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Keystone bacterial functional module activates P-mineralizing genes to enhance enzymatic hydrolysis of organic P in a subtropical forest soil with 5-year N addition

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Abstract

Microorganisms play an integral role in driving P transformation in forest soils; however, studies on soil P cycling and the molecular mechanisms of microbes activated in response to elevated nitrogen (N) deposition are limited. In this study, we conducted a multilevel field N enrichment experiment in a subtropical P-deficient Moso bamboo (*Phyllostachys heterocycla*) system to evaluate the microbial ecological traits of P transformation (e.g., organic P mineralization and inorganic P solubilization) over three consecutive years. N addition significantly decreased available and organic P levels in the soil and increased the microbial biomass C:P and N:P ratios, indicative of severe microbial P limitation. Consequently, N addition increased the absolute abundance of P starvation response regulation genes (*phoU* and *phoR*), which further induced an increase in organic P mineralization (*phoN*, *phoD*, *appA*), but not that of inorganic P solubilization genes (*ppx* and *gcd*). This suggests that microbes enhance P availability by organic P mineralization rather than inorganic P solubilization to ameliorate reduced P availability. Furthermore, a bacterial functional module (B_Mod#0) consisting of *Proteobacteria*, *Actinobacteria*, and *Firmicutes* accounted for more than 60% of the changes in the abundance of genes responsible for organic P-mineralization in the soil, suggesting that B_Mod#0 acts as a keystone phylotype in enhancing functional P-cycling potential. This study provides novel insights into microorganism-driven P cycling in P-deficient forest soils with N addition.

Keywords: Nitrogen deposition, Phosphorus limitation, Microbial modules, Functional P-cycling potential, Phosphorus mineralization

1. Introduction

Over the past century, the increase in N deposition has resulted in a global N/P imbalance caused by the combustion of fossil fuels, application of fertilizers, and expansion of N₂-fixing crops (Davidson, 2009; Peñuelas et al., 2013; Peñuelas and Sardans, 2022). This imbalance in nutrient stoichiometry threatens terrestrial ecosystems and has received great attention (Peñuelas and Sardans, 2022). Increasing evidence from studies reporting increased N:P ratios in soils and microbial biomass, accompanied by decreased levels of available P and microbial biomass P in the soil caused by N addition, supports the hypothesis that N deposition exacerbates P limitation in terrestrial ecosystems (Deng et al., 2017; Fan et al., 2020; Li et al., 2016; Wang et al., 2023). However, N deposition could partly increase P supply by enhancing plant and microbial phosphatase activity (Chen et al., 2020). These contrasting findings indicate that the response of soil P availability to N deposition remains uncertain, highlighting the necessity to determine the underlying mechanisms of such discrepancies and assess the effects of N deposition on forest ecosystems at the global and regional scales.

Microorganisms play a fundamental role in soil P cycling and regulating P availability in forest ecosystems (Liang et al., 2020; Richardson and Simpson, 2011). Low P availability may considerably affect soil community composition by strengthening the interaction between certain microbial taxa with common habitats and functions, leading to the formation of microbial ecological modules capable of ameliorating P limitation (Fan et al., 2021a). A recent study used co-occurrence network analysis to construct a microbial ecological module and identify the keystone phylotypes that regulate P cycling in agroecosystems under long-term

fertilization (Fan et al., 2021a). Forest ecosystems are more prone to P limitation than agroecosystems because the former lack artificial fertilization, and the weathering of P-containing rocks is the key P source (Reinhard et al., 2017; Walker and Syers, 1976). Microbes play an essential role in P transformation in forest soils (Dai et al., 2020; Liang et al., 2020; Richardson and Simpson, 2011). However, information regarding the dominant ecological module phylotypes and their roles in P transformation remains limited.

Microorganism-mediated soil P transformation is mainly attributed to four microbial gene categories: genes involved in P starvation response regulation (e.g., *phoU* and *phoR*), inorganic P solubilization (e.g., *gcd*), organic P mineralization (e.g., *phoD* and *phoA*), and P uptake and transport (e.g., *pst* and *pit*) (Bergkemper et al., 2016; Liang et al., 2020). The *phoR-phoB* two-component regulatory system detects extracellular changes in P concentration and regulates the expression of genes involved in P cycling. Upon encountering P limitation, the expression of the sensor *phoR* activates the response regulator *phoB*, which consequently activates P-starvation response regulation, enabling microorganisms to use alternative P sources (Hsieh and Wanner, 2010; Lubin et al., 2016; Wanner, 1993). Although this paradigm of microbial molecular mechanisms affecting P utilization has been demonstrated in *Escherichia coli* *in vitro* (Hsieh and Wanner, 2010), further studies are required to determine whether two-component regulatory systems mediate P cycling *in situ* (forest soil).

Microbial functional genes are expected to regulate inorganic P solubilization and organic P mineralization processes in the soil when P availability is limited (Alori et al., 2017; Dai et al., 2020). Liang et al. (2020) reported that the enhanced levels of available P found in mining degraded land after ecological restoration are mainly attributed to the solubilization of

inorganic P driven by enhanced *gcd* gene expression. Another study revealed that N addition increased organic P mineralization because of the upregulation of the *phoD* gene (phosphatase) in microbes (Wang et al., 2018a). Thus, microbes may trigger different functional genes to regulate P availability in soils with varying accessibility to inorganic or organic P. Nevertheless, in subtropical forest soils with a high phosphate sorption capacity (Vitousek and Farrington, 1997), the group of microbial functional genes that is predominantly involved in P transformation to counter the N deposition-induced soil P limitation remains unclear.

Owing to severe weathering and the immobilization of P by iron or aluminum oxides, subtropical forest soils are characterized by low P availability (Walker and Syers, 1976; Zhou et al., 2018). Furthermore, N deposition stimulates plant growth and can further increase P demand (Fan et al., 2018), thereby intensifying the gap between P supply and demand in the soil. Therefore, evaluating soil P availability in response to N deposition and the underlying mechanisms are critical for predicting the productivity of subtropical forest ecosystems. In the present study, Moso bamboo (*Phyllostachys heterocycla* cv. *Pubescens*), one of the most important non-timber forest plants in the world (Peng et al., 2013), was selected to conduct a simulated N deposition experiment to explore the mechanism of microorganism-driven P cycling in soil. We hypothesized that: 1) N deposition may induce stoichiometric imbalance of soil microbial biomass and thus enhance microbial function in mobilizing soil P; 2) microbes mainly upregulate organic P mineralization rather than inorganic P solubilization genes to alleviate P limitation because organic P is a major P source in subtropical forest soil.

2. Material and methods

2.1. Site description and sampling

This study was conducted in the Daiyun Mountain Nature Reserve located in Dehua County, Fujian Province, southeast China (118°05'22"–118°20'15" E, 25°38'07"–25°43'40" N; 1200 m above sea level). This region is characterized by a subtropical humid monsoon climate (Zeng et al., 2022), which can be simultaneously hot and humid, with minimum and maximum average temperatures of 8.5 and 25.5 °C in January and July (30-year average), respectively. The average annual precipitation is 1,800 mm, primarily occurring during the monsoon period (March–September), while the average relative humidity is 80%. The soil in this region is classified as Ultisol according to the USDA Soil Taxonomy. The experimental site consisted of a pure Moso bamboo forest with few trees or shrubs where the bamboo trees were mostly regenerated during 2015–2016.

