	-	-
<b>PW</b>	<i>146.0</i>	-	-	-	-
<b>BAW</b>	<i>90.0</i>	-	-	-	-
<b>HLmax</b>	-	-	<i>340.0</i>	<i>310.0</i>	-
<b>HAP</b>	-	<i>90.0</i>	-	-	-
<b>HDV</b>	-	<i>70.0</i>	<i>77.0</i>	-	-
<b>HLmin</b>	<i>206.0</i>	<i>226.0</i>	<i>223.0</i>	<i>215.0</i>	-
<b>HBWmax</b>	<i>326.0</i>	<i>335.0</i>	-	-	-
<b>HBWmin</b>	<i>250.0</i>	<i>270.0</i>	-	-	-
<b>FFW</b>	<i>144.0</i>	-	-	-	-
<b>HOD</b>	<i>77.0</i>	-	-	-	-
<b>FWmin</b>	<i>259.0</i>	-	-	-	-
<b>FFLmin</b>	<i>231.0</i>	-	-	-	-
<b>FLmax</b>	<i>157.0</i>	-	-	-	-
<b>FLmin</b>	<i>157.0</i>	-	-	-	-
<b>PPW</b>	58.0	67.0	69.0	59.0	66.0
<b>APW</b>	33.0	33.0	38.0	34.0	36.0

**Table 3.** Measurements (mm) of the metacarpals of *Bison (Eobison) degiulii* from Pietrafitta. Estimated measurements are in italics. Abbreviations as in Fig. S1 and Table S2.

ID Specimen	Side	Lmax	Lf	PEW	PET	DW	DT	DDW	DEW	DEtm	DETI	PFWm	PFWI	ABETm	ABETI	AETm	AETI
SABAP UMB 20. 1.7324	L	249.6	239.1	74.2	51.1	42.5	31.6	67.2	68.0	39.4	38.8	43.1	29.3	30.8	28.0	37.3	36.8
SABAP UMB 20. 1.7325	L	220.8	211.6	59.7	<i>40.8</i>	<i>36.7</i>	24.9	58.2	57.5	33.3	32.9	35.1	21.9	26.6	23.4	31.0	31.0
SABAP UMB 22. 4.247/7	L	<i>245.0</i>	-	66.4	<i>45.4</i>	42.5	30.0	61.8	64.0	35.7	35.7	-	-	27.6	26.1	33.8	33.1
SABAP UMB 20. 1.7329/1	R	<i>250.2</i>	-	-	-	<i>50.0</i>	-	<i>70.0</i>	<i>69.0</i>	<i>38.0</i>	-	-	-	<i>30.8</i>	<i>28.5</i>	<i>35.9</i>	<i>36.5</i>
SABAP UMB 22. 4.247/8	R	249.0	240.4	66.4	42.8	44.9	31.0	60.1	64.0	35.7	37.5	37.8	26.3	-	-	-	-
SABAP UMB 20. 1.7326	R	247.3	238.0	<i>73.1</i>	43.5	<i>44.6</i>	<i>30.5</i>	66.7	69.2	39.2	38.6	37.0	33.0	28.6	27.3	36.5	35.6
SABAP UMB 130038	R	218.9		60.2	38.9	37.0	26.0	57.8	57.7	33.5	33.3	34.7	24.9	25.8	24.8	31.0	30.8
IGF 1011	R	241.0	230.0	69.9	44.4	40.5	30.5	65.2	69.8	39.5	39.4	39.5	28.6	31.6	31.1	37.9	37.3

**Table 4.** Measurements (mm) of the metatarsals of *Bison (Eobison) degiulii* from Pietrafitta. Estimated measurements are in italics. Abbreviations as in Fig. S1 and Table S2.

