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26 Late spring frost delays tree spring phenology by reducing photosynthetic

27 **productivity**

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Abstract

Under climate warming, earlier spring phenology has heightened the risk of late spring frost (LSF) damage. However, the intricate interplay among LSF, spring phenology, and photosynthetic carbon uptake remains poorly understood. Utilizing 286,000 ground phenological records involving 870 tree species and remote-sensing data across the Northern Hemisphere, we show that LSF occurrence in a given year reduces photosynthetic productivity by 13.6%, resulting in a delay in spring phenology by approximately 7.0 days in the subsequent year. Our experimental evidence, along with simulations using modified process-based phenology models, further supports this finding. This frost-induced delay in spring phenology subsequently leads to a decrease in photosynthetic productivity during the next year following an LSF event. Therefore, it is essential to integrate this frost-induced delay in spring phenology into current Earth System Models to ensure accurate predictions of the impacts of climate extremes on terrestrial carbon cycling under future climate change.

Main

Tree phenological events, encompassing seasonal biological activities like budbreak and leafout, wield substantial influence over carbon, water, nutrient cycling, and trophic interactions within forest ecosystems^{1,2}. Late spring frost (LSF), characterized by frost events taking place during or after leaf emergence, emerges as a significant threat to tree growth and the functionality of forest ecosystems³⁻⁵. Trees undergo deep physiological dormancy, termed endodormancy, during winter, which maximizes frost resistance until adequate chilling units accumulate. As spring arrives, trees transition to ecodormancy, accompanied by diminishing freezing resistance and the emergence of new leaves, marking the lowest freezing resistance⁶. The occurrence of LSF after leaf-out can inflict severe harm upon these nascent leaves^{7,8}. Hence, trees strive to initiate leaf growth as early as possible, optimizing the production of photosynthetic carbohydrates while sidestepping the perils of LSF^{9,10}. The past decades have witnessed the propagation of earlier spring phenology due to global warming across the Northern Hemisphere^{11–13}. This warming-induced advancement in spring phenology contributes to extended carbon assimilation periods but also heightens the risk of LSF^{14–16}. Consequently, it becomes imperative to comprehend and model the intricate dynamics between LSF, tree productivity, and climate change to establish effective forest management strategies.

Sufficient carbon assimilation and storage during the prior growing season are vital for winter survival and supporting early spring leaf growth^{17,18}. Nonstructural carbohydrates (NSC), including soluble sugars and starch, peak in autumn before dormancy but dwindle after new leaves emerge in spring^{19,20}. Frost events amplify stem respiration, intensify carbon depletion,

and trigger the formation of ice crystals, potentially damaging biomembranes and dehydrating cells, ultimately diminishing stored carbon reserves^{7,21}. A significant drop in carbon assimilation renders available carbon reserves inadequate for fueling the energy and osmotic needs of leaf-out^{22,23}. Consequently, more chilling or forcing units might be needed to initiate spring leaf-out, leading to delayed spring phenology^{22,23}. LSF may disturb the uptake of photosynthetic carbon and impact the accumulation of chilling and forcing units as supplementary energy sources, thus influencing spring phenology in the subsequent year. However, the interactions between frost, carbon uptake, and phenology spanning multiple growing seasons in trees have remained largely unexplored.

In this study, we endeavor to unravel the intricate interplay between LSF, photosynthetic productivity, and spring phenology across the Northern Hemisphere by combining multiple complementary long-term and large-scale datasets and frost-controlled experiments. Specifically, we embark on deciphering the carry-over effect of LSF on subsequent-year spring phenology and photosynthetic productivity across subtropical, temperate, and boreal regions in the Northern Hemisphere. Our hypothesis postulates that the reduction in carbon uptake due to frost triggers an augmented demand for (1) chilling units, (2) forcing units, or (3) both, thereby delaying the onset of spring phenology in the subsequent year. This frost-induced delay, in turn, would decrease productivity through a feedback process.

Carry-over effect of LSF on spring phenology

To investigate the carry-over effect of LSF on spring phenology, we compiled leaf-out records of 870 tree species spanning from 1950 to 2020. These records were sourced from five largescale ground-based phenological and PhenoCam networks situated in Europe, the USA, China, and Russia, spanning diverse climate zones within the Northern Hemisphere (Fig. 1a). The leafout dates were employed to denote the start of season (SOS) and LSF events were defined as occurrences of daily minimum temperatures dropping below -2°C after the SOS. We test the impact of LSF on SOS, and computed the difference in SOS (ΔSOS) for consecutive two-year periods, per species at each site. Subsequently, we contrasted the differences in ΔSOS when LSF transpired versus when it did not, for each species at each site. Positive and negative values of ΔSOS indicated delayed and advanced SOS in the subsequent year (SOS_{next}) relative to the present year (SOS_{current}). Through the utilization of ground-based phenological networks, we found that, in the absence of LSF, the mean Δ SOS remained close to zero, while it was positive $(5.70 \pm 0.06 \text{ days})$ during LSF occurrences (Fig. 1b). This observation indicated a significant delay in SOS_{next} attributable to LSF in the current year. Similar results were obtained across diverse ground-based phenological networks and through analysis using linear mixed models (Supplementary Fig. 1 and Supplementary Table 1).

Generalizing our findings, we calculated ΔSOS based on phenological metrics obtained from GIMMS NDVI3g and MODIS datasets spanning 1982 to 2019 (Fig. 1c–h). The analysis employing GIMMS NDVI3g highlighted that 85.7% of forested areas depicted earlier SOS_{next} , indicated by negative ΔSOS values, when LSF did not occur in the Northern Hemisphere (Fig. 1c). Conversely, in the presence of LSF, 90.6% of forested areas displayed delayed SOS_{next} (Fig. 1d). These trends were consistent when utilizing the MODIS data (Fig. 1f, g). On the whole, SOS_{next} was found to advance by 2.30 ± 0.01 days when LSF was absent, whereas it exhibited a delay of 7.18 ± 0.03 days when LSF was present (Fig. 1e, h). To ensure the robustness of our results, we further excluded the influence of preseason temperatures on spring phenology (see Methods). We also observed that SOS_{next} was delayed by LSF after excluding the effect of spring temperature (Supplementary Fig. 2 and Supplementary Table 2).