Nine 3 × 10 m experimental plots were randomly selected in the study area. These plots had similar slopes and elevations and were facing the same direction. Three different levels of N were applied to the plots: control (CT, 0 kg N ha⁻¹·yr⁻¹), low N application (LN, 20 kg N ha⁻¹·yr⁻¹), and high N application (HN, 80 kg N ha⁻¹·yr⁻¹). Each treatment had three replicates, resulting in a total of nine experimental plots. The experiment was initiated in July 2014. For the LN and HN plots, appropriate amounts of NH₄NO₃ were dissolved in 20 L of deionized water. The solution was sprayed evenly on these plots monthly during the growing season (March to September). The CT plots received an equivalent volume of deionized water. No other cultivation activities, such as fertilizing with other nutrients, weeding, or felling, were performed during the experimental period.

Leaf sampling was conducted following the methods described previously (Li et al., 2021a). In July 2019, 20 healthy leaves from the mid-upper canopy on the south-facing side of each bamboo plant were collected. These leaves were mixed to create a pooled sample for each plot. Immediately after collection, the leaves were flash-frozen in dry ice and stored at -80 °C during transportation to the laboratory. Subsequently, the samples were dried at 105 °C for 30 min, followed by incubation at 65 °C until a constant weight was achieved. The dried samples were then milled before measuring the P concentration.

For soil sampling, five topsoil cores from a depth of 0 to 15 cm were randomly collected from each of the three replicate plots in July 2017, 2018, and 2019. After removing the litter layer, the soil samples were mixed and homogenized to create a pooled sample for each plot. A total of 27 soil samples were collected, sealed in plastic bags, and transported to the laboratory. The soil samples were sieved (< 2 mm) to remove visible roots and stones. The soil samples were divided into two parts. One part was stored at 4 °C for determining microbial characteristics and extracting DNA within two weeks. The other part was air-dried and used for the analysis of available, organic, and total P in the soil, as well as other soil properties. Fine roots (0–2 mm in diameter) of Moso bamboo were carefully cleaned and collected. The root samples were dried at 105 °C for 30 min, then dried to a constant weight at 65 °C. They were then ground and sieved through a 0.149 mm mesh sieve.

2.2. Soil Chemical Analysis

Foliar and root P concentrations were evaluated by digesting the samples with H₂SO₄ and HClO₄ (4:1) (Xu et al., 2017) and measured using a continuous flow analyzer (San++; Skalar,

the Netherlands). Soil pH was measured in a 1:2.5 (w/v) soil-to-water extract using a pH meter (STARTER300; OHAUS, USA). Soil organic C (SOC) and total N (TN) were determined using an elemental analyzer (Elementar Vario EL III; Elementar Analysensysteme GmbH, Germany). The NH_4^+ -N and NO_3^- -N were extracted with 2 mol L⁻¹ KCl, and determined using a continuous-flow analytical system (Skalar San++, Breda, the Netherlands) (Carter and Gregorich, 2007). Mineral N (MN) was calculated as the summary of NH_4^+ -N and NO_3^- -N. Soil available P was extracted using 0.5 M NaHCO₃ (Watanabe and Olsen, 1965), whereas soil organic P was extracted with 1 M H₂SO₄ after ignition at 500 °C for 1 h (Walker et al., 1958). Finally, the total soil P was evaluated using sulfuric acid and perchloric acid digestion. A continuous-flow analytical system was used to determine phosphate concentrations in the soil extract samples. Soil microbial biomass C, N, and P concentrations (MBC, MBN, and MBP, respectively) were measured using the chloroform fumigation–extraction method (Vance et al., 1987). Soil microbial biomass C, N, and P (MBC, MBN, and MBP, respectively) were determined using the chloroform fumigation–extraction method. The conversion factors for MBC, MBN, and MBP were 0.45, 0.54, and 0.40, respectively (Brookes et al., 1982; Li et al., 2021b; Vance et al., 1987). P nitrophenyl phosphate was used as a substrate to assay the potential activities of acid phosphatase (ACP) and alkaline phosphatase (ALP) in a modified universal buffer (pH = 5.0, pH = 11.0) (Saiya-Cork et al., 2002; Zhang et al., 2012). Furthermore, bis-*p*-nitrophenyl phosphate was used as a substrate to assay the potential activity of phosphodiesterase (PD) in a THAM-acetate buffer (pH = 8) (Tian et al., 2021). Phosphatase activity was measured using 365 nm excitation and 450 nm emission filters on a SpectraMax M5 Microplate Reader (MDS Analytical Technologies, Molecular Devices, USA).

2.3. DNA extraction, PCR amplification, and sequencing

Soil DNA was extracted using the E.Z.N.A.[®] soil DNA Kit (Omega Bio-tek, Norcross, GA, USA), according to the manufacturer's instructions. DNA was visualized on 1% agarose gels and measured using a NanoDrop 2000 UV-vis spectrophotometer (Thermo Scientific, Wilmington, USA).

The variable V3-V4 regions of the bacterial 16S rRNA gene were amplified using the primers 338F (ACTCCTACGGGAGGCAGCAG) and 806 R (GGACTACNNGGGTATCTAAT) (Caporaso et al., 2011). For fungi, the ITS1 region of fungal rRNA was amplified using ITS1 (CTTGGTCATTTAGAGGAAGTAA) and ITS2 (TGC GTTCTTCATCGATGC) primer sets (Wang et al., 2022a). PCR analysis was performed in triplicates and consisted of 20 μ L reaction mixture containing 4 μ L of 5 $\times$ FastPfu Buffer, 2 μ L of 2.5 mM dNTPs, 0.8 μ L of each primer (5 μ M), 0.4 μ L of FastPfu Polymerase, and 10 ng of DNA template. Sequencing was performed on an Illumina MiSeq platform (Illumina, San Diego, CA, USA), according to the standard protocols provided by Majorbio Bio-Pharm Technology Co., Ltd. (Shanghai, China). Metagenomic sequencing was performed as follow. DNA extract was fragmented to an average size of about 400 bp using Covaris M220 (Gene Company Limited, China) for paired-end library construction. Paired-end library was constructed using NEXTFLEX Rapid DNA-Seq (Bioo Scientific, Austin, TX, USA). Paired-end sequencing was performed on Illumina NovaSeq (Illumina Inc., San Diego, CA, USA) at Majorbio Bio-Pharm Technology Co., Ltd. (Shanghai, China) using NovaSeq 6000 S4 Reagent Kit v1.5 (300 cycles) according to the manufacturer's instructions.

2.4. Processing of sequencing data

Raw FASTQ files were quality-filtered by Trimmomatic and merged by FLASH using the following criteria: (i) reads with an average quality score < 20 over a 50 bp sliding window were truncated at any site; (ii) sequences overlapping more than 10 bp were merged according to their overlap with a mismatch of no more than 2 bp; and (iii) sequences of each sample were separated according to barcodes (exactly matching) and primers (allowing 2 nucleotide mismatching), and reads containing ambiguous bases were removed. Operational taxonomic units (OTUs) with $\geq 97\%$ similarity were clustered using UPARSE (version 7.1; <http://drive5.com/uparse/>; Edgar, 2013). Phylogenetic taxonomy was conducted using the RDP Classifier algorithm (<http://rdp.cme.msu.edu/>) against the SILVA database (Quast et al., 2013) for bacteria and UNITE database (Nilsson et al., 2019) for fungi, using a 70% confidence threshold. After quality filtering, 1,543,733 and 790,479 high-quality 16S rRNA genes and ITS sequences were obtained, with a minimum of 47,504 and 29,277 sequences per sample, respectively. After clustering at the 97% level, 16S rRNA genes and ITS sequences were grouped into 5,833 and 1,490 OTUs and used for bacterial and fungal community analyses, respectively. Metagenomic data processing including the procedures for quality control, contig assembly, open reading frames prediction, and taxonomic classification and functional annotation of genes can be found in Text S1 for reference in the Supporting information.

