

Pyrrhotite nano-inclusions in apatite from Les Guilleries lamprophyres (NE Iberian Peninsula)

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1. INTRODUCTION

- Les Guilleries lamprophyre dykes (LD) (spessartites) represent the last pulses of the least modified magmas at the end of Variscan magmatism prior to Triassic extension.
- **Apatite** is present in accessory proportions but is ubiquitous in the Les Guilleries LD.
- ❖ Apatite can form by a variety of processes, from **primary** igneous crystallization to secondary alteration products, and its examination, such as grain sizes, textures, and composition provide useful information about petrogenetic processes.

