INTRODUCTION

The Ollo de Sapo (Toad’s Eye) Formation is the largest accumulation of Cambrian-Ordovician magmatic rocks in the Iberian Massif. It comprises variably metamorphosed felsic peraluminous volcanic rocks and high-level granites. The Formation crops out in the core of a Variscan anticlinorium extending for about 600km from the Cantabrian coast to central Spain in the northern part of the Central Iberian Zone (Fig. 1; Parga-Pondal et al., 1964; Díez Montes et al., 2004, and references therein). Rocks of similar composition, structure, and age are also found in other localities: i) forming a batholithic complex in the Eastern part of the Spanish Central System (Vialette
et al., 1987; Wildberg et al., 1989; Talavera et al., 2013), ii) scattered in the anatectic complexes of the Central Iberian Zone (Bea et al., 2003; Talavera et al., 2013), and iii) forming the Urra Formation at the southern boundary of the Central Iberian Zone with the Ossa Morena Zone (Fig. 1; Sola et al., 2008, and references therein).

The Ollo de Sapo gneisses have two notable characteristics. First, they are not obviously connected with any major tectonic or metamorphic event (Gutiérrez Marco et al., 2002). Second, they have an unusually high proportion of inherited zircons: in the 20 massifs we have studied so far, no less than 70-80%, and in some samples nearer 100%, of zircon grains contain pre-magmatic cores (Bea et al., 2006b; Montero et al., 2007; Montero et al., 2009; Talavera et al., 2013). This high degree of inheritance seems to be a common feature of the Cambrian-Ordovician rocks of Western Europe (e.g. Laumonier et al., 2004; Teipel et al., 2004; Helbing and Tiepolo, 2005) that undoubtedly reflects a fundamental fact of their formation.

The whole-rock geochemistry of the Ollo de Sapo gneisses is felsic, markedly peraluminous and alkaline-calcic —in the Frost et al. (2001) rather than in the

![Image of paleogeographic zones of Iberia and the Ollo de Sapo Anticlinorium](image)

**FIGURE 1.** A) Distribution of the paleogeographic zones of Iberia; the gray area represents the Iberian Massif, and the black areas represent exposures of Cambrian-Ordovician rocks. CIZ = Central Iberian Zone; CXG-D = Schist-Graywacke Complex Domain; CZ = Cantabric Zone; GTOMZ = Galicia Trás os Montes Zone; OMZ = Ossa Morena Zone; OS-D = Ollo de Sapo Domain; SPZ = South Portuguese Zone; and WALZ = Western-Asturian Leonian Zone. B) The Ollo de Sapo Anticlinorium. Stars represent the three samples studied for zircon oxygen and hafnium isotopes. Their composition is given in Table 1.
Peacock (1931) sense— with trace element features that bear similarities with those shown by arc magmas (see below). Based on this, and despite the lack of petrological or regional evidence, some authors have proposed that the Ollo de Sapo magmas originated in a supra-subduction setting (e.g. Gebauer et al., 1993; Valverde-Vaquero and Dunning, 2000; Navidad and Bea, 2004; Del Greco et al., 2016). Other authors, however, disagreed and favored a rifting environment. Bea et al. (2007) for example, based on the abnormally high zircon inheritance, proposed that the Ollo de Sapo magmas resulted from extremely fast melting rates in the middle-lower crust followed by fast melt transportation to the upper crust in an extensional regime. Díez Montes et al. (2010) reached analogous conclusions based on field and structural data.

It is evident that the controversies about the petrogenesis and geodynamic significance of the Ollo de Sapo gneisses mainly arise from the lack of understanding of the nature of the magmas involved. This paper attempts to identify the magmatic sources which, we believe, is the first step in that direction. To this end we combined the inferences derived from exploitation of a large dataset of whole-rock element and Sr and Nd isotopes obtained by the authors over the years, with new oxygen and hafnium isotope data on the abnormally high zircon inheritance, proposed that the Ollo de Sapo magmas resulted from extremely fast melting rates in the middle-lower crust followed by fast melt transportation to the upper crust in an extensional regime. Díez Montes et al. (2010) reached analogous conclusions based on field and structural data.