Analogous outcomes were evident across diverse tree species, forest types, and climate zones, as inferred from both ground-based phenological datasets and remote sensing data (Fig. 2a, b and Supplementary Figs. 3–5). The observed delay in SOS_{next} was significantly more pronounced in broadleaf tree species compared to conifer species and in late-leafing species compared to early-leafing species (Fig. 2a, b). Additionally, we observed a significant escalation in frost-induced delayed SOS_{next} as the severity of LSF heightened across various species (Extended Data Fig. 1). To assess the robustness of our results, we also examined the impact of LSF occurring one to eight weeks after SOS and before July 15th on SOS_{next} using a one-week smoothing window. Consistently, we observed a delay in SOS_{next}, regardless of the temperature thresholds or timing of LSF occurrence (Supplementary Fig. 6). Moreover, we found that the delay in SOS_{next} due to frost was more pronounced when LSF occurred later compared to earlier in the season. The risk of LSF was also higher for later occurrences compared to earlier ones (Supplementary Fig. 7).

Species-specific phenological responses to LSF

To glean greater clarity about disparities in the response of SOS_{next} to LSF across tree species, we calculated the lethal temperature at which 50% of leaves were impaired (LT₅₀). We selected long-term and extensive SOS records from the PEP725 network for the calculation of LT₅₀. The LT₅₀ was estimated for all eight species, based on median SOS dates for each species within the PEP725 dataset and six experimentally derived LT₅₀ values from pertinent literature sources (Supplementary Table 3). We observed a strong positive correlation between measured and predicted LT₅₀ values (Supplementary Fig. 8). The outcomes revealed an LT₅₀ range of -8° C to -2° C across the eight species, with early-leafing species exhibiting a significantly lower LT₅₀ than their late-leafing counterparts (Fig. 2c). Employing linear regression, a positive linear correlation was evident between LT₅₀ and Δ SOS for consecutive two-year periods (Fig. 2d).

This finding suggested that tree species with low frost resistance experienced a more pronounced delay in SOS_{next} compared to those with high frost resistance when LSF occurred.

Interactions between LSF, phenology and productivity

To expound upon the frost-induced alterations in SOS_{next}, we examined the discrepancies in gross primary productivity (GPP) and net primary productivity (NPP) when LSF occurred or did not. We found the GPP of the current year (GPP_{current}) decreased by 13.6% when LSF occurred (Fig. 3a, b). We obtained similar results utilizing remote-sensing vegetation indices and tree-ring width in the Northern Hemisphere (Supplementary Figs. 9 and 10). To investigate the effect of LSF on photosynthetic carbon uptake, we conducted frost experiments using seedlings of four widely distributed tree species in subtropical and temperate forests (see Methods). Our findings revealed a significant decrease in leaf photosynthetic rate, starch, soluble sugar, and NSC in tree seedlings when LSF occurred (Fig. 3c–j and Supplementary Figs. 11–14). These results provide experimental evidence that LSF significantly reduces leaf photosynthetic rate and carbon storage.

 To bolster these findings, a piecewise structural equation model (SEM) was constructed to probe the direct and indirect effects of LSF on GPP_{current} and SOS_{next}. The SEM outcomes unveiled a negative correlation between LSF and GPP_{current} (Fig. 3k, l and Supplementary Table 4), affirming the substantive reduction in GPP_{current} caused by LSF. Additionally, SOS_{next} exhibited positive correlations with LSF and SOS in the current year (SOS_{current}), while demonstrating a negative correlation with GPP_{current}. This finding indicated that a decline in GPP_{current} led to a delayed SOS_{next}. Consequently, the SEM elucidated that LSF exerted a dampening effect on GPP_{current}, thereby contributing to the postponement of SOS_{next}. Analogous outcomes were discerned within NPP_{current} (Supplementary Fig. 15 and Supplementary Table 5). In addition, we found a significant decrease in GPP_{current} when LSF occurred in both drought and non-drought years (Extended Data Fig. 2).

To unravel the mechanisms underpinning the correlation between frost-induced carbon reduction and subsequent-year spring phenology, we postulated that trees would necessitate additional chilling units, forcing units, or a combination thereof, to initiate spring phenology following decreased carbon assimilation in the preceding year. To test these hypotheses, a range of process-based spring phenological models were modified by integrating the impact of LSF-induced GPP reduction on spring phenology (see Methods). We found that the modified phenology models outperformed their respective original versions, and the modified two-phase sequential model (SM), specifically adjusted in the forcing phase, emerged as the optimal model for simulating spring phenology (Fig. 3m–p).

Projected impact of LSF on future spring phenology

Using the modified SM, we predicted shifts in spring phenology during LSF occurrences and compared these simulations with projections from temperature-driven CMIP6 models under diverse future emission scenarios (see Methods). We found that, in the absence of LSF, the mean \triangle SOS was negative (-1.70 \pm 0.02 days) (Fig. 4). However, when LSF events occurred, the mean \triangle SOS remained close to zero (0.67 \pm 0.01 days) across all four climatic scenarios simulated by CMIP6 models. Using the modified SM, we found that the mean Δ SOS became positive (5.27 \pm 0.03 days) in the presence of LSF under all four future climatic scenarios, with the highest mean \triangle SOS (7.33 \pm 0.07 days) observed under the high-emission SSP5-8.5 scenario (Fig. 4i–l). To reduce uncertainties arising from relying on a single phenological dataset, we further performed the same analysis using daily leaf area index (LAI) data from CMIP6 models under four future climatic scenarios, and obtained consistent results (Supplementary Fig. 16). Furthermore, utilizing data from all ground-based phenological networks, we observed a rising trend in the mean SOS_{next} attributed to LSF between 1950 and 2020, accompanied by an increased LSF risk (Extended Data Fig. 3 and Supplementary Fig. 17). These analyses highlight that the CMIP6 models underestimated the delaying impact of LSF on spring phenology under future emission scenarios.