Bacterial functional profiles were predicted using the PICRUSt2 software (<https://github.com/picrust/picrust2/wiki>) and KEGG orthology (de Scally et al., 2016; Douglas et al., 2020). A previously reported dataset was selected to evaluate soil microbial

genes involved in P transformation (Bergkemper et al., 2016; Dai et al., 2020; Mcgrath et al., 2013).

2.5. Real-time fluorescence quantitative PCR (qPCR)

qPCR using primers listed in Table S1 was performed to determine the absolute abundance of predominant genes responsible for P starvation response (*phoU* and *phoR*), inorganic P solubilization (*gcd* and *ppx*) (Liang et al., 2020; Zhang et al., 2023), and organic P mineralization (*phoN*, *phoD*, *phnX*, and *appA*) (Long et al., 2018; Lu et al., 2022). The reaction mixtures included 1.0 µL DNA, 10 µL 2 × Taq PCR Master Mix (Allwegene, Beijing, China), and 0.5 µL (10 µM) of the forward and reverse primers. Sterile water was added to the reactions until the mixture volume was fixed to 18 µL. The amplification conditions were as follows: 94 °C for 5 min, 30 cycles at 94 °C for 30 s, 55 °C for 30 s, 72 °C for 30 s. Standard curves were plotted using a 10 × plasmid DNA dilution series containing the target gene.

2.6. Statistical analysis

A repeated measures ANOVA was performed conducted to assess the impact of N addition level, duration, and their interaction on various variables, including soil available P, organic P, microbial characteristics (biomass and phosphatase), and microbial diversity and richness indices. The significance level was set at $p < 0.05$. The normality of residuals and the homogeneity of variance were assessed using the Shapiro–Wilk and Levene's tests, respectively. Furthermore, significant differences between groups ($p < 0.05$) were determined using Duncan's test. Pearson's correlations were used to examine linear relationships between

variables. All statistical analyses were conducted using IBM SPSS Statistics for Windows (version 21.0; IBM Corp., Armonk, NY, USA).

OTUs with relative abundances exceeding 0.01% in bacterial and fungal sequences were selected for constructing a co-occurrence network (Feng et al., 2021). Initial network relationships were inferred according to the Spearman correlation matrix calculated using the “WGCNA” package in R, version 4.1.2 (<http://cran.r-project.org/>). In the network, each node represents an OTU, and the edges connecting nodes represent the correlations between OTUs. To control for false-positive results, all *p*-values were adjusted for multiple correlations using the Benjamini and Hochberg false discovery rate (FDR) (Benjamini et al., 2006). Strong correlations with Spearman correlation coefficients > 0.60 and FDR-adjusted *p*-values < 0.01 were selected to construct the co-occurrence networks. The network was visualized using Gephi, version 0.9.2 (<http://gephi.github.io/>). Subsequently, we identified the major microbial modules of strongly correlated OTUs in the co-occurrence network (modularity value > 0.4), assuming the network had a modular structure (Shi et al., 2016). The relative abundance of each module was calculated by averaging the standard relative abundances (z-scores) of all taxa within the module (Delgado-Baquerizo et al., 2018).

We calculated Spearman correlations between soil enzyme activities and the major bacterial and fungal modules. Additionally, we used the multiple regression model (lm function in the “stats” package in R) (Jiao et al., 2020) with variance decomposition analysis (calc.relimp function in the “relaimpo” package in R) (Grömping et al., 2006) to estimate the importance of the differences in major modules in explaining the dissimilarities in the absolute abundance of P cycling genes. Afterward, to identify the statistically significant predictors in

the absolute abundance of P cycling genes, a random forest analysis was used by incorporating all species in keystone modules into the model using the “rfPermute” package (Feng et al., 2021).

Based on known and potential relationships, a structural equation model (SEM) was used to evaluate the direct or indirect effects of microbial P-cycling genes on soil organic P and available P under N addition (Grace, 2006). The SEM was fitted by maximum likelihood estimation using IBM SPSS Amos 21.0 software (Amos Development Corporation, Chicago, IL, USA). The models were considered to have a good fit when the p -value > 0.05 , standardized root mean square residual (SRMR) < 0.10 , and root mean square error of approximation (RMSEA) < 0.08 (Ma et al., 2021). Box plots were generated using the "ggplot2" package in R.

3. Results

3.1. Foliar and root P concentration

Five years of N addition significantly increased the aboveground biomass of Moso bamboo (Table S2) and P concentration in the leaves, especially under the HN treatment (Fig. S1; $p < 0.05$). This was also the case for the root P concentration, which significantly increased under N addition treatments during the experimental period ($p < 0.05$).

3.2. Soil properties and P status

N addition did not affect soil pH, SOC, and TN concentrations during the experimental period (Table S3), whereas significantly increased the N availability ($p < 0.05$). N addition also caused

a significant decline in available and organic P concentrations in the soil ($p < 0.05$; Figs. 1 and S2), whereas no significant changes were observed in inorganic and total P ($p > 0.05$). Available P and organic P decreased by 32.54% and 27.88%, respectively, in the LN treatment and by 5.14% and 24.63%, respectively, in the HN treatment compared with those in the control group.

3.3. Microbial biomass C, N, and P

N addition induced a significant reduction in MBP (Table 1; $p < 0.01$); however, it did not affect MBC or MBN ($p > 0.05$). Similarly, N addition did not alter the MBC:MBN ratio but increased the MBC:MBP and MBN:MBP ratios, especially in the 4th and 5th year of N addition (Table 1).

3.4. Phosphatases activities

The activities of ACP ($p < 0.01$) and ALP ($p < 0.05$) were significantly increased by N addition, but not those of PD in the soil ($p > 0.05$; Figs. 2 and S3). The duration of N addition (T) did not affect ACP activity ($p > 0.05$) but significantly increased PD and ALP activities ($p < 0.05$). No significant interactions ($T \times N$) were observed between the N addition duration and level on the activities of the three phosphatases ($p > 0.05$).

3.5. P-related microbial functional genes

The results of qPCR revealed that N addition significantly increased the absolute abundance of genes involved in P starvation response regulation (*phoU* and *phoR*) and organic P mineralization (*phoN*, *phoD*, and *appA*) ($p < 0.05$; Figs. 3 and S4). N addition also significantly increased the absolute abundance of *pstS* ($p < 0.05$), but did not changed *pitA*. N addition had

minor or negative effects on the absolute abundance of genes involved in inorganic P solubilization (*ppx* and *gcd*).