GEOLOGICAL SETTING AND PETROGRAPHY

The Ollo de Sapo Formation crops out in the core of a Variscan anticlinorium extending for about 600km from the west Cantabrian Coast to central Spain (Fig. 1; see an overview in Díez Montes et al., 2004). It consists of metavolcanic and metagranitic rocks that are overlain by a clastic succession topped by the Armorican Quartzite, and underlain by another succession of Early to Middle Cambrian sequences underlying the Ollo de Sapo (González Lodeiro, 1981; Iglesias Ponce de Leon and Ribeiro, 1981; Díez Montes et al., 2010). The metagranites are composed of large K-feldspar megacrysts embedded in a coarse-grained matrix of quartz, plagioclase, K-feldspar, biotite, muscovite that may be accompanied by rare tourmaline, cordierite and garnet. The accessories assemblage consists of apatite, zircon, monazite, rare xenotime and Fe-Ti oxides.

SAMPLES AND METHODS

For this work we have used 32 samples collected all throughout the Ollo de Sapo Formation. All samples were analyzed for major and trace elements, and 28 of them were also analyzed for Sr and Nd isotopes. To avoid inter-laboratory effects all analyses, except Hf isotopes, have been undertaken in the authors’ laboratories at the Centro de Instrumentación Científica of the University of Granada (CIC-UGR). Zircons from 11 samples have been dated previously using single-grain evaporation and SIMS techniques (Montero et al., 2007; 2009; Talavera et al., 2013). For this work we have separated zircons from three new samples representing the most common facies of the Ollo de Sapo rocks: one metagranite (R19) and two metavolcanic rocks, one coarse-grained (R32) and the other fine-grained (R31; Fig. 1). Using the SHRIMP, we dated them and determined their oxygen isotope composition. Hf isotopes were determined using a MC-LA-ICPMS system.

Whole-rock major-element and Zr determinations were done by XRF after fusion with lithium tetraborate. Typical precision was better than ±1.5% for an analyte concentration of 10 wt.%. Trace-element determinations, except Zr, were done by ICP-MS using Rh as an internal standard. Samples were dissolved in HNO₃ + HF at a pressure of 180p.s.i. in a microwave digestor, dried and redissolved in 4% HNO₃. Precision, estimated from the analyses of 10 replicates of one sample, was better
than ±2% and ±5% for analyte concentrations of 50 and 5ppm respectively.

Samples for Sr and Nd isotope studies were digested in a clean room using ultra-clean reagents and analyzed by TIMS in a Finnigan Mat 262 spectrometer after chromatographic separation with ion-exchange resins. Normalization values were 86Sr/88Sr=0.1194 and 146Nd/144Nd=0.7219. Blanks were 0.6 and 0.09ng for Sr and Nd, respectively. The external precision (2σ), estimated from the results of the last 10 replicates of the standard WS-E (Govindaraju et al., 1994), which is routinely analyzed each 10 unknown samples, was better than 0.003% for 87Sr/86Sr, and 0.0015% for 146Nd/144Nd. 87Rb/86Sr and 147Sm/144Nd were directly determined by ICP-MS (Montero and Bea, 1998) with a precision, estimated by analyzing 10 replicates of the standard WS-E, better than 1.2% and 0.9% (2σ), respectively.

Zircon was separated using panning, first in water and then in ethanol. After eliminating the magnetic fraction from the concentrates with a neodymium magnet, zircons were handpicked under a binocular microscope and put in a SHRIMP megamount (Ickert et al., 2008). Once mounted and polished, zircon grains were studied by optical and cathodoluminescent imaging, coated with a 10nm thick gold layer, and analyzed for U-Th-Pb and Oxygen isotopes using a SHRIMP IIe/mc ion microprobe at the IBERSIMS laboratory of the Centro de Instrumentación Científica del University of Granada (CIC-UGR). The SHRIMP U-Th-Pb analytical method roughly followed the standard method described in detail in www.ugr.es/ibersims. Uranium concentration was calibrated using the SL13 reference zircon (417Ma; Black et al., 2002) which was measured every 4 unknowns. When required, common lead was corrected from the standard WS-E, better than 1.2% and 0.9% (2σ), respectively.