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Interactions between LSF, phenology, and productivity

To decipherer the interactions between LSF, SOS, and productivity, we further examined the impact of LSF on the GPP in the next year (GPP_{next}) when LSF occurred. Results showed that GPP_{next} decreased by 11.2 % when LSF occurred (Extended Data Fig. 4a, b). Furthermore, we observed significantly higher Δ GPP_{next} values when SOS_{next} was advanced, irrespective of LSF occurrence (Extended Data Fig. 4e, f). Conversely, no substantial disparities were apparent in Δ GPP_{next} when LSF events transpired or not, regardless of the advancement or delay in SOS_{next}. Similar trends were observed within NPP_{next} (Extended Data Fig. 4c, d, g and h). This suggested that changes in GPP_{next} and NPP_{next} were closely linked to the timing of SOS_{next}.

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Discussion

In addition to the warming-induced earlier spring phenology, the risk of LSF is also increasing ^{16,24–26}. Numerous studies have demonstrated that LSF can severely damage newly developed buds and leaves of trees ^{7,8}. However, the carry-over effect of LSF on vegetation growth and productivity has been largely neglected. By combining multiple long-term and large-scale ground and remote sensing datasets, we discovered that the onset of spring phenology was significantly delayed in the year following an LSF event, resulting in reduced carbon assimilation in the subsequent year. Furthermore, our results suggested that the delayed spring phenology in the subsequent year, induced by LSF, was driven by reduced carbon uptake

in the year of occurrence. As a consequence, photosynthesis was reduced for two years following an LSF event (Fig. 5).

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Effect of LSF on spring phenology

In winter and spring, trees typically undergo periods of endodormancy and ecodormancy before initiating spring growth²⁷. Trees need to accumulate chilling units (i.e., cold temperatures) to fulfill endodormancy, while the accumulation of forcing units (i.e., warm temperatures) is necessary to break ecodormancy¹³. Warming springs can lead to earlier spring phenology by accelerating the accumulation of forcing units required for trees to end dormancy^{12,28}. However, trees also need to assimilate and store adequate carbohydrates through photosynthesis to maintain metabolic processes during dormancy, protect cells from winter freezing damage, and support bud development and early spring leaf growth in cold regions^{17,18,29}. When occurring after leaf-out, frost events can decrease photosynthetic pigments (e.g., chlorophyll), and even lead to complete and irreversible damage to photosynthesis and photosystems, resulting in decreased carbon uptake^{30–32}. Our analysis, based on multiple data sources and frost-controlled experiments, provides direct and robust evidence that LSF can significantly impact the photosynthetic carbon uptake of trees.

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If trees experience a significant decline in carbon assimilation when LSF occurs, available carbon reserves might be insufficient to fulfill the energetic and osmotic requirements to support the growth of buds and leaves^{22,23,33}. Consequently, trees may adopt a conservative strategy to reduce and avoid the risk of frost damage, resulting in longer periods of deep physiological dormancy and ecodormancy. When LSF occurs, ending the prolonged and deep dormancy period requires greater accumulations of chilling units, forcing units, or both in order to compensate for the carbon loss, ultimately leading to a delayed spring leaf-out in the following year. Therefore, LSF in a given year may delay spring phenology in the following year by altering the accumulation of chilling and forcing units required for trees to break dormancy. To provide explanations for the delayed spring phenology in the subsequent year due to frostinduced carbon reduction, therefore, we proposed three hypotheses: the reduced carbon uptake would lead to an increase in the requirements of (i) chilling units, (ii) forcing units, or (iii) both chilling and forcing units to trigger the onset of spring phenology. To test these hypotheses, we modified a series of process-based phenological models by incorporating the effect of frostinduced decrease in carbon assimilation on spring phenology. We observed that the modified process-based phenology model, with adjustments applied solely in the forcing phase, showed superior performance compared to other models. This finding suggests that, in order to supplement carbon loss, trees might require higher forcing units as an additional energy source to trigger the onset of spring leaf-out. This trade-off between carbon uptake and forcing units well explains the observed delay in spring phenology in the subsequent year when LSF occurred. By integrating various complementary data sources and conducting an additional frost-controlled experiment, our study provides compelling evidence that LSF can significantly postpone spring phenology in the following year by decreasing photosynthetic carbon assimilation in the current year.

Projected LSF effect on spring phenology

The global average temperature is projected to rise by around 2°C to 4°C within this century, increasing the risk of LSF³⁴. However, we found that CMIP6 models underestimate the effect of LSF on spring phenology under future climate warming. Although the impact of LSF on vegetation growth has recently been incorporated into Dynamic Global Vegetation Models (DGVMs)^{35–37}, the phenological response to frost events is not yet coupled in Earth System Models (ESMs). Therefore, ESMs have limitations in accurately depicting extreme events and their impacts on vegetation phenology. Due to the absence of intricate interactions between carbon uptake and spring phenology, traditional ESMs fail to detect and even underestimate the delaying effect induced by LSF³⁸. In contrast, our modified phenology model incorporates the impact of LSF-induced carbon loss, significantly enhancing the simulation of projected delays in spring phenology under future climatic scenarios. By integrating the interactions among LSF, carbon assimilation, and spring phenology, our modified phenology model offers a novel and comprehensive framework for understanding and predicting changes in spring phenology and terrestrial carbon cycling in a warming world.

Variation among tree species and regions

We observed that the magnitude of the delayed effect of LSF in the current year on spring phenology in the next year varied among different tree species and regions. Specifically, spring phenology of broadleaf tree species was delayed more than that of conifer tree species. It is reported that thick and narrow needles in conifers are often more resistant to extreme climate conditions, such as frost and drought, compared to wide and flat leaves in broadleaf trees^{39,40}. Broadleaf tree species are therefore more vulnerable to frost damage compared to conifer tree species. In addition, conifer trees carry multiple generations of needles, and old needles could compensate for photosynthesis once new needles are damaged by frost⁴¹. The clade-specific foliar traits may explain why spring phenology in conifers is less sensitive to frost than it is in broadleaved trees.