To obtain more information, we also detected 38 genes regulating P transformation based on PICRUSt2 software; these genes and their corresponding KO numbers are shown in Table S4 and Fig. S5. Similarly, the results of PICRUSt2 showed that N addition induced significant higher relative abundance of P starvation response regulation and organic P mineralization genes, although overall, it had minor effects on P-uptake and P-transport genes and inorganic P solubilization genes (Fig. S6).

3.6. Microbial P cycling ecological module

To further investigate the microbial functional potential associated with soil P cycling, soil microbial communities were grouped into different modules based on co-occurrence networks. Five microbial modules were created for both bacterial (B_Mod#0–5) and fungal (F_Mod#0–5) communities (Fig. 4a). The bacterial modules included eight phyla (*Firmicutes*, *Proteobacteria*, *Actinobacteria*, *Chloroflexi*, *Acidobacteria*, *Planctomycetes*, *Gemmatimonadetes*, and *Bacteroides*). In contrast, the fungal modules contained five specific (*Ascomycota*, *Basidiomycota*, *Rozellomycota*, *Mortierellomycota*, and *Mucoromycota*) and one unassigned phylum (Fig. S7).

N addition significantly affected the bacterial modules (B_Mod#0–5) and particularly increased the abundance of B_Mod#0 ($p < 0.05$; Fig. S8). In contrast, no significant changes were observed in the abundance of fungal modules (F_Mod#0–5) in response to N addition ($p > 0.05$; Fig. S8). B_Mod#0 was the microbial module that contributed the most to changes in P

cycle functional genes (> 60%; Fig. 4b), with a significant positive correlation with P starvation response regulation (e.g., *phoU* and *phoR*) and mineralization (e.g., *phoN*, *phoD*, *phnX*, and *appA*) genes ($p < 0.05$; Figs. 4b and S9). *Proteobacteria*, *Firmicutes*, and *Actinobacteria* were the dominant phyla in B_Mod#0, and their relative abundances increased with N addition (Table S5). Moreover, random forest analysis revealed that these taxa were the most important taxonomic phyla responsible for the increase in P mineralization genes abundance (Fig. 4c; Table S5). These findings suggest that the B_Mod#0 act as the keystone modules controlling potential P cycling.

3.7. Relationships between soil P status, microbial functional traits and enzyme activities

SEM was used to explore the direct and indirect pathways by which microorganisms regulated soil P availability under N addition (Fig. 5). Results showed that the MBN:MBP ratio was positively correlated with the absolute abundance of P starvation response regulation genes. The increased expression of P starvation response regulation genes mostly upregulated the expression of P mineralization genes but not P solubilization genes, consequently enhancing the activities of phosphatase, which indirectly accelerated organic P mineralization (Fig. 5).

4. Discussion

4.1. Soil P status and microbial biomass P in response to N addition

In this study, a significant decrease in soil P availability induced by N addition was observed (Fig. 1), which is in agreement with studies reporting the effect of N addition on subtropical *Pinus elliottii* plantations and *Castanopsis* forests (Fan et al., 2020; Kou et al., 2018). These

findings support the hypothesis that N deposition exacerbates P limitation in subtropical forest soils with low P availability (Fan et al., 2019a; Li et al., 2016; Vitousek et al., 2010). N addition usually leads to the leaching of exchangeable base cations (e.g., Ca^{2+} and Mg^{2+} ; Li et al., 2018; Pereg et al., 2018), which may activate Al and hydrous oxides and thereby enhancing phosphate fixation (Fan et al., 2019a; Wang et al., 2022b; Zhou et al., 2018). Another reason is the rapid uptake of P caused by N deposition-stimulated plant growth (Fan et al., 2018; Kimmel et al., 2020; Yang et al., 2019), which was also evident by the fact that the aboveground biomass and P concentrations in Moso bamboo increased after N addition (Table S2 and Fig. S1).

However, organic P concentration declined significantly after N addition (Fig. 1). Similarly, previous studies have reported that organic P decreased because of experimental N enrichment in a subtropical natural forest and deciduous forests in the Bear Brook (Fan et al., 2018; Heuck et al., 2018). Considering that organic P serves as an important source of available P after decomposition (Fan et al., 2021b), its decrease implies that continuous N deposition may deplete potential P sources in subtropical forest soils stored in the form of organic P (Yang et al., 2021). Contrastingly, no significant changes in inorganic P were observed following N addition. One possible reason is that less than 2% of the inorganic P belongs to bioavailable P and can be utilized, while majority of the inorganic P is fixed by Fe and Al oxides and minerals (Carreira et al., 2000; Fan et al., 2019a). Therefore, inorganic P may be inaccessible to microbes and plant roots and insensitive to environmental changes, thereby contributing to the unchanged inorganic P observed. Furthermore, the decomposition of organic P releases inorganic P, which is subsequently absorbed by soil minerals, compensating the loss of inorganic P and maintaining the equilibrium of inorganic P in the soil.

Microbes play a critical role in regulating nutrient cycling in the soil, and soil microbial biomass is sensitive to environmental changes and inputs such as N deposition (Carrara et al., 2021; Li et al., 2014; Moore et al., 2021; Wang et al., 2018b; Zhang et al., 2018). We found that N addition significantly decreased MBP but increased MBC:MBP and MBN:MBP (Table 1). Besides, the results of enzymatic stoichiometry revealed that the vector angle was $> 45^\circ$ and increased with N addition (Fig. S10; unpublished data). These results indicate that additional N input could increase microbial P demand and may even intensify the P limitation for microbes at the study site. This can be further supported by an increase in the ACP and ALP activities (Fig. 2) and an experiment with P addition in vitro (Zeng et al., 2024). Nevertheless, no significant changes in MBC and MBN were observed, despite that N addition directly increased soil N availability (Table S3). It is possible that microbial growth was more limited by P, leading to a greater investment of C and N in acquiring P (e.g., producing phosphatase) rather than storing them in microbial biomass. An alternative or complementary explanation is that N addition did not reduce the pH (Table S3), owing to Al^{3+} serving as the dominant cation in buffering soil acidification when the pH approaches 4.5 (Bowman et al., 2008; Tian and Niu, 2015). The stronger buffering capacity of Al^{3+} compared to that by exchangeable based cations maintains stable pH levels, thereby protecting microbes from the harmful effects induced by a low pH (Vitousek and Farrington, 1997; Zhang et al., 2018).

4.2. Soil microbial P function potential in response to N addition

Consistent with the increase in phosphatase activity, N addition significantly enhanced the absolute abundance of P cycling genes, including those involved in P starvation response

regulation (e.g., *phoU* and *phoR*) and organic P mineralization (e.g., *phoN*, *phoD*, and *appA*) (Figs. 3, S4, and S5). Previous studies have also found that soil microorganisms upregulate the expression of functional genes harboring phosphatases (e.g., *appA*, *phoN*, and *phoD*) under soil P deficiency conditions (Jiang et al., 2021; Liang et al., 2020; Ragot et al., 2017; Wei et al., 2019). These findings suggest that N deposition stimulates the microbial P starvation response and enhances organic P hydrolysis by microbes (Fan et al., 2018; Heuck et al., 2018). Contrastingly, Dai et al. (2020) reported that N fertilizer increased the abundance of *phoR* but reduced *phoD* in agroecosystem soil, attributed to a decline in pH. This decline suggests an inhibitory effect of soil acidification on microbial-driven organic P mineralization. Therefore, the sustained pH in this study is conducive to microbial functional performance, which offers an explanation for the increased abundance of organic P mineralization genes. Furthermore, N addition significantly increased the abundance of *pstS*, which encodes high-affinity phosphate-specific transporter, but had minor impact on *pitA*, a gene that encodes a transporter with low-affinity. This suggests that in low P soil, microbes prioritize the expression of high-affinity *pstS* over *pitA* to facilitate phosphate transportation (Hsieh and Wanner, 2010).