ID Specimen	Side	Lmax	Lf	PEW	PET	DW	DT	DDW	DEW	DETm	DETI	PFOm	PFOI	ABETm	ABETI	AETm	AETI
SABAP UMB 20. 1.7330	R	258.2	248.2	47.8	48.2	34.8	33.3	54.2	52.9	33.8	33.0	33.9	31.8	24.9	24.1	29.8	31.2
SABAP UMB 20. 1.7331	R	284.3	270.5	58.8	55.0	36.9	36.0	63.0	61.5	37.6	35.0	36.6	34.0	28.5	23.0	33.5	33.3
SABAP UMB 20. 1.7332	R	295.3	283.9	64.5	56.8	42.8	36.9	-	63.3	38.5	36.9	40.1	-	28.4	27.1	33.3	34.7
SABAP UMB 20. 1.7333	L	254.3	-	-	-	37.0	-	56.8	57.4	33.5	35.0	43.7	-	25.1	26.7	32.7	30.6
SABAP UMB 20. 1.7335	R	288.0	-	-	-	-	-	-	-	40.0	37.6	-	-	28.7	28	36.2	36.4
SABAP UMB 20. 1.7336	R	296.0	-	-	-	-	-	62.3	63.0	38.5	37.4	41.4	37.6	29.2	26.7	33.9	35.6
SABAP UMB 20. 1.7338	R	-	-	-	-	-	-	63.3	65.6	41.2	40.0	-	-	30.6	29.3	36.1	37.9
SABAP UMB 20. 1.7328	R	250.3	-	53.2		34.5	35.7	57.1	55.7	34.3	33.8	-	-	26.1	25.9	30.8	31.7
SABAP UMB 20. 1.7327	R	-	-	-	-	-	-	66.0	71.3	39.1	39.0	-	-	31.5	28.5	36.8	37.9
SABAP UMB 130037	L	283.0	-	58.2	35.9	36.6	36.5	62.2	61.5	38.7	36.6	40.0	35.9	29.4	25.7	32.7	34.5
SABAP UMB 22. 4.247/15	R	278.9	-	59.5	54.8	38.2	40.6	59.7	58.7	35.5	35.1	39.5	39.8	26.3	25.4	31.6	33.9

**Table 5.** Comparative cranial measurements (mm) of selected *Leptobos* and *Bison* s.l. samples. Abbreviations of taxa: LF, *Leptobos falconeri*; LSEM, *Leptobos gr. stenometopon-elatus-merlai*; LE, *Leptobos etruscus*; LV, *Leptobos vallisarni*; AC, *Adjiderebos cantabilis*; EP, *Bison (Eobison) palaeosinensis*; EGDm, *Bison (Eobison) georgicus* from Dmanisi; EDPf, *Bison (Eobison) degiulii* from Pietrafitta; EDPr, *Bison (Eobison) degiulii* from Pirro Nord; EDKm, *Bison (Eobison) degiulii* from Kalamoto; EcfDSO, *Bison (Eobison) cf. degiulii* from Salita di Oriolo; BM, *Bison (Poephagus) menneri*; BS, *Bison (Bison) schoetensacki*; BP, *Bison (Bison) priscus* ssp. Estimated measurements are in italics. Abbreviations of measurements as in Fig. S1 and Table S2. Data taken from: <sup>a</sup>Masini (1989); <sup>b</sup>Dubrovo and Burchak-Abramovich (1986); <sup>c</sup>Teilhard de Chardin Piveteau (1930); <sup>d</sup>Bukshianidze (2005); <sup>e</sup>this work; <sup>f</sup>Kostopoulos et al. (2018); <sup>g</sup>Bukshianidze (2020); <sup>h</sup>Sala (1986); <sup>i</sup>Anfossi et al., 1999; <sup>j</sup>Prat et al. (2003); <sup>k</sup>Vercoutère and Guerin (2010); <sup>l</sup>Castañes et al. (2012).