Early-leafing species often exhibit an "opportunistic" strategy to maximize photosynthetic carbon assimilation despite the risk of frost damage, compared to late-leafing species that are more sensitive to photoperiod ^{16,42,43}. We observed that spring phenology of late-leafing species

in the next year was delayed more than that of early-leafing species when LSF occurred. Generally, species that leaf emergence occurred earlier experience a higher risk of frost damage compared to late-leafing species. Therefore, early-leafing species might have evolved a higher frost resistance to reduce the damage when LSF occurred. To test this hypothesis, we further examined the relationship between frost resistance, indicated by the LT₅₀, and the timings of leaf-out in eight widely distributed temperate broadleaf tree species. Our results indicated that species with earlier leaf-out had a significantly stronger frost resistance than those with later leaf-out. Moreover, frost-induced delayed days of spring phenology showed a significant decrease with an increase in frost resistance. Therefore, species-specific frost resistance well explains why the early-leafing species are less sensitive to frost compared to the late-leafing species.

Combining long-term and large-scale ground-based phenological records across the Northern Hemisphere, we found that LSF significantly delayed spring phenology in the subsequent year by reducing photosynthetic carbon assimilation. This conclusion was further reinforced by the evidence gathered from our frost-controlled experiment and the simulations of process-based phenology models. We further observed a significant decrease in GPP in the subsequent year due to the frost-induced delay in spring phenology, indicating the feedback effect of frost-induced delay in spring phenology on GPP. Our results reveal complex interactions between spring-frost risk, phenology, and terrestrial carbon cycling. In light of the increasing risk of LSF, it is crucial to integrate our findings into Earth System Models to ensure accurate projections of the impact of climate extremes on terrestrial carbon cycling under future climate change, which can further inform decision-making regarding the ecological and economic impacts of these frost events on land management, forestry, and agriculture.

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Author contributions

- L.C. and C.W. designed the research. J.W., J.G. and H.H. performed the data analysis. J.G. and
- J.W. wrote the paper with the inputs of H.H., X. H., X. Z., Y.Y., D.W., X. G., N.G.S., S.R.,
- 348 J.P., P.C., C.W. and L.C. All authors contributed to the interpretation of the results, and
- 349 approved the final manuscript.

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Competing interests

352 The authors declare no competing interests.

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- Fig. 1 | Effect of late spring frost (LSF) on start of growing season (SOS) in the next year.
- a-h, Locations of the ground-based phenological and PhenoCam network sites (a), changes in
- 356 ΔSOS between LSF occurred and non-occurred across all phenological networks, Pan European
- 357 Phenology Network (PEP725), USA National Phenology Network (USA-NPN), China
- 358 Phenological Observation Network (CPON), Russian 'Chronicles of Nature' Network (RCNN),
- and PhenoCam network (b), spatial distribution of mean \triangle SOS from phenological products of
- 360 GIMMS NDVI3g (c, d) and MODIS (f, g) when LSF non-occurred (c, f) and occurred (d, g)
- over the Northern Hemisphere forests, and changes in ΔSOS between LSF occurred and non-
- occurred from phenological products of GIMMS NDVI3g (e) and MODIS (h). In b-h, ΔSOS
- 363 indicates difference in SOS for each consecutive two-year period at each site or pixel. In b, e
- and **h**, the Y and N indicate LSF occurred or did not in each year for each species at each site
- 365 (pixel), respectively. Statistical significance between groups was determined using two-tailed t
- 366 test. The asterisk indicates a significant difference in ΔSOS between LSF occurred and non-
- occurred (P < 0.01). The box spans from the first to the third quartile, with intermediate values
- marked as the black line in the middle of the box. In c, d, f and g, the histograms present the
- 369 frequency distributions of the percentages.

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Fig. 2 | Effect of late spring frost (LSF) on start of growing season (SOS) in the next year for broadleaf and conifer species, and for early- and late-leafing species. a,b, Changes in ΔSOS between broadleaf and conifer species (a), and between early- and late-leafing species (b) when LSF occurred or did not using ground-based phenological networks. c,d, The LT₅₀ searched and compiled from literature data for species of PEP725 dataset and predicted by the four-parameter logistic model based on the median dates of leaf-out (c), and the linear relationship between the predicted LT₅₀ and Δ SOS (**d**). In **a**, **b** and **d**, Δ SOS indicates difference in SOS for each consecutive two-year period for each species at each site. In a and b, the Y and N indicate LSF occurred or did not in each year for each species at each site, respectively. Statistical significance was determined using two-tailed t test. The asterisk indicates a significant difference in \triangle SOS between groups (P < 0.01). The box spans from the first to the third quartile, with intermediate values marked as the black line in the middle of the box. In c and d, black and red dots represent the LT₅₀ collected in the literature and predicted by the fourparameter logistic model, respectively. The black lines represent fits from a four-parameter logistic model in c and a linear regression model in d, respectively. The shading represents 95% confidence interval.

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Fig. 3 | Relationships between late spring frost (LSF), spring phenology, and photosynthetic productivity. a,b, Changes in annual gross primary productivity in the current year (GPP_{current}) between LSF occurred and non-occurred using phenological products of GIMMS NDVI3g (a) and MODIS (b). c-f, Changes in photosynthetic rate (c), starch (d), soluble sugar (e), and nonstructural carbohydrates (f) across all tree species in control (CK) and LSF group. g-j, LSF treatment using seedlings from four tree species. k,l, Piecewise structural equation model (SEM) considering both LSF and GPP_{current} using phenological products of GIMMS NDVI3g (k) and MODIS (l). m-p, The percentages of pixels with significant correlations (Freq) (m), average correlation coefficient (R) (n), Root Mean Square Error (RMSE) (o), and Kling-Gupta Efficiency (KGE) (p) between model estimates and satellitederived observations using the growing degree days (GDD), spring warming (SW), sequential model (SM), parallel model (PM), and corresponding modified models with GPP effects. In a and **b**, the Y and N indicate LSF occurred or did not in each year at each pixel, respectively. Statistical significance was determined using two-tailed t test. The asterisk indicates a significant difference in GPP_{current} between LSF occurred and non-occurred (P < 0.01). In **c–f**, the asterisk indicates a significant difference in photosynthetic rate, starch, soluble sugar, and NSC between CK and LSF group (n = 36, P < 0.01). The box spans from the first to the third quartile, with intermediate values marked as the black line in the middle of the box. In k and l, LSF, GPP_{current}, and start of growing season in the current year (SOS_{current}) were incorporated into the SEM to explore the direct (arrows from each factor directly pointing to the SOS_{next}) or

indirect (arrows from each factor firstly directly pointing to $GPP_{current}$ then to the SOS_{next}) effects of factors on spring phenology in the next year. The calculated P values based on two-sided test and other statistics were listed in Supplementary Table 4. In $\mathbf{m}-\mathbf{p}$, the error bars indicate standard errors of the mean (n = 10698).