In contrast to the increased abundance of P mineralization genes, the abundance of inorganic P solubilization genes (*gcd* and *ppx*) did not increase after N addition (Fig. 3). This indicates that microbes at the study site utilize P from organic P mineralization rather than inorganic P solubilization again, which is contrary to the earlier studies reporting that inorganic P desorption serves as an important resource of bioavailable P (Ghodszad et al., 2022; Liang et al., 2020). These conflicting results may partly be attributed to the availability of inorganic or organic P and their ease of accessibility by soil microbes. In the experimental plots, due to its

strong adsorption to Fe and Al oxides, most inorganic P was hard to be accessed and utilized by microbes and thus non-bioavailable (Vitousek and Farrington, 1997; Walker and Syers, 1976). Contrastingly, organic P accounts for approximately 40% of soil total P (Fan et al., 2019a; Fan et al., 2021b) and is more easily desorbed from soil oxides and enzymatically hydrolyzed to release available P for microorganisms compared to that by inorganic P (Fan et al., 2018; Landeweert et al., 2001; Richardson and Simpson, 2011). Furthermore, earlier studies focusing on the influence of N deposition on soil P fractions emphasized that organic P is critical in supplying available P to plants and microbes in subtropical areas (Fan et al., 2018; Fan et al., 2019a). This is further supported by the positive correlation between organic P and MBP observed in this study (Fig. S11). Furthermore, we observed that ACP had higher explanation ($R^2 = 0.54$) than that by ALP ($R^2 = 0.37$) for the decrease in organic P (Fig. S12). This result supports the fact that ACP plays a more important role than that by ALP in organic P mineralization in acidic soils, as ACP demonstrates maximal activities in acidic conditions (Tazisong et al., 2015; Zheng et al., 2021). This is further corroborated by the increased expression of the genes *phoN*, which encode ACP, following N addition. Overall, we speculated that microorganisms respond to P deficiency primarily by increasing the abundance of P mineralization genes and the associated organic P decomposition.

4.3. Keystone microbial ecological modules associated with P transformation under N addition

To cope with nutrient deficiency (e.g., N and P), soil microbes may reestablish their potential associations and form modules to ameliorate nutrient limitation (Fan et al., 2019b; Fan et al., 2021a; Han et al., 2021). This was confirmed in the present study, revealing that B_Mod# was

the keystone module regulating P cycling functional potential. Furthermore, the abundance of some *Proteobacteria*, *Actinobacteria*, and *Firmicutes* in B_Mod#0 increased with N addition and were the key phyla contributing to the increase in organic P mineralization genes abundance (Fig. 4c; Table S5). *Proteobacteria*, *Actinobacteria*, and *Firmicutes* were demonstrated to contain phosphatases encoding genes such as *phoN* (coding for ACP) and *phoD* (coding for ALP) (Alori et al., 2017; Franco-Correa et al., 2010; Ragot et al., 2017; Wei et al., 2016), and their abundance is positively correlated with phosphatase enzyme activity (Cui et al., 2018). These findings indicate that N addition increased the abundances of *Proteobacteria*, *Actinobacteria*, and *Firmicutes* phyla, and strengthened their association to form B_Mod#0, ultimately contributing to activate soil P.

To clarify the molecular mechanism of microbial-driven P mineralization, we further identified the keystone taxa in the present study. Our analysis revealed that the *Alphaproteobacteria* of *Proteobacteria* phylum covered most *phoD* and *phoN*-harboring bacteria, and were significantly increased after N addition (Fig. S13). Such findings implied that the N input increased *Alphaproteobacteria*, enhancing expression of P-mineralization genes, which subsequently regulated the production of P-cycling enzymes (e.g., ACP and ALP) and thus contributed to soil P transformation in the experimental plot.

4.4. Pathway of soil microorganisms involved in enzymatic hydrolysis of organic P under N application

In this study, we identified the key modules and phyla of soil microorganisms involved in organic P decomposition (Fig. 4). However, the mechanism by which N deposition triggers

these processes remains unclear. SEM analysis revealed that N addition led to an imbalance in soil nutrient stoichiometry and intense P deficiency, as evidenced by the increased MBN:MBP ratio and decrease in soil P availability. In response to environmental P shortages, soil microbes effectively assimilate alternative forms of P from the soil by activating the *phoR/phoB* two-component signal transduction system (Eder et al., 1996; Hsieh and Wanner, 2010; Lubin et al., 2016). For example, under P limitation, *E. coli* upregulates the *phoR* sensor, which phosphorylates the response regulator *phoB* and activates the phosphate regulon. However, under excessive P conditions, *phoR*, *pstS*, and *phoU* dephosphorylate *phoB* and deactivate the phosphate regulons (Hsieh and Wanner, 2010; Wanner, 1993). Similarly, our study found that N addition enhanced the absolute abundance of P starvation response regulation genes, suggesting that the two-component signaling system might regulate the expression of P starvation response genes in microbes under P limitation conditions in the forest soil.

Furthermore, we found that an increase in the absolute abundance of P starvation response genes (path coefficient = 0.91) had a positive effect on organic P mineralization genes (e.g., *phoN* and *phoD*) (Fig. 5). These results are consistent with the findings of a previous study, in which the expression of the P mineralization gene *phoD* was dependent on the sensor gene *phoR* (Eder et al., 1996), suggesting that the enhanced expression of organic P mineralization genes might be a result of the activation of P starvation response genes because of N application. Microorganisms secrete extracellular enzymes, such as ACP and ALP, as products of the corresponding encoding genes, *phoN* and *phoD* genes, respectively (Eder et al., 1996; Liang et al., 2020; Luo et al., 2017; Ragot et al., 2017). Therefore, the increased abundance of the organic P-mineralization genes under N application might enhance microbial phosphatase

production (Eder et al., 1996). These results could support the causal relationship between organic P mineralization genes (*phoN* and *phoD*) and phosphatase activities (ACP and ALP) (Fig. 5).

Furthermore, the SEM results showed that phosphatase activity was negatively correlated with organic P (Fig. 5), consistent with the results of previous studies that showed that higher phosphatase activity facilitates organic P hydrolysis (Fujita et al., 2017; Sharma et al., 2020). Thus, the decomposition of organic P resulted in the release of available P. Collectively, our findings suggest that microbial-driven P transformation plays a critical role in meeting increased P demand of Moso bamboo, a rapid-growing species with well-developed underground root system. This, in turn, induces an intense microbial P "starvation response", thereby driving soil P-cycling (Fig. 6). However, the suitability of such mechanisms for driving organic P mineralization in other forest soils still requires further field-based evidence, especially in tropical and subtropical forests with high species diversity.