Once analyzed for U-Th-Pb, the megamounts were burned for about 5 minutes before measurements. During this time the secondary beam and the e-gun were fully optimized to maximize the 18O signal. Measurements were done in two sets of 10 scans each. The scans were of 10 seconds each so that the real data collection time was 200 seconds per spot. The EISIE (electron induced secondary ion emission) background was recorded during 10s before and after each set, and subtracted from the 18O and 16O counts. As a standard we used the TEMORA-II zircon measured every four unknowns and cross-checked again the 91500 zircon every 20 unknowns. The reproducibility of the standards was excellent: δ18O=8.17±0.34 (2σ) for the TEMORA-II and δ18O=9.98±0.26 (2σ) for the 91500, respectively. Data reduction was done with the POXY program developed by P. Lanc and P. Holden at the Australian National University.

Hf isotopes were done at the Geochronology and Isotope Geochemistry-SGIker facility of the University of the Basque Country using a Thermo-Fisher Scientific Neptune MC-ICP-MS coupled to the New Wave Research UP-213 laser system with a SuperCell laser cell. Data were collected in static mode during 50s of ablation with a spot size of 40μm. Masses 171, 173, and 175 were simultaneously monitored during each step to correct the isobaric interferences of Lu and Yb isotopes on mass 176. 176Yb and 176Lu were calculated assuming a 176Yb/172Yb of 0.796179 and a 176Lu/172Lu of 0.02655 (Chu et al., 2002) and taking into account the instrumental mass fractionation of each individual analysis. To correct Instrumental Mass Fractionation (IMF) Yb isotope ratios were normalized to 173Yb/171Yb=1.132685 and Hf isotope ratios to 176Hf/177Hf=0.7325 (ibid.) using an exponential law. The Lu IMF was assumed to follow that of Yb. Data reduction was performed using the Iolite 2.5 software package for deconvolution of time resolved data (Patton et al., 2011).

**GEOCHEMICAL AND ISOTOPIC SIGNATURES**

The major element compositions of the Ollo de Sapo gneisses (Table 1) are highly variable, with significant differences in the proportions of normative feldspars and quartz. The proportions of normative feldspars correspond to granites sensu stricto (Fig. 4A) and the proportions of normative albite, orthoclase and quartz are identical to ordinary corundum-normative (peraluminous) granites (Fig. 4B). We have found no meaningful compositional differences between metavolcanic and metaprotolith facies.
The major elements are notably decoupled; TiO$_2$, Al$_2$O$_3$, MgO, and FeO$_{tot}$ are well correlated with SiO$_2$ but CaO, Na$_2$O, K$_2$O and P$_2$O$_5$ are not. In contrast to differentiated calc-alkaline granitoids, in these rocks the saturation in alumina decreases with increasing silica so that the less silicic samples are notably more aluminous than the most silicic ones (Fig. 3). Noticeably, these lack such geochemical features as low K/Rb, low Zr/Hf, etc., that characterize highly differentiated granite rocks (e.g. Taylor et al., 1956; Bea et al., 2006a).
Isotopes in zircons from the Ollo de Sapo Formation

Chondrite-normalized REE patterns are parallel for LREE and MREE but diverge considerably for the heaviest REE (Fig. 5). La/N is typically between 100 and 200, decreasing to Sm/N between 30 and 60, with a small negative Eu anomaly between Eu/Eu*=0.53-0.65 and Lu/N between 2 and 20. La/N/Lu/N is between 7 and 60.

NMORB-normalized spidergrams (Fig. 6) show a marked enrichment with increasing incompatibility, especially for the mica-hosted elements Rb and Cs, with notable troughs in Nb, Ti and Sr, and peaks in K and Pb. They however, lack the enrichment in Sr and moderate depletion in Zr characteristic of true arc-magmas.

TABLE 1. (Cont.)