Fig. 4 | Projected effect of late spring frost (LSF) on start of growing season (SOS) in the next year at the end of this century (2080-2100) across different models under four climatic scenarios. a–d, Spatial distribution of mean ΔSOS when LSF occurred from the Coupled Model Intercomparison Project Phase 6 (CMIP6) models under SSP1-2.6 (a), SSP2-4.5 (b), SSP3-7.0 (c), and SSP5-8.5 (d). e–h, Spatial distribution of mean ΔSOS when LSF occurred from the sequential model (SM) with GPP effects under SSP1-2.6 (e), SSP2-4.5 (f), SSP3-7.0 (g), and SSP5-8.5 (h). i–l, Changes in ΔSOS across different models under SSP1-2.6 (i), SSP2-4.5 (j), SSP3-7.0 (k), and SSP5-8.5 (l). In a–l, ΔSOS indicates difference in SOS for each consecutive two-year period at each pixel. Y-CMIP indicates LSF occurred using CMIP6 models, Y-SM indicates LSF occurred using modified SM, and N indicates LSF did not occur using CMIP6 models. The SOS dates from 2080 to 2010 were extracted by daily gross primary productivity (GPP) dataset from CMIP6 models. In i–l, different letters indicate significant differences among different models based on Tukey's HSD test (*P* < 0.01). The detailed statistics were listed in Supplementary Table 6. The box spans from the first to the third quartile, with intermediate values marked as the black line in the middle of the box.

Fig. 5 | A schematic diagram of the interactions between late spring frost (LSF), photosynthetic carbon uptake, and spring phenology. The LSF in the current year could delay spring phenology in the next year by reducing photosynthetic carbon uptake in the current year. This LSF-induced delayed spring phenology in the next year can further reduce photosynthetic carbon uptake in the next year through feedback control.

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Methods

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Phenological data

Ground-based phenology network

The start of growing season (SOS) was obtained from the Pan European Phenology Network 541 (PEP725), USA National Phenology Network (USA-NPN), China Phenological Observation 542 Network (CPON), and Russian 'Chronicles of Nature' Network (RCNN). The PEP725 contains 543 in situ phenological observations of many temperate species across central Europe since 1868⁴⁴, 544 The USA-NPN includes plant phenological records of 400 species at 3000 sites in conterminous 545 United States⁴⁵. The CPON involves phenological data at 44 sites in China starting in 1963⁴⁶. 546 The RCNN has 506,186 phenological observations at 471 sites in Russian Federation, Ukraine, 547 Uzbekistan, Belarus and Kyrgyzstan⁴⁷. Due to the different definitions of spring phenological 548 events for each ground-based network, we screened the phenological observations of all four 549

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PhenoCam network

PhenoCam network (https://phenocam.sr.unh.edu/) is a cooperative continental-scale repository. 553 The network uses digital phenocamera imagery to track the timing of phenological transition 554 from 2000 to 2018 in North America and around the world⁴⁸. In the PhenoCam network, the 555 50^{th} , 75^{th} and 90^{th} percentile of the Green Chromatic Coordinate (G_{CC}) were calculated to extract 556

ground phenological network according to distinct criteria (Supplementary Note 1).

- the date of greenness rising and falling based on the following formula: 557

$$G_{CC} = \frac{G_{DN}}{R_{DN} + G_{DN} + B_{DN}}$$
 (1)

where R_{DN} , G_{DN} and B_{DN} denotes the average digital numbers (DN) of red, green and blue color 559 560 channels across the region of interest (ROI), respectively.

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We used the 50% threshold of 90th percentile G_{CC} as spring leaf-out⁴⁹. We removed outliers more than 2.5 times of MAD⁵⁰. We obtained three forest types: evergreen needleleaf forests (ENF), deciduous needleleaf forests (DNF), and deciduous broadleaf forests (DBF). Ultimately, we selected 529 records at 67 sites with at least five years of observations between 2000 and 2018.

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GIMMS NDVI3g phenological product

Over recent decades, satellite remote sensing data have been widely used to detect long-term 569 and large-scale plant phenological events⁵¹. We derived SOS metrics from the GIMMS NDVI3g 570 dataset (http://ecocast.arc.nasa.gov) produced by Advanced Very High Resolution Radiometer 571 (AVHRR) instruments between 1982 and 2014 at 0.08° spatial resolution and 15-day temporal 572 resolution⁵². We excluded non-vegetated areas with annual average NDVI < 0.1 to reduce bias. 573

We used Savitzky-Golay filter to smooth the time series and minimize the noise of atmospheric interference and satellite sensor before the estimation of spring phenology⁵³. We then applied a double logistic function (Eq. 2) to fit annual NDVI time series curve, and calculated the second-order derivative of the fitted curve⁵¹. The SOS was defined as the date when the rate of change in curvature reached its first local maximum value in the first half year.

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$$y(t) = a + b\left(\frac{1}{1 + e^{c(t-d)}} + \frac{1}{1 + e^{e(t-f)}}\right)$$
 (2)

where b, c, d, and f are parameters of logistic function, and a is the initial background NDVI value, a + b denotes the maximum NDVI value, t is time in days, and y(t) is the NDVI value at time t.