5. Conclusion

In this study, we observed that N addition exacerbated microbial P limitation, as indicated by imbalanced nutrient stoichiometry (MBC:MBP and MBN:MBP) and significantly decreased soil available and organic P. To cope with P limitations, microorganisms formed microbial modules, with *Proteobacteria*, *Actinobacteria*, and *Firmicutes* being the predominant bacterial phyla in the functional module (B_Mod#0). These bacterial phyla acted as keystone phylotypes, regulating the potential for P cycling. Bacteria harboring P starvation response genes (*phoU* and *phoR*) responded to P deficiency by increasing the absolute abundance of organic P

mineralization genes (*phoN* and *phoD*), thereby enhancing the hydrolysis of organic P. Our findings provide insights into the molecular mechanisms underlying microbe-driven P cycling in subtropical forest soils under N deposition, specifically in the enzymatic hydrolysis of organic P.

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Figures legends

Fig. 1. Variations in soil available, organic, and total P concentrations after N addition. Box plots show the median (black line), 25 and 75% (box edge), and 10 and 90% of the data set (error bars). T, time effect; N, N addition effect; T × N, interactive effect of time and N addition. ** indicate a significant difference at $p < 0.01$ level. n.s., no significance ($p > 0.05$). Different lowercase letters indicate significant differences between treatments ($n = 9$).

Fig. 2. Variations in soil phosphatase activities after N addition. ACP, acid phosphatase; PD, phosphodiesterase; ALP, alkaline phosphatase. Box plots show the median (black line), 25 and 75% (box edge), and 10 and 90% of the data set (error bars). T, time effect; N, N addition effect; T × N, interactive effect of time and N addition. Asterisks (*) indicate a significant difference (* $p < 0.05$, ** $p < 0.01$). n.s., no significance ($p > 0.05$). Different lowercase letters indicate significant differences between treatments ($n = 9$).

Fig. 3. Abundance of genes responsible for P starvation response regulation (*phoU* and *phoR*), organic P mineralization (*phoN*, *phoD*, *phnX*, *appA*), inorganic P solubilization (*ppx* and *gcd*), and P-uptake and P-transport (*pitA* and *pstS*). Box plots show the median (black line), 25 and 75% (box edge), and 10 and 90% of the data set (error bars). T, time effect; N, N addition effect; T × N, interactive effect of time and N addition. Asterisks (*) indicate a significant difference (* $p < 0.05$, ** $p < 0.01$, and *** $p < 0.001$). n.s., no significance ($p > 0.05$). Different lowercase letters indicate significant differences between treatments ($n = 9$).

Fig. 4. (a) Major bacterial and fungal modules, (b) correlations between modules and P cycling genes, and (c) predominant bacterial phyla (top 10) that influence P cycling genes. (a) Representation of microbial modules induced by N addition. The network diagram nodes represent OTUs colored by each of the major modules within co-occurrence network of bacterial and fungal communities, respectively; (b) identifying the key modules that predominantly contributed to the changes of the abundance of P cycling genes based on correlation and multiple regression models. The size of the circle represents the importance of the variable (i.e., proportion of explained variability calculated via multiple regression modeling and variance decomposition analysis). Colors represent Spearman correlations; (c) the keystone species of key bacterial module (B_Mod#0), responsible for the changes of the abundance of P cycling genes based on random forest analysis. Grey bars indicate species that had no significant effect on the abundance of P cycle genes ($p > 0.05$). Asterisks (*) indicate a significant difference (* $p < 0.05$, ** $p < 0.01$, and *** $p < 0.001$).

Fig. 5. Pathway by which soil microorganisms involved in enzymatic hydrolysis of organic P combat soil P deficiency under N addition, revealed by structural equation models (SEM). Red, blue, and grey solid arrows indicate significant positive, negative, and non-significant effects, respectively. * $p < 0.05$, ** $p < 0.01$, and *** $p < 0.001$. MBN:MBP, the ratio of microbial biomass N to microbial biomass P; Phosphatase, represents the summary activities of acid and alkaline phosphatases. GFI, goodness-of-fit index; RMSEA, root mean square error of approximation; SRMR, standardized root mean square residual.

830

831 **Fig. 6.** Conceptual framework of enzymatic hydrolysis of organic P co-regulated by keystone
832 bacterial functional module and genes responsible for P mineralization induced by N addition
833 in a subtropical bamboo forest ecosystem.