Meas.	Stat.	LSEM <sup>a</sup>	LE <sup>a</sup>	LV <sup>a</sup>	AC <sup>b</sup>	EP <sup>a,c</sup>	EGDm <sup>d,e</sup>	EDPf <sup>e</sup>	EDPr <sup>a,e</sup>	EDKm <sup>f</sup>	EcfDSO <sup>e</sup>	BM <sup>g</sup>	BS <sup>e,h</sup>	BP <sup>e,h,i,j,k,l</sup>
IBW	Mean (N)	61.5 (5)	66.2 (13)	86.36 (5)	65.0 (1)	92.0 (2)	94.0 (1)	108.5 (2)	106.0 (1)	111.0 (1)	122.0 (1)	100.8 (1)	146.0 (1)	172.7 (25)
	Min-Max	48.4-77.0	41.0-95.0	78.5-99.0	-	90.0-94.0	-	102.0-115.0	-	-	-	-	-	125-230
IBH	Mean (N)	48.8 (5)	30.6 (13)	27.32 (5)	30 (1)	25.8 (1)	26.0 (1)	25.0 (2)	29.0 (1)	17.0 (1)	20.0 (1)	37.0 (1)	28.0 (1)	60.3 (16)
	Min-Max	44.0-59.0	22.5-40.0	23.4-32	-	-	-	22-28	-	-	-	-	-	13-147
CW	Mean (N)	92.4 (3)	89.5 (11)	93.75 (4)	106 (1)	100.3 (2)	99.8 (1)	108.4 (2)	-	-	118.0 (1)	110.2 (1)	136.7 (11)	135.3 (39)
	Min-Max	84.5-103.0	56.5-101	88.3-97	-	100-100.5	-	101.3-115.5	-	-	-	-	129.0-145.0	107.5-156.0
FMH	Mean (N)	35.5 (4)	35.8 (10)	36.6 (3)	-	42.5 (1)	30.0 (1)	35.5 (2)	-	-	46.0 (1)	-	40.3 (11)	44.0 (37)
	Min-Max	30.3-41.4	32.0-41.0	31.9-40.0	-	-	-	35.0-36.0	-	-	-	-	30.0-52.0	33.6-53.0
FMW	Mean (N)	42.9 (4)	36.4 (10)	38.2 (3)	-	42.8 (1)	33.0 (1)	36.95 (2)	-	-	43 (1)	40.7 (1)	44 (11)	41.9 (37)
	Min-Max	37.5-46.7	28.9-39.4	35.1-42.1	-	-	-	34.0-39.9	-	-	-	-	36.0-52.0	19.6-59.0
MW	Mean (N)	162.7 (4)	193.5 (8)	177.96 (5)	220.0 (1)	217.0 (2)	200.0 (1)	209.7 (2)	204 (1)	165 (1)	246 (1)	-	269.1 (7)	274.4 (35)
	Min-Max	131.9-200.0	155.3-220	145-200.8	-	210.0-224.0	-	202-217.4	-	-	-	-	208-322	218-350
OSHmin	Mean (N)	101.9 (4)	90.9 (12)	92.1 (2)	-	94.8 (1)	100.0 (1)	90.0 (1)	90.7 (1)	-	120.0 (1)	97.4 (1)	104 (8)	122.8 (30)
	Min-Max	96.8-106.8	75.4-103.0	91.2-93.0	-	-	-	-	-	-	-	-	72.0-130.0	93.0-151.0
OSHmax	Mean (N)	134.5 (4)	128 (8)	108.5 (4)	102.0 (1)	131.5 (2)	134.0 (1)	130.0 (1)	-	-	143.0 (1)	131.7 (1)	136.0 (8)	160.9 (31)
	Min-Max	125.8-141.0	115.6-141.0	85.0-130.5	-	129.0-134.0	-	-	-	-	-	-	120-164	132.6-188.0
PW	Mean (N)	125.3 (5)	144.4 (11)	155.5 (2)	-	208.0 (1)	143.0 (1)	146.0 (1)	177.0 (1)	-	205.0 (1)	-	233.0 (1)	239.4 (16)
	Min-Max	94.0-161.0	85.5-220.0	151.0-160.0	-	-	-	-	-	-	-	-	-	205.0-259.0
BAW	Mean (N)	86.5 (5)	87.1 (10)	70.375 (4)	-	55.3 (2)	63.0 (1)	90.0 (1)	-	48.0 (1)	-	47.6 (1)	74.0 (1)	98.6 (17)
	Min-Max	80.0-99.0	61.0-101.6	53.0-78.0	-	50.0-60.5	-	-	-	-	-	-	-	69.0-149.0
HLmax	Mean (N)	435.0 (1)	490.0 (3)	322.0 (3)	280.0 (1)	450.0 (1)	344.0 (1)	325.0 (2)	435.0 (1)	-	-	-	344.7 (9)	440.6 (9)
	Min-Max	-	440.0-540	270.0-360.0	-	-	-	310.0-340.0	-	-	-	-	310.0-420.0	285.0-570.0
HAP	Mean (N)	79.3 (3)	82.2 (4)	80.3 (3)	65.0 (1)	85 (1)	92.1 (1)	90 (1)	75 (1)	58 (1)	106 (1)	67.8 (1)	105.1 (8)	110.2 (10)
	Min-Max	72.0-87.8	79.5-84.5	70-86	-	-	-	-	-	-	-	-	91.0-116.0	74.5-140.0