MODIS phenological product

The MODIS land surface phenology product (MCD12Q2 V6) was downloaded from the Land Processes Distributed Active Archive Center (LPDAAC) (https://lpdaac.usgs.gov/)⁵⁴. The MCD12Q2 product provides annual characteristics of vegetation phenology at a spatial resolution of 500 m between 2001 and 2019 on a global scale. The phenological metrics were derived from time series of the two-band enhanced vegetation index (EVI2), which is calculated from MODIS nadir BRDF adjusted surface reflectance (NBAR-EVI2). We selected the "Greenup" phase to define SOS and to represent spring leaf-out. The SOS was defined as the date when the EVI2 first crossed 15% of the segment EVI2 amplitude.

Future phenological data

Due to the absence of future NDVI and EVI data, we extracted phenological metrics using daily gross primary productivity (GPP) and leaf area index (LAI) dataset to predict future phenology shifts at the end of this century (2080-2100) under four climatic scenarios (SSP1-2.6, SSP2-4.5, SSP3-7.0 and SSP5-8.5). GPP and LAI data for the period 2080-2100 was obtained from the NorESM2-MM model at 1° spatial resolution, which is a component of the Coupled Model Intercomparison Project Phase 6 (CMIP6, https://esgf-node.llnl.gov/projects/esgf-llnl/). We first used Savitzky-Golay filter to smooth the time series and exclude abnormal daily GPP (or LAI) values before estimating spring phenology⁵³. We then applied a double logistic function (Eq. 2) to fit annual GPP (or LAI) time series and extracted SOS⁵¹, which is defined as the date when the rate of change in curvature reached its first local maximum value in the first half year.

Gross primary productivity and net primary productivity data

The annual GPP and net primary productivity (NPP) datasets between 1982 and 2018 were obtained from the global land surface satellite (GLASS) product (http://www.glass.umd.edu/index.html) at 0.05° spatial resolution⁵⁵.

611 Remote-sensing vegetation indices

We utilized two complementary remote-sensing vegetation indices, the Normalized Difference Vegetation Index (NDVI) and the Enhanced Vegetation Index (EVI), to quantify the impact of late spring frost (LSF) on tree growth. NDVI and EVI data from 2001 to 2019 were derived from Terra MODIS vegetation indices products (MOD13A1) at 500 m spatial resolution and 16-day temporal resolution.

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Tree-ring width data

We also used raw tree-ring width (TRW) and basal area increment (BAI) to examine the effect 619 of LSF on tree growth and productivity⁵⁶. TRW data were obtained from the International Tree-620 Ring Data Bank (ITRDB) database and ref.⁵⁷. In total, we remained 35,716 TRW records of 66 621 622 tree species at 1855 sites with at least 15 years of observations during 1982-2014 in the Northern 623 Hemisphere. BAI, a two-dimensional measure, provides a more accurate quantification of wood production compared to TRW58. Due to the lack of field-measured diameters in the ITRDB 624 database, we calculated diameters from the sum of annual raw TRWs in accordance with 625 previous studies^{59,60}. The BAI of each tree-ring core was calculated using Eq. 3. 626

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$$BAI = \pi (R_t^2 - R_{t-1}^2)$$
 (3)

where R denotes the radius, and t represents the year of tree ring formation.

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MODIS land cover product

The MODIS land cover product (MCD12Q1) with the International Geosphere-Biosphere Programme (IGBP) classification scheme at 500m spatial resolution was used to define the vegetation types⁶¹. The IGBP classification scheme includes 17 land cover types. We excluded non-forest vegetation types and reclassified the original forest types into the needleleaf forests, broadleaf forests, mixed forests, and woody savannas.

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Global climate zones product

The Global ecological zone (GEZ) of the Forest Resources Assessment with a spatial resolution of 0.05° was used to define the climate zones⁶². We only kept areas outside the tropics (latitudes >30 °N) that have clear seasonal phenological cycles, and classified forest biomes into subtropical forests, temperate forests, and boreal forests based on GEZ 2010 map.

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Climate data

644 Gridded daily minimum temperature during 1950-2013 in Europe were obtained from the 645 database E-OBS (http://www.ecad.eu/) with 0.25° spatial resolution to match the PEP725 phenological dataset. We also used global daily minimum and average temperatures during 1950-2020 from the CRU JRA v2.2 dataset at 0.5° spatial resolution to match other phenological datasets. Monthly Standardized Precipitation-Evapotranspiration Index (SPEI) between 1982 and 2019 were extracted from the global 0.5° gridded SPEI dataset (SPEIbase v.2.9) (http://hdl.handle.net/10261/332007). Monthly Palmer Drought Severity Index (PDSI) data from 2001 to 2019 were derived from TerraClimate dataset at 4 km spatial resolution⁶³. Future daily minimum and average temperatures during 2080-2100 were obtained from the CMIP6 model (NorESM2-MM) under four climatic scenarios (SSP1-2.6, SSP2-4.5, SSP3-7.0 and SSP5-8.5) with 1° spatial resolution. To address the difference in the spatial resolutions of various remote sensing datasets, all satellite remote sensing datasets were resampled to 0.25° spatial resolution. Then, the bilinear interpolation method was used to extract climate data for each site or pixel using the "raster" package⁶⁴ in R⁶⁵.

Spring phenology models

We used four process-based phenology models to simulate and predict spring phenology, including one-phase growing degree days model (GDD)⁶⁶, spring warming model (SW)⁶⁷, two-phase sequential model (SM)⁶⁸ and parallel model (PM)⁶⁹ (Supplementary Note 2).

Statistical analysis

We first investigated the carry-over effect of LSF on spring phenology using five ground-based phenological and PhenoCam networks, and remotely sensed data from 1950 to 2020 in the Northern Hemisphere. To clarify the frost-induced changes in spring phenology, we then examined the effect of LSF on photosynthetic productivity using remote sensing data (i.e., GPP and NPP) and frost-controlled growth chamber experiments. To test the robustness of our results, we also quantified the impact of LSF on tree growth and productivity using remote-sensing vegetation indices (i.e., NDVI and EVI), and tree-ring width (i.e., TRW and BAI) in the Northern Hemisphere. Considering that GPP and NPP data are more closely connected to the growth of leaves and canopies compared to tree-ring width records, GPP and NPP data were further utilized to examine the effect of LSF on photosynthetic productivity in the next year. Lastly, we projected the effect of LSF on spring phenology based on CMIP6 models and modified SM under four future climatic scenarios (Extended Data Fig. 5).