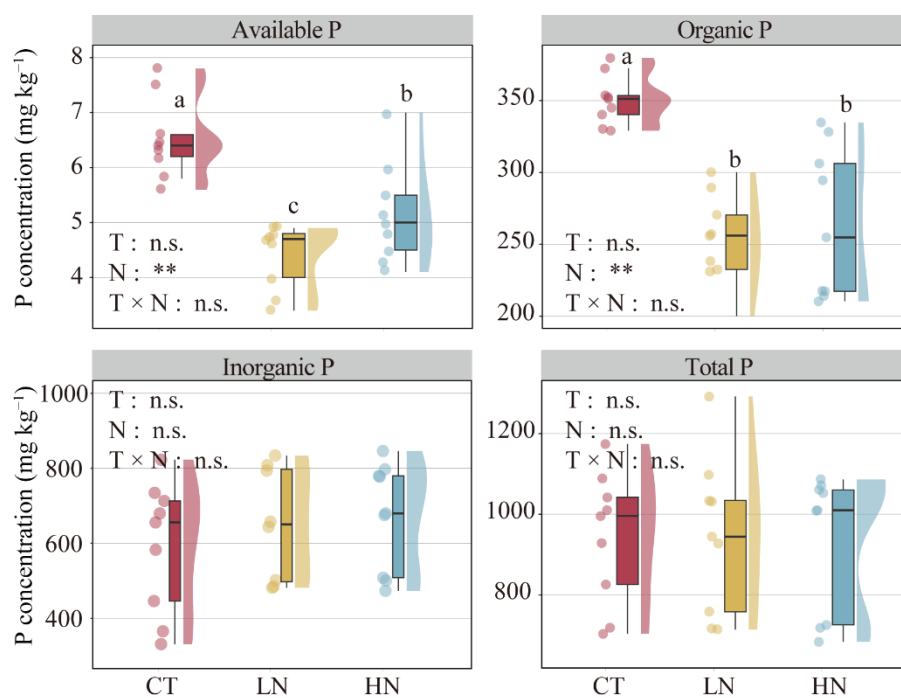


Fig. 1

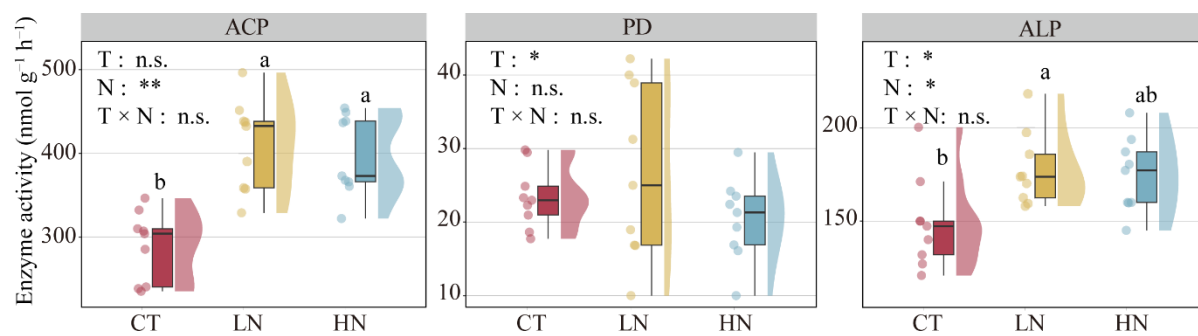


Fig. 2

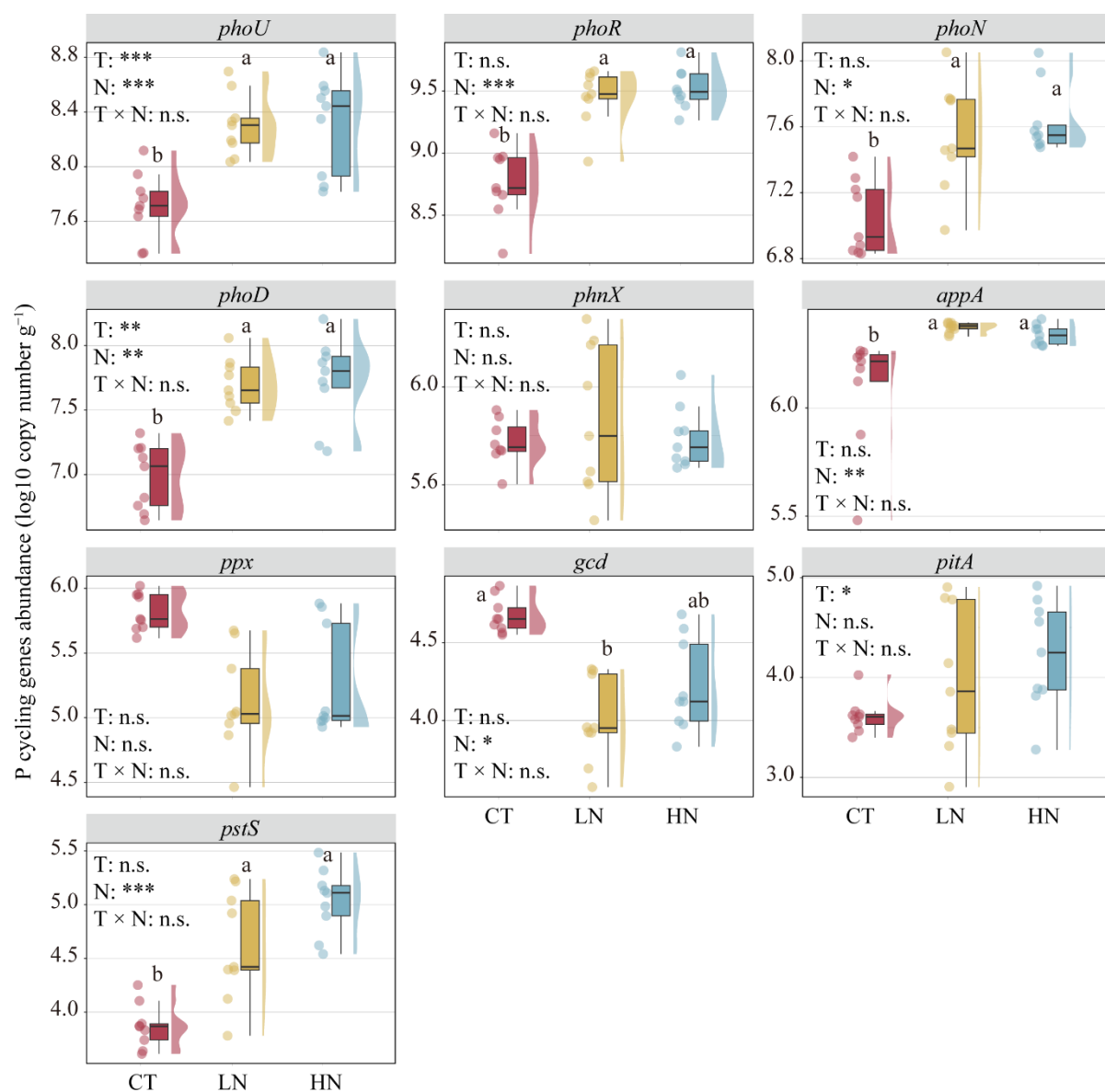


Fig. 3

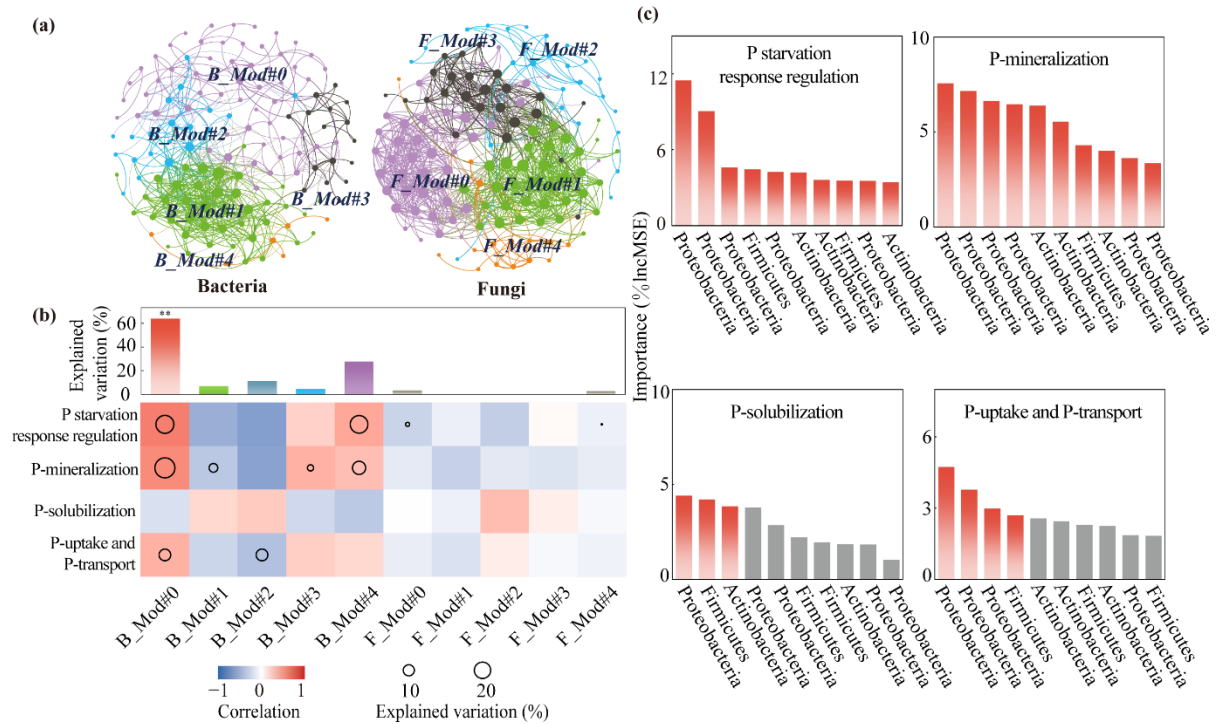
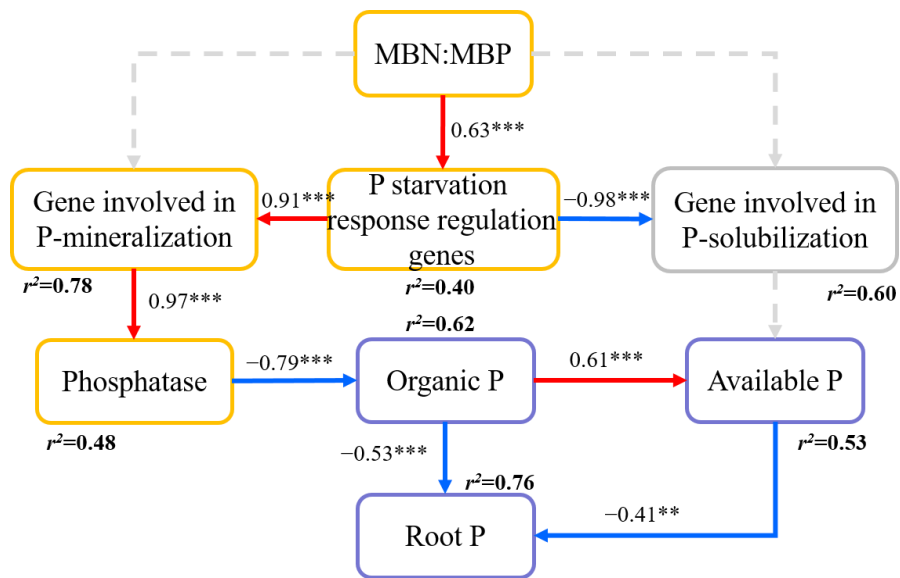