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<td>1.86</td>
<td>0.59</td>
<td>1.66</td>
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DOI: 10.1344/GeologicaActa2017.15.4.1
The Sr and Nd isotope compositions of the Ollo de Sapo (Table 2) are clearly crustal: 87Sr/86Sr485Ma is within the 0.7094 to 0.7113 interval (at 95% confidence) and clusters at 0.7103; the εNd485Ma is within the -4.76 to -4.23 interval (at 95% confidence) and clusters around -4.5 (Fig. 7A). It is worth mentioning that in a 87Rb/86Sr vs. 87Sr/86Sr plot most points scatter along the 485Ma reference line (Fig. 7C) so that, excluding the sample with the highest 87Rb/86Sr, the remaining 27 samples fit an errorchron at 493±27Ma which nearly matches the average U-Pb zircon age (see next section). The goodness-of-fit of the errorchron despite the large sampled area and variable metamorphic grade indicates that the Ollo de Sapo rocks were not heavily disturbed during the Variscan metamorphism.

The Nd TDM model age distribution is asymmetric; it peaks a 1.5Ga and shows a large tail towards higher values with a minor peak at 1.8Ga (Fig. 7B). The 1.5Ga value is very common in Iberia and northern Africa (e.g. Bea et al., 2010) but finds no correspondence with the U-Pb zircon ages. Bea et al. (2011) interpreted this value as a “mixed” Nd model age that is characteristic of most Pan-African granitoids generated throughout the Saharan metacraton.

ZIRCON DATA

U-Pb ages

As mentioned in previous sections, the Ollo de Sapo zircons are distinctive because ~70-80%, in some cases up to 100%, contain pre-magmatic zircons (Bea et al., 2006b; Montero et al., 2007; Montero et al., 2009; Talavera et al., 2013). The data obtained by these authors in the zircon rims indicate that the magmatism of the Ollo de Sapo Formation started at about 495Ma, reached its maximum at 485Ma, and ended at about 470Ma (Fig. 8). The pre-magmatic cores show a wide range of inherited dates from Ediacaran to Mesoarchean but most of them are Neoproterozoic, mainly Ediacaran, with ages between 590Ma and 620Ma.

The three new samples studied here fit this pattern. The orthogneiss R19 (Fig. 8) yielded an age of 483±3Ma (2s) in the rims (22 data), which we consider the crystallization age. Most cores cluster around 600Ma but there are minor populations at ca. 680Ma, 890Ma, 970Ma and 2525Ma. The coarse-grained metavolcanic rock R32 (Fig. 8) yielded a rim age of 482±2Ma (51 data). The Ediacaran cores define a broad peak which likely indicates provenance from several sources with slightly different ages. A few cores yielded older-than-Ediacaran ages, up to Mesoarchean. Sample R31, a metavolcanic rock from the top of the sequence has the youngest rims with an age of 478±4Ma (Fig. 8). The Ediacaran cores peak at 615Ma, but the peak asymmetry suggests the contribution of another population at 590Ma-600Ma.

Oxygen isotopes

Zircons from the three low-grade samples studied for oxygen isotopes show the same picture: the Cambrian-Ordovician rims are more enriched in the heaviest oxygen than the Ediacaran and older cores (Fig. 9). The rims yielded unimodal, nearly gaussian distributions that cluster around δ18O=10, values which are typical of S-type magmas formed from melting of altered crust. The cores, however, are considerably impoverished in the heaviest oxygen, with δ18O clustering around 6.5. The oxygen isotope distribution in the cores is unimodal for samples R31 and R32 but shows a minor δ18O mode around 8.5 in the orthogneiss R19. The 6.5
mode is about one per mil higher than typical mantle values but of the same order as subduction-related granites (Valley, 2003); they indicate that the cores of the Ollo de Sapo zircons crystallized from magmas that were either derived from primitive sources slightly altered with low temperature fluids or were contaminated with metasediments.