Effect of LSF on spring phenology

Previous studies showed that spring temperatures fall below -2° C may significantly affect the growth and survival of trees^{5,10}. Here, we defined the LSF event as the last frost event (i.e., when daily minimum temperature falls below -2° C) that occurred between the period after SOS and before 15 July⁷.

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To examine the carry-over effect of LSF on spring phenology in the next year, we calculated and compared the changes in SOS of the next year (SOS_{next}) at each site or pixel using multisource observational data when LSF occurred or did not (ground phenological network, PhenoCam network, and satellite remote sensing products) across the Northern Hemisphere. To this end, we first calculated the difference in SOS (Δ SOS) for each consecutive two-year period according to Eq.4. Then, we identified whether LSF occurred at each site (pixel) in each year, and compared the difference of Δ SOS when LSF occurred or not.

$$\Delta SOS = DOY_{t+1} - DOY_{t} \tag{4}$$

where the DOY_t is the timing of SOS in year t, and DOY_{t+1} is the timing of SOS in the next year t+1. One-way analysis of variance (ANOVA) was used to test the difference in the Δ SOS when LSF occurred or not for each phenological dataset. Using linear mixed model, we further examined the overall effect of LSF on SOS_{next} by pooling all the ground-based phenological datasets for a global test. In the model, the response variable was Δ SOS, the fixed effect was the occurrence of LSF, with random intercepts among sites and species.

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In order to eliminate the influence of low preseason temperatures on spring phenology, we categorized preseasons into either cold or warm based on whether the mean minimum temperature during winter (November to January), spring (February to April), and the combined winter and spring months (November to April) of the specific year at the particular site fell below or above their long-term averages from 1982 to 2018, respectively. We employed oneway ANOVA to assess the variation in ΔSOS when LSF occurred or did not occur during cold and warm preseasons. To test the robustness of our results, we further examined the effects of both spring temperature and LSF on timing of leaf-out using a linear mixed model. In this model, the response variable is ΔSOS, the predictors are the occurrence of LSF (a categorical variable with two levels: occurrence and non-occurrence) and the difference in spring temperature, with species and sites are set as random effects. To address the species-specific frost response, we divided the tree species into early-leafing species and late-leafing species based on the mean SOS date of all species in each ground-based phenological dataset. One-way ANOVA was used to test the difference in the Δ SOS between early- and late-leafing species, and between broadleaf and conifer species using ground-based phenological data, also among different forest types and climate zones using remote sensing and ground-based phenological datasets. We also set the minimum temperature threshold of LSF at 0°C, -2°C, and -4°C to examine and compare the carry-over effect of different levels of LSF on spring phenology in the next year in eight deciduous tree species based on the PEP725 dataset, the largest ground phenological network used in our study. To ensure the reliability of our findings, we further

incorporated the temporal dimension of frost damage into the LSF definition. Specifically, we redefined LSF to be the period from one to eight weeks after SOS and prior to July 15th, employing a range of temperature thresholds (0°C, -2°C, and -4°C) based on ground phenological networks. Using one-week smoothing window, we re-examined carry-over effect of LSF on spring phenology based on ground-based phenological data. The accumulation of thermal units prior to the LSF potentially increases the susceptibility of immature leaves to subzero temperatures. Using the smoothing window, we also calculated the accumulated Growing Degree Days (GDD) exceeding 0°C from January 1st to the occurrence of LSF for assessing the temporal change in LSF risk¹⁶.

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Estimates of the lethal temperature

Among all the ground-based phenological networks, the period of phenological records in the PEP725 spanned nearly six decades, and was distributed at more than 1000 sites. To ensure the robustness of results, the long-term and large-scale SOS records from the PEP725 network were selected to calculate the lethal temperature of 50% leaves (LT₅₀) to further clarify differences in SOS_{next} in response to LSF among tree species. The LT₅₀ is commonly determined by electrolyte leakage tests or visual damage assessment in the field or in the laboratory^{70,71}. We conducted an extensive literature search and compiled LT₅₀ in as many tree species as possible in the PEP725 dataset. Since LT50 is greatly affected by the leaf development stage, we only retained LT₅₀ measured in spring leaf-out^{5,7}. Ultimately, we obtained determined LT₅₀ values in five tree species. Also, we used a four-parameter logistic model with two fixed asymptotes and two free parameters to estimate the LT₅₀ of all eight selected temperate species in the PEP725 dataset (Eq. 5)⁷². Specifically, we used two asymptotes to limit the estimate below -1°C, which corresponds to the highest LT₅₀ of woody plants observed in temperate regions⁷³, and above -7.5°C, which is close to the lowest measured LT₅₀ obtained from literature. To ensure reliability, we extracted the median dates of long-term and multi-site SOS for each species to estimate the LT₅₀.

$$LT_{50} = \frac{a-b}{1+\exp\left[\left(c-x\right)/d\right]} + b \tag{5}$$

where a and b are fixed parameters, a = -7.5 and b = -1, c and d are free parameters (estimated c = 130, d = -10), and x is the predictor variable (median date of leaf unfolding for each species). We conducted a non-linear model fitted using the "nlme" package⁷⁴ in R⁶⁵. Correlation analysis was used to examine the relationship between the measured and calculated LT₅₀ values. Linear regression model was used to examine the relationship between the LT₅₀ predicted by the model and Δ SOS for each species in the PEP725 dataset.