<b>HDV</b>	<b>Mean (N)</b>	62.0 (3)	75.9 (4)	70.8 (3)	85.0 (1)	78.2 (1)	73.6 (1)	73.5 (2)	60.0 (1)	50.0 (1)	80.0 (1)	49.4 (1)	75.1 (8)	108.2 (10)
	<b>Min-Max</b>	59.0-65.0	71.5-79.0	63.0-77.5	-	-	-	70.0-77.0	-	-	-	-	58-87	77.5-136.5
<b>HLmin</b>	<b>Mean (N)</b>	245.0 (1)	374.3 (3)	275.5 (3)	173.0 (1)	210 (2)	245 (1)	217.5 (4)	-	-	-	159.9 (1)	276 (10)	395.1 (8)
	<b>Min-Max</b>		343.0-400.0	250.0-296.5	-	200.0-220.0		206.0-226.0	-	-	-	-	235.0-325.0	255.0-505.0
<b>HBWmax</b>	<b>Mean (N)</b>	238.5 (2)	306.3 (3)	303.3 (1)	-	317.0 (1)	294.3 (1)	330.5 (2)	245.0 (1)	-	-	-	314.1 (10)	-
	<b>Min-Max</b>	232.0-245.0	283.0-320.0	-	-	-	-	326.0-335.0	-	-	-	-	280.0-350.0	-
<b>HBWmin</b>	<b>Mean (N)</b>	135.5 (2)	268 (3)	202 (1)	-	247.5 (2)	195 (1)	260 (2)	153 (1)	-	310 (1)	-	-	314.4 (6)
	<b>Min-Max</b>	118.0-153.0	224.0-350.0	-	-	220.0-275.0	-	250.0-270.0	-	-	-	-	-	236.5-390
<b>FFW</b>	<b>Mean (N)</b>	102.4 (6)	105.7 (8)	126.9 (3)	-	152 (2)	127.4 (1)	144 (1)	134.3 (1)	122.0 (1)	132.0 (1)	142.1 (1)	-	195.8 (15)
	<b>Min-Max</b>	84.0-123.0	77-159	107.6-144		143.0-161.0	-	-	-	-	-	-	-	180.5-217.0
<b>HOD</b>	<b>Mean (N)</b>	93.8 (2)	102.8 (2)	92.0 (1)	75.0 (1)	-	72.0 (1)	77.0 (1)	65.0 (1)	60.0 (1)	90.0 (1)	-	-	97.1 (18)
	<b>Min-Max</b>	85.0-102.5	95.0-110.5	-	-	-	-	-	-	-	-	-	-	73.0-111.0
<b>FWmin</b>	<b>Mean (N)</b>	159.2 (5)	205.3 (7)	249.5 (3)	210.0 (1)	237.0 (1)	241.8 (1)	259.0 (1)	239.5 (1)	202.0 (1)	244.0 (1)	223.5 (1)	290.9 (11)	296.6 (37)
	<b>Min-Max</b>	132.0-201.0	133.7-267.4	223.4-280.0	-	-	-	-	-	-	-	-	255.0-335.0	232.0-344.0
<b>FFLmin</b>	<b>Mean (N)</b>	134.3 (3)	150.2 (8)	166.0 (1)	-	150.0 (1)	154.5 (1)	231.0 (1)	182.9 (1)	175.0 (1)	-	-	-	208.4 (15)
	<b>Min-Max</b>	105.0-153.0	132.1-174.4	-	-	-	-	-	-	-	-	-	-	167.0-249.0
<b>FLmax</b>	<b>Mean (N)</b>	234.1 (2)	226.2 (3)	-	-	-	223.0 (1)	157.0 (1)	230.0 (1)	-	225.0 (1)	214.4 (1)	-	274.6 (16)
	<b>Min-Max</b>	231.2-237.0	196-282	-	-	-	-	-	-	-	-	-	-	233.0-303.0
<b>FLmin</b>	<b>Mean (N)</b>	148.0 (3)	160.9 (4)	157.7 (3)	130.0 (1)	-	147.5 (1)	157.0 (1)	194.0 (1)	175.0 (1)	159.0 (1)	166.8 (1)		196.0 (12)
	<b>Min-Max</b>	133.0-173.5	126.4-202	150-170	-	-	-	-	-	-	-	-	-	145.4-244.0
<b>PPW</b>	<b>Mean (N)</b>	58.3 (6)	64.5 (7)	65.5 (5)	-	65.5 (1)	61.8 (1)	63.84 (5)	-	38 (1)	-	34.1 (1)	-	73.8 (20)
	<b>Min-Max</b>	50.3-63.0	49.7-76.3	62.0-69.5	-	-	-	58-68.9	-	-	-	-	-	55.5-90.0
<b>APW</b>	<b>Mean (N)</b>	29.7 (6)	39.8 (6)	34.9 (5)	-	42.5 (1)	33.6 (1)	34.68 (5)	-	-	-	-	-	46.8 (17)
	<b>Min-Max</b>	24.8-36.6	35.0-55.5	29.0-43.6	-	-	-	33.0-38.0	-	-	-	-	-	32.0-56.6