Fig. 4



GFI = 1.000, $p = 0.591$, RMSEA = 0.061, SRMR < 0.001

Fig. 5

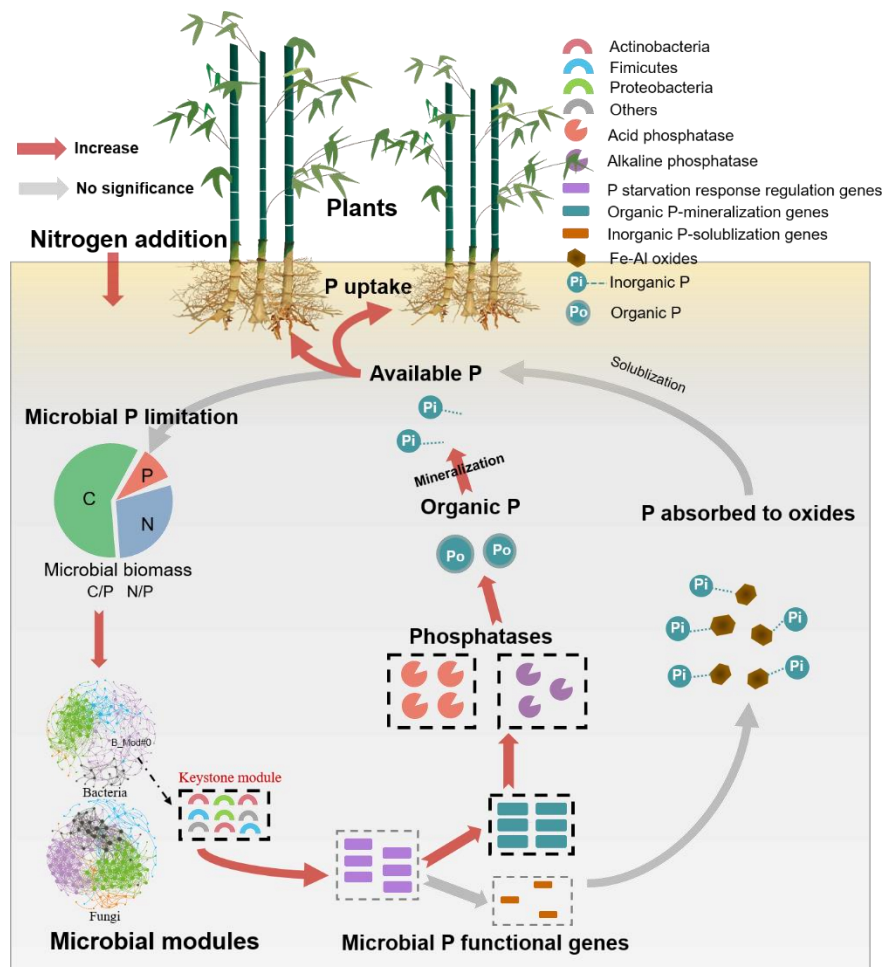


Fig. 6

Table 1. Variations in microbial biomass and its stoichiometry with N addition.

Year	Treatment	MBC (mg kg ⁻¹)	MBN (mg kg ⁻¹)	MBP (mg kg ⁻¹)	MBC:MBN	MBC:MBP	MBN:MBP
3 rd -year	CT	780 (156)	96 (7.8)	43 (6.5) a	8.2 (1.5)	18.5 (3.7) a	2.3 (0.2) a
	LN	1078 (283)	79 (8.2)	45 (8.5) a	14.1 (4.2)	25.5 (7.3) a	1.8 (0.2) a
	HN	1509 (54)	114 (16.7)	47 (3.6) a	13.1 (1.8)	32.5 (1.5) a	2.4 (0.2) a
4 th -year	CT	1119 (205)	104 (9.8)	69 (4.9)a	10.5 (0.7)	16.2 (3.0) b	1.5 (0.1) b
	LN	966 (128)	91 (5.8)	41 (4.1)b	10.8 (1.1)	24.6 (5.1) ab	2.2 (0.2) ab
	HN	1161 (183)	106 (10.9)	38 (4.5)b	11.3 (1.2)	30.5 (4.4) a	2.8 (0.2) a
5 th -year	CT	1255(128)	139 (2.5)	50 (3.3) a	8.9 (0.4)	24.8 (1.3)b	2.8 (0.1) c
	LN	1140 (147)	124 (21.3)	35 (2.1) b	10.0 (2.5)	32.6 (4.4)ab	4.2 (0.3) b
	HN	1277 (74)	116 (16.2)	31 (3.6) b	11.6 (1.4)	42.1 (5.3)a	3.7 (0.4) a
Repeated	T	n.s.	*	n.s.	n.s.	*	***
measures	N	n.s.	n.s.	**	n.s.	***	**
ANOVA	T×N	n.s.	n.s.	*	n.s.	n.s.	**

Note: All values are means $\pm$ SE ($n = 3$). MBC, microbial biomass C; MBN, microbial biomass N; MBP, microbial biomass P; CT, control; LN, low N addition; HN, high N addition. T, time effect; N, N addition effect; T $\times$ N, interactive effect of time and N addition. Asterisks (*) indicate a significant difference (* $P < 0.05$, ** $P < 0.01$, and *** $P < 0.001$). n.s., no significance ($P > 0.05$). When only main effect of N addition or its interactive effect with time was significant at $p < 0.05$, Duncan's test was used to evaluate the significant differences (indicated by different lowercase letters) among N addition treatments in the same year.

Highlights

- N addition reduced soil P availability and intensified microbial P limitation
- N addition activated microbial P starvation response and enhanced expression of P mineralization genes
- B_Mod#0 acts as a keystone phylotype in enhancing functional potential for organic P mineralization
- Organic P mineralization is more important than inorganic P desorption for microbial amelioration of P limitation.

Declaration of Competing Interest

The authors declare that they have no known competing financial interests or personal relationships that could have appeared to influence the work reported in this paper.