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Effect of LSF on photosynthetic productivity

We first matched GLASS GPP (or NPP) data with GIMMS NDVI3g and MODIS phenological products. We then analyzed the changes in the GPP and NPP in the current year (GPP current, NPP_{current}) when LSF occurred or did not occur for each pixel using remote sensing data. Oneway ANOVA was used to test the difference in the GPP_{current} (or NPP_{current}) when LSF occurred or not in the current year. We also confirmed the LSF damage utilizing long-term and largescale remote-sensing vegetation indices (NDVI and EVI) and tree-ring data (TRW and BAI) in the Northern Hemisphere. We further used structural equation models (SEM) to examine the relationships between LSF, GPP_{current} (or NPP_{current}) and SOS_{next} using remote sensing data. We constructed a conceptual model that includes both the direct and indirect effects of LSF in the current year on SOS_{next}. In the SEM model, we hypothesized that LSF in current year is likely to directly affect the SOS_{next}, indicated by the arrows from LSF directly point to the SOS_{next}, or indirectly affect SOS_{next} by altering the photosynthetic carbon assimilation, indicated by the arrows from LSF firstly directly point to GPP_{current} (or NPP_{current}) then to the SOS_{next}. Incidentally, SOS in the current year (SOS_{current}) may directly alter carbon storage, and then affect SOS_{next}. Therefore, SOS_{current} is also included in SEM as a direct and indirect factor in the model. The piecewise SEM was fit using the "piecewiseSEM" package⁷⁵ in R⁶⁵. To eliminate the potential impact of drought on GPP_{current}, we conducted an analysis of GPP_{current} changes concerning the occurrence or absence of summer droughts (between July and September) in the current year, utilizing both 6- and 12-month SPEI data as well as PDSI data. For our analysis, a summer drought event was defined as when the mean SPEI dropped below −2 or when the mean PDSI fell below -4 during the summer months. We employed one-way ANOVA to assess the disparity in \triangle GPP_{current} between instances where LSF occurred or did not occur in the current year. Using a frost-controlled experiment with four tree species widely distributed in temperate and subtropical forests (Supplementary Note 3), we further investigated the effect of LSF on photosynthetic rate and nonstructural carbohydrates (NSC) of stem and leaves to understand the mechanisms underlying the relationship between LSF and photosynthetic productivity. We also calculated species-specific shifts in photosynthetic rate, starch, soluble sugar, and total NSC between the control (CK) and LSF groups. we employed one-way ANOVA to analyze the differences in photosynthetic rate, starch, soluble sugar, and total NSC between the CK and LSF treatment,

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Feedback effect of spring phenology on photosynthetic productivity

To evaluate whether LSF have an effect on carbon uptake in the subsequent year, one-way ANOVA were used to examine the differences in the GPP (or NPP) of the subsequent year (GPP_{next}, NPP_{next}) when LSF occurred and did not. In order to decipher the interactions between LSF, SOS and GPP, we further calculated and compared the differences in the GPP_{next} and

NPP_{next} anomaly (Δ GPP_{next}, Δ NPP_{next}) relative to the long-term average when SOS_{next} was advanced or delayed compared to SOS_{current} for each pixel. One-way ANOVA was used to test the difference in Δ GPP_{next} (or Δ NPP_{next}) when SOS_{next} was advanced and delayed.

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Projected effect of LSF on spring phenology under future climatic scenarios

We similarly used daily minimum temperature (below -2° C) to define LSF during 2080-2100 from the CMIP6 models under four climatic scenarios. We then used daily GPP data to extract future phenological metrics, and examined the effect of LSF on spring phenology based on CMIP6 models under four future climatic scenarios. Due to the superior performance of the modified SM, with solely the adjusted forcing phase, in simulating spring phenology when considering frost-induced in photosynthetic productivity, the modified SM was further used to predict the potential effects of LSF on shifts in spring phenology under future climatic scenario. One-way ANOVA followed by Tukey's HSD test was used to test the difference in mean ΔSOS when LSF occurred or not for different climatic scenarios using modified SM and original CMIP6 models. To minimize the uncertainties arising from a single phenological data, we also performed the same analysis to investigate effect of LSF on spring phenology using daily LAI data from CMIP6 models under four future climatic scenarios. We further assessed the temporal change in the carry-over effect of LSF and its risk utilizing ground phenological networks.

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All data analyses were conducted using R version 4.1.165.

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Data availability

All ground-based and remote-sensing derived phenological data are freely available on the 813 following websites: Pan European Phenology Network (PEP725), http://www.pep725.eu/; 814 815 USA National Phenology Network (USA-NPN), https://www.usanpn.org/results/data; China 816 Phenological Observation Network (CPON), http://www.cpon.ac.cn/; Russian 'Chronicles of Nature' https://doi.org/10.1038/s41597-020-0376-z; 817 Network (RCNN), PhenoCam, https://phenocam.sr.unh.edu/; 818 **GIMMS** NDVI3g phenological product, 819 https://globalecology.unh.edu/; MODIS phenological product, https://lpdaac.usgs.gov/. Gross 820 primary productivity and net primary productivity data available 821 http://www.glass.umd.edu/index.html. Gridded daily minimum and average temperatures can 822 be downloaded from E-OBS (http://www.ecad.eu/) and CRU JRA v2.2 dataset 823 (https://catalogue.ceda.ac.uk/uuid/4bdf41fc10af4caaa489b14745c665a6). Future minimum 824 and average temperatures, GPP, and leaf area index (LAI) data were from the CMIP6 models 825 (https://esgf-node.llnl.gov/projects/esgf-llnl/). Standardized Precipitation-Evapotranspiration Index (SPEI) dataset (SPEIbase v.2.9) are available on http://hdl.handle.net/10261/332007. 826 827 Palmer Drought Index (PDSI) Severity data are available at

- 828 https://climate.northwestknowledge.net/TERRACLIMATE/. MODIS vegetation indices and
- land cover data are available at https://ladsweb.modaps.eosdis.nasa.gov/. Raw tree-ring width
- 830 data were obtained from the International Tree-Ring Data Bank (ITRDB)
- 831 (https://www1.ncdc.noaa.gov/pub/data/paleo/treering/). Global ecological zone (GEZ) map is
- 832 available at https://www.fao.org/forest-resources-assessment/remote-sensing/global-
- 833 ecological-zones-gez-mapping/en/.

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Code availability

- All data analyses and modeling were performed using R version 4.1.1. The codes used for the
- 837 model simulations in this study are available at
- 838 https://doi.org/10.6084/m9.figshare.25844587.v1⁷⁶.

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