**Table 6.** Comparative measurements (mm) of the metacarpal in selected *Leptobos* and *Bison* s.l. samples. Abbreviations as in Fig. S1 and Table S2. Data taken from: <sup>a</sup>Masini (1989); <sup>b</sup>Rodrigo (2011); <sup>c</sup>Garrido (2008); <sup>d</sup>Tong et al. (2016); <sup>e</sup>Bukshianidze (2005); <sup>f</sup>this work; <sup>g</sup>Maniakas and Kostopoulos (2017); <sup>h</sup>Kostopoulos et al. (2018); <sup>i</sup>Sher (1997); <sup>j</sup>Brugal (1994); <sup>k</sup>Moullé (1992); <sup>l</sup>Brugal and Fosse (2005); <sup>m</sup>Schertz (1936); <sup>n</sup>Sala (1986); <sup>o</sup>Sorbelli et al. (2021a); <sup>p</sup>Breda et al. (2010); <sup>q</sup>Vercoutère and Guerin (2010); <sup>r</sup>Prat et al. (2003); <sup>s</sup>Drees (2005); <sup>t</sup>Castaños et al. (2012); <sup>u</sup>Empel and Roskoz (1963).

Taxon	Site	Statistics	Lmax	PEW	PET	DW	DT	DEW	DEtm
<i>Leptobos elatus</i> <sup>a, b</sup>	Les Étouaires, Villarroya (France, Spain)	Mean (N)	253.6 (25)	57.5 (32)	37.8 (27)	33.9 (20)	26.1 (18)	55.1 (23)	34.5 (23)
		Min-Max	242.2-270	52.0-64.7	32.1-42.9	29.6-39.4	22.0-30.0	49.1-63.6	31.5-36.9
		SD	8.1	3.7	2.9	3.0	2.2	3.8	1.6
<i>Leptobos merlai</i> <sup>a, b</sup>	St. Vallier (France)	Mean (N)	245.5 (8)	58.2 (13)	39.2 (14)	36.8 (7)	27.1 (7)	57.5 (14)	33.8 (13)
		Min-Max	231.6-255.8	51.1-62.5	34.5-46.4	32.5-40.0	23.4-29.1	51.9-62.9	30.3-37.4
		SD	9.2	3.9	3.6	3.0	2.1	3.7	2.1
<i>Leptobos etruscus</i> <sup>a, c</sup>	Senèze, Faella, Matassino, Olivola, Fonelas-1 (France, Italy, Spain)	Mean (N)	248.3 (32)	58.1 (40)	45 (40)	39.3 (37)	28.5 (37)	62.6 (49)	35.7 (43)