**Hafnium isotopes**

The Hf isotope compositions of rims and cores overlap, but the former are considerably more homogeneous (Fig. 10). In the rims, the $\varepsilon_{Hf}^{485Ma}$ spans from -6.4 to 2.6 with a mean of -2.1. The distribution in the two metavolcanic rocks is sharply unimodal (Fig. 11) albeit peaking at different values (-0.4 the fine-grained, and -3.4 the coarse-grained metavolcanites). The metagranite R19, however, is bimodal, with a lower mode at $\varepsilon_{Hf}^{485Ma}$=-0.2 and a higher mode at $\varepsilon_{Hf}^{485Ma}$=1.6. On the other hand, the distribution of $\varepsilon_{Hf}^{600Ma}$ in the Ediacaran cores is polymodal, notably more irregular than in the rims.

The Hf model ages (calculated assuming a depleted mantle that changed from chondritic at $\varepsilon_{Hf}^{4.56G}=0$ to $\varepsilon_{Hf}^{0G}=+16$, and a source with $^{177}Lu^{176}Hf=0.007$, see Bea et al., Submitted) range between 1.1 and 1.5Ga with a mean of 1.3Ga but, again, the average values for each sample are markedly different, mimicking somehow the distribution of $\varepsilon_{Hf}^{485Ma}$.

The cores are more heterogeneous than rims. Twenty-nine Pan-African cores yielded $\varepsilon_{Hf}^{485Ma}$ from -9.4 to 9.8 with a mean of -0.02, and Hf model ages from 0.8 to 2.4Ga with mean at 1.34Ga. Remarkably, the average of rims and cores are nearly identical despite the largest variation of the latter. Lastly, eleven older-than-Panafrican cores yielded $\varepsilon_{Hf}^{T}$ from -22 to 4.6 and the Hf $T_{DM}$ from 1.7 to 3.2Ga so indicating the contribution of Paleo-Proterozoic to Archean components to the source.
DISCUSSION

As mentioned in the Introduction, some authors have proposed a supra-subduction setting for the Ollo de Sapo and compared these rocks, either implicitly or explicitly, with I-type calc-alkaline subduction-related granite-rhyolite series. This hypothesis is based solely on the general enrichment in the most incompatible trace elements and depletion in Nb and Ti. It, however, is inconsistent with major element and Sr and Nd isotopes.

The lack of correlation between CaO, Na$_2$O, K$_2$O and P$_2$O$_5$ with SiO$_2$ rules out that the Ollo de Sapo rocks might represent variably differentiated members of magma batches with a similar initial composition. Of particular relevance are the elevated peraluminosity of the less silicic samples and the negative correlation between ASI and SiO$_2$. These features, which are never seen in arc-related calc-alkaline granite-rhyolite series (Fig. 3), suggest different degrees of restite unmixing in anatectic magmas derived from peraluminous sources. Moreover, Sr and Nd isotopes reveal that the magmatic sources contained no detectable juvenile components and so must have resided in the crust long before the Cambrian-Ordovician melting event.

It might be argued that the original chemical and isotopic features of the Ollo de Sapo rocks were profoundly disturbed by the Variscan metamorphism, or that the spatial heterogeneity of the magmatic sources masked the inter-element correlations characteristic of felsic arc-magmas. However, none of these ideas find support from Sr isotopes as shown by the excellent match between the twenty-seven samples Rb-Sr errorchron age (493±27Ma; Fig. 7) and the SHRIMP zircon U-Pb ages (Fig. 8) mentioned before.

Geochemical evidence, therefore, is inconsistent with a supra-subduction setting but rather points to an intra-continental origin. The magmatic arc-like trace element features can be attributed to inheritance from the magma source which consisted of rocks with abundant Ediacaran zircons. The question is now to understand whether these sources consisted of Ediacaran igneous rocks with some older zircon components, or younger immature sediments derived from them, a question first raised by Montero et al. (2007).