		Min-Max	234.5-265.5	36.2-67.9	36.7-65.6	31.8-43.7	25.0-31.5	50.4-68.6	32.0-41.1
		SD	6.9	9.5	9.2	3.0	1.6	3.9	1.7
<i>B. (Eobison) palaeosinensis</i> <sup>d</sup>	Nihowan Basin (China)	Mean (N)	234.8 (4)	68.4 (4)	40.4 (3)	39.4 (1)	30.6 (1)	63.5 (4)	34.6 (3)
		Min-Max	226.0-248.1	60.0-73.8	35.0-47.5	-	-	58.8-65.2	34.2-34.9
		SD	9.4	6.1	6.4	-	-	3.1	0.4
<i>B. (Eobison) georgicus</i> <sup>e, f</sup>	Dmanisi (Georgia)	Mean (N)	237.0 (3)	66.9 (4)	42.4 (4)	39.9 (3)	29.2 (3)	65.3 (4)	35.4 (4)
		Min-Max	230.0-244.7	62.8-70.8	37.7-45.2	34.3-44.6	25.8-31.1	61.0-71.6	33.2-37.8
		SD	7.3	3.6	3.3	5.2	3.0	4.5	2.0
<i>B. (Eobison) sp.</i> <sup>g</sup>	Venta Micena (Spain)	Mean (N)	225.6 (30)	58.5 (43)	36.6 (44)	35.3 (31)	26.9 (30)	60.1 (36)	33.0 (35)
		Min-Max	208.2-248.7	52.3-69.5	32.8-46.5	31.4-44.2	24.1-32.1	53.5-69.1	30-36.3
		SD	10.2	4.6	3.4	3.7	2.1	4.7	1.6
<i>B. (Eobison) degiulii</i> <sup>e</sup>	Pietrafitta (Italy)	Mean (N)	240.0 (8)	67.1 (7)	43.8 (7)	42.3 (8)	29.5 (7)	65.0 (8)	36.8 (8)
		Min-Max	218.9-250.2	59.7-74.2	38.9-51.1	36.7-50.0	24.9-32.9	57.5-70	33.3-39.5
		SD	12.8	5.7	3.9	4.4	2.9	5.2	2.6
<i>B. (Eobison) degiulii</i> <sup>a, e</sup>	Pirro Nord (Italy)	Mean (N)	222.2 (3)	62.1 (4)	38.7 (4)	38.5 (3)	27.4 (3)	64.8 (4)	33.1 (5)
		Min-Max	211.5-230.0	58.0-69.5	35.5-41.7	29.8-46.2	25-30.1	62.0-66.4	29.3-37.3
		SD	9.6	5.3	3.2	8.3	2.6	1.9	3.2
<i>B. (Eobison) cf. degiulii</i> <sup>e</sup>	Capena	Mean (N)	222.5 (2)	68.4 (2)	43.1 (2)	44.4 (2)	29.7 (1)	64.6 (2)	35.4 (2)

	(Italy)	<b>Min-Max</b>	220.0-225.0	68.1-68.7	42.0-44.0	44.3-44.4	-	64.1-65.0	34.1-36.7
		<b>SD</b>	3.5	0.4	1.4	0.1	-	0.6	1.8
<i>B. (Eobison) degiulii</i> <sup>h</sup>	Mygdonia Basin	<b>Mean (N)</b>	238.5 (19)	67.4 (24)	40.1 (24)	42.4 (23)	29.4 (23)	70.3 (22)	39.3 (22)
	(Greece)	<b>Min-Max</b>	220.7-258.1	57.7-78.3	32.9-48.4	31.8-51.2	23.5-35.2	59.5-88.8	34.4-48.2
		<b>SD</b>	10.6	4.9	3.0	5.1	2.9	6.7	3.0
<i>B. (Eobison) cf. degiulii</i> <sup>e</sup>	Capena	<b>Mean (N)</b>	222.5 (2)	68.4 (2)	43.1 (2)	44.4 (2)	29.7 (1)	64.6 (2)	35.4 (2)
	(Italy)	<b>Min-Max</b>	220.0-225.0	68.1-68.7	42.0-44.0	44.3-44.4	-	64.1-65.0	34.1-36.7
		<b>SD</b>	3.5	0.4	1.4	0.1	-	0.6	1.8
<i>B. (Poephagus) menneri</i> <sup>f, i</sup>	Untermassfeld	<b>Mean (N)</b>	276.4 (37)	77.0 (50)	46.5 (48)	44.8 (43)	32.8 (40)	75.6 (41)	38.2 (38)
	(Germany)	<b>Min-Max</b>	255.8-296.0	64.37-91.4	39.27-52.6	35.3-57.7	27.8-39.0	65.4-86.0	29.5-47.3
		<b>SD</b>	10.2	5.6	3.2	4.7	3.0	6.3	5.8
<i>B. (Bison) schoetensacki</i> f, g, j, k, l, m, n, o, p	Durfort, Le Vallonnet, Le Vassirie, Mauer, Süßenborn, Cesi, Vallparadís, Cromer Forest-bed	<b>Mean (N)</b>	252.2 (54)	78.4 (68)	46.8 (61)	49.3 (49)	33.9 (51)	76.9 (67)	42.1 (54)
	(France, Germany, Italy, UK)	<b>Min-Max</b>	225.0-277.0	46.9-93.0	37.6-56.0	36.3-64.0	27.9-46.6	65.9-88.5	35.3-52.4
		<b>SD</b>	12.8	7.0	3.6	5.7	3.9	6.0	3.3
<i>B. (Bison) cf. schoetensacki</i> <sup>f, m</sup>	Mosbach	<b>Mean (N)</b>	256.0 (22)	84.5 (22)	48.9 (22)	49.9 (22)	32.6 (22)	82.9 (22)	42.8 (22)
	(Germany)	<b>Min-Max</b>	241.1-277.8	72.0-97.3	43.4-56.4	43.3-55.1	30.2-35.1	73.0-97.1	40.1-46.0
		<b>SD</b>	22.0	18.0	19.0	4.0	4.0	19.0	4.0
<i>B. (Bison) priscus priscus</i> <sup>e, g, q</sup>	Chatillon-Saint-Jean, Romain-La-Roche, Taubach	<b>Mean (N)</b>	254.5 (40)	87.9 (39)	53.1 (38)	55.7 (38)	36.3 (38)	87.9 (23)	46.1 (40)
	(France, Germany)	<b>Min-Max</b>	232.0-286.5	71.4-102.0	44.0-63.0	42.3-63.0	30.6-42.5	75.5-98.1	37.0-51.5
		<b>SD</b>	13.0	8.1	5.2	6.1	3.6	6.9	3.1
<i>B. (Bison) priscus</i> g, m, n, r, s, t	Habarra, Roter Berg, Cava Filo, North Sea, Kiputz IX, Joint Mitnor Cave	<b>Mean (N)</b>	234.5 (58)	81.5 (66)	48.2 (31)	50.3 (40)	33.4 (29)	84.7 (56)	42.6 (50)
	(France, Germany, Italy, North Sea, Spain, UK)	<b>Min-Max</b>	216.8-260.0	60.0-96.5	42.5-56.0	39.5-59.0	28.8-40.0	70.5-101.5	31.1-53.0
		<b>SD</b>	10	7.1	3.5	5.9	2.8	7.1	5.8
<i>B. (Bison) bonasus</i> <sup>u</sup>	Białowieża	<b>Mean (N)</b>	213.9 (32)	73.5 (31)	45.7 (30)	42.9 (31)	27.1 (32)	69.9 (32)	39.7 (30)
	(Poland)	<b>Min-Max</b>	202.0-231.0	64.0-88.0	39.0-54.0	35.0-55.0	23.0-32.0	61.0-80.0	30.0-46.0
		<b>SD</b>	7.6	6.6	3.8	5.7	2.5	5.5	3.3

**Table 7.** Summary list of morphological characters that define the four main groups of *Leptobos* and *Bison* studied in this work.