The answer comes from the zircon oxygen isotope compositions. All rims without exception have an elevated $\delta^{18}$O of around 10. This implies a source with whole-rock $\delta^{18}$O close to 12 which is a typical "sedimentary" value absolutely uncharacteristic of fresh igneous rocks from magmatic arcs (Valley, 2003, and references therein). The oxygen isotope composition of the 600-615Ma cores, in contrast, is around $\delta^{18}$O≈6, thus compatible with Ediacaran magmas being arc-derived. Accordingly, the contrasting oxygen isotope composition of Cambrian-Ordovician rims and Ediacaran cores indicates that the Ollo de Sapo magmatic sources likely consisted of detrital sediments mainly derived from Ediacaran igneous rock rather than the Ediacaran igneous rocks themselves. Considering zircon ages and composition (Talavera et al., 2012) the most likely candidates are the graywackes of the Schist-Graywacke-Complex, the dominant Central Iberian meta-sedimentary formation (e.g. Rodríguez Alonso et al., 2004, and references therein).

A sedimentary source explains the large variability of U-Pb ages and Hf isotopes in the Ediacaran cores (Figs. 8 and 10). If we accept that the Ollo de Sapo magmas were swiftly produced and emplaced high in the crust, as suggested by the abnormally high zircon inheritance (Bea...
et al., 2007), we must also accept that the so produced magmas could not have mixed over large volumes. Had the source been an igneous rock, the Hf isotope composition of their zircons would have been nearly uniform. If, on the other hand, the source consisted of sediments in which zircons were randomly accumulated as detrital particles coming from isotopically different Ediacaran rocks, the magmas generated from them would have had isotopically heterogeneous pre-magmatic zircons. However, as a result of homogenization of the melt phase, they would have crystallized nearly isotopically homogenous rims with a Hf composition close to the core average. This can explain the distribution of εHf in rims and cores of zircons from the Ollo de Sapo rocks depicted in Figure 11.

Lastly, a sedimentary source may also explain why the Hf TDM of the zircons (1.3Ga) is significantly younger than the whole-rock Nd TDM (1.5Ga). This effect suggest that the source, in addition to detrital materials coming from Ediacaran igneous rocks, also contained a small fraction of older materials with higher REE/Zr the remnants of which are the Mesoproterozoic to Archean zircon cores of the Ollo de Sapo. Nevertheless, this suggestion must be considered with caution because calculating Hf TDM model ages of zircons requires an assumption regarding average crust composition, and this introduces a difficult-to-evaluate uncertainty factor (see Vervoort and Kemp, 2016, and references therein). In the present case we tuned the calculations for the average composition of the pre-Variscan Iberian crust (Lu/Hf=0.045; submitted) which is remarkably uniform, but even using an average crust composition (Lu/Hf=0.09; Rudnick and Fountain, 1995) the Hf TDM yields 1.35Ga, which is still younger than the Nd TDM.

Accordingly, we infer that the Ollo de Sapo magmas were produced by melting of (meta) sedimentary sources that mostly, but not solely, derived from 600-615Ma polygenic igneous rocks. Currently, no examples of such materials are known in Central Iberia. The only Ediacaran rocks precisely dated are the 580±3Ma Aljucén gabbronoritites (Talavera et al., 2008) and, close to the Cambrian boundary, the 543±6 Almohalla granodiorites (Bea et al., 2003) so that the main component of the Ollo de Sapo magmatic sources remains hidden to direct observation. Nonetheless, considering the

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TABLE 2. Isotope composition of the Ollo de Sapo gneisses. See text.
FIGURE 7. A) $^{87}\text{Sr}/^{86}\text{Sr}_{485\text{Ma}}$ vs. $\varepsilon\text{Nd}_{485\text{Ma}}$ for the Ollo de Sapo gneisses. No juvenile components are detected. B) Frequency plot of the Nd TDM model age (DePaolo, 1981). The distribution peak at 1.5Ga, is asymmetric toward the highest values and has a minor mode at 1.8Ga. C) Rb-Sr errorchron built with 27 Ollo de Sapo gneisses. The age is 493±27Ma, i.e. within the error range of the zircon U-Pb crystallization age. See text for discussion.
FIGURE 8. Frequency diagram of the U-Pb ages in the three studied samples (raw data in ELECTRONIC APPENDIX I (available at www.geologica-acta.com), Tables I, II, and III). Ages younger than 500Ma correspond to zircon rims; ages older than 500Ma are always pre-magmatic cores. The crystallization ages, 483±3Ma for the metagranite, 482±2Ma for the lower metavolcanite and 478±4Ma for the upper metavolcanites are close to the age of maximum magma production, ca. 485Ma (Montero et al., 2009, and references therein). Note how the most abundant pre-magmatic cores are Ediacaran although minor Cryogenian and Tonian populations can also be recognized.
FIGURE 9. $\delta^{18}$O distribution in zircon rims and cores of the studied samples (raw data in ELECTRONIC APPENDIX I, Table IV). The rims are enriched in $^{18}$O thus indicating derivation from metasediments. The composition of the cores, however, is consistent with magmas rich in juvenile components and little participation of sedimentary materials. See text for discussion.
The chemical and isotopic composition of the Ollo de Sapo rocks and the ages of the inherited zircons, we propose that this component would have not be very different to the Pan-African granitoids of North-East Africa from which Central Iberia detached during the Cambrian-Ordovician (see Bea et al., 2010; 2011, and references therein).

SUMMARY AND CONCLUSIONS

The Ollo de Sapo gneisses are Cambrian-Ordovician S-type granitoids that contain an abnormally elevated proportion of zircons with premagmatic cores, most of which are Ediacaran.

The gneiss Sr and Nd isotope compositions (\(87^{\text{Sr}}/86^{\text{Sr}}\) at 485 Ma = 0.7104 ± 0.0008, \(\varepsilon_{\text{Nd}}\) at 485 Ma = -4.50 ± 0.26) are decidedly crustal, indicating their sources contained little, if any juvenile material.

The trace element compositions bear some similarities to high-K calc-alkaline rocks from subduction settings such as the enrichment in the most incompatible elements, depletion in Nb and Ti, and peaks in K and Pb but lack the enrichment in Sr and moderate depletion in Zr characteristic of true arc-magmas.

The Ollo de Sapo zircons are markedly bimodal with respect to oxygen isotopes. Whereas the Cambrian-Ordovician rims yielded unimodal distributions that cluster around \(\delta^{18}\text{O} = 10\), typical of S-type magmas formed from melting of altered crust, the Ediacaran cores cluster around \(\delta^{18}\text{O} = 6.5\), typical of arc-magmas. Rims and cores have the same average Hf isotope composition but the former are much more homogeneous.

The elevated \(\delta^{18}\text{O}\) of the Cambrian-Ordovician zircon rims reveals that the Ollo de Sapo gneisses are in fact S-type granites and rhyolites that resulted from anatexis of younger-than-600Ma immature sediments. These sediments mostly, but not solely, derived from different Ediacaran igneous rocks with a wide range of Hf isotope composition. Despite the lack of direct field evidence, zircon data suggest that the metasedimentary sources of the Ollo de Sapo were the graywackes that form part of the Schist-Graywacke Complex.

The abundance of inherited zircons indicates that melting and emplacement occurred with little fractionation and discharge of suspended solids. This indicates that the...
arc-like trace element features were inherited from the sedimentary sources formed from erosion of Ediacaran arc-like igneous rocks.

Our data preclude that the Ollo de Sapo rocks might have resulted from arc-related magmatism and lend support to the idea of generation in an extensional regime as proposed by Bea et al. (2007) and Díez Montes et al. (2010).

ACKNOWLEDGMENTS

This paper has been financed by the Spanish Grant CGL2013-40785-P, and the Andalusian Grant P12.RNM.2163. This is the IBERSIMS publication nº 39.

REFERENCES


Manuscript received December 2016; revision accepted September 2017; published Online October 2017.
ELECTRONIC APPENDIX I

Supplementary material available at:

- Table I: revistes.ub.edu/public/journals/13/article1/TableI.xlsx
- Table II: revistes.ub.edu/public/journals/13/article1/TableII.xlsx
- Table III: revistes.ub.edu/public/journals/13/article1/TableIII.xlsx
- Table IV: revistes.ub.edu/public/journals/13/article1/TableIV.xlsx
- Table V: revistes.ub.edu/public/journals/13/article1/TableV.xlsx