	<i>Leptobos</i> gr. LSEM	<i>Leptobos</i> gr. LEV	<i>Bison</i> ( <i>Eobison</i> )	<i>Bison</i> ( <i>Bison</i> )
<b>Horn-core shape</b>	Long with double bending and strong torsion. Markedly compressed dorsoventrally at the base.	Long with single bending and feeble torsion. Circular section at the base.	Short and stout with double bending and feeble torsion. Dorsoventrally compressed at the base.	From short and stout to very long. Single or double bending with feeble torsion. Circular to dorsoventrally compressed at the base.
<b>Horn-core position</b>	Emerging more laterally than posteriorly (i.e., forming a wide angle with the sagittal axis). Pedicle long.	Emerging more posteriorly than laterally (i.e., forming a small angle with the sagittal axis). Pedicle long.	Emerging more laterally than posteriorly (i.e., forming a wide angle with the sagittal axis). Pedicle long to short.	Emerging laterally (i.e., forming a very wide angle with the sagittal axis). Pedicle short.
<b>Occipital squama</b>	Small. Bell shaped, wide base and narrow top.	Trapezoidal to semicircular shaped, wide base and narrow top.	Trapezoidal to bell shaped, wide base and narrow top.	Trapezoidal to square shaped, wide base and top. Dorsoventrally compressed.
<b>Intertemporal bridge</b>	Well developed. High and narrow.	Poorly developed. Short and wide.	Poorly developed to absent.	Absent.
<b>Temporal fossae</b>	High, developing posteriorly. The distance between the two posterior ends is short.	High, developing posteriorly. The distance between the two posterior ends is relatively short.	Relatively low, developing posteriorly. The distance between the two posterior ends is relatively long.	High, developing posteriorly. The distance between the two posterior ends is very long.
<b>Frontals</b>	Relatively small, long and narrow. Concave, poorly pneumatized at the horn-core bases.	Relatively large, long and narrow. Concave, poorly pneumatized at the horn-core bases.	Relatively large, short and wide. Concave, poorly pneumatized at the horn-core bases.	Large, very short and wide. Concave to convex, strongly pneumatized at the horn-core bases.
<b>Ethmoidal fenestrae</b>	Present.	Present.	Absent.	Absent.
<b>Orbits</b>	Not tubular nor protruding.	Slightly protruding, not tubular.	Protruding and slightly tubular.	Extremely protruding and tubular.
<b>Nasals</b>	Short, ending anteriorly to the anterior margin of the orbits.	Short, ending anteriorly to the anterior margin of the orbits.	Long, ending at the anterior margin of the orbits or slightly posteriorly.	Long, ending posteriorly the anterior margin of the orbits.
<b>Premaxillae</b>	In contact with the nasals.	In contact with the nasals.	Not in contact with the nasals.	Not in contact with the nasals.
<b>Teeth</b>	Low degree of hypsodonty.	Medium degree of hypsodonty.	Medium degree of hypsodonty.	High degree of hypsodonty.
<b>Limbs</b>	Very slender. Radius short. Metatarsals slightly longer than metacarpals.	Slender. Radius relatively short. Metatarsals slightly longer than metacarpals.	Relatively robust. Radius relatively short. Metatarsals relatively longer than metacarpals.	To relatively to extremely robust. Radius long. Metatarsals from relatively to extremely longer than metacarpals.
<b>Body size</b>	Small.	Medium.	Medium to large.	Large to very large.

---

**Sexual dimorphism**

Female hornless. Size difference between male and females reduced.

Female hornless. Size difference between males and females reduced.

Female horned. Size difference between male and females accentuated.

Female horned. Size difference between male and females extremely accentuated.

---

**Declaration of interests**

The authors declare that they have no known competing financial interests or personal relationships that could have appeared to influence the work reported in this paper.

The authors declare the following financial interests/personal relationships which may be considered as potential competing interests:

**Author contributions**

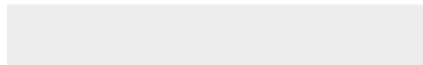
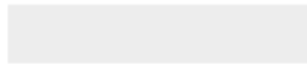
L.S, M.C and J.M.-M conceived the research and wrote the paper; L.S, performed the analyses; L.S, M.C, D.S.K., R.F., B.M., V.P., M.P.V., B.A., S.V.B. and J.M.-M. collected the data, revised all the manuscript versions and made improvements in the manuscript.



Click here to access/download

**e-Component/Supplementary data**

Supplementary material\_Sorbelli\_et\_al.docx













Click here to access/download  
**e-Component/Supplementary data**  
Fig.S5.tif





Click here to access/download  
**e-Component/Supplementary data**  
Fig.S6.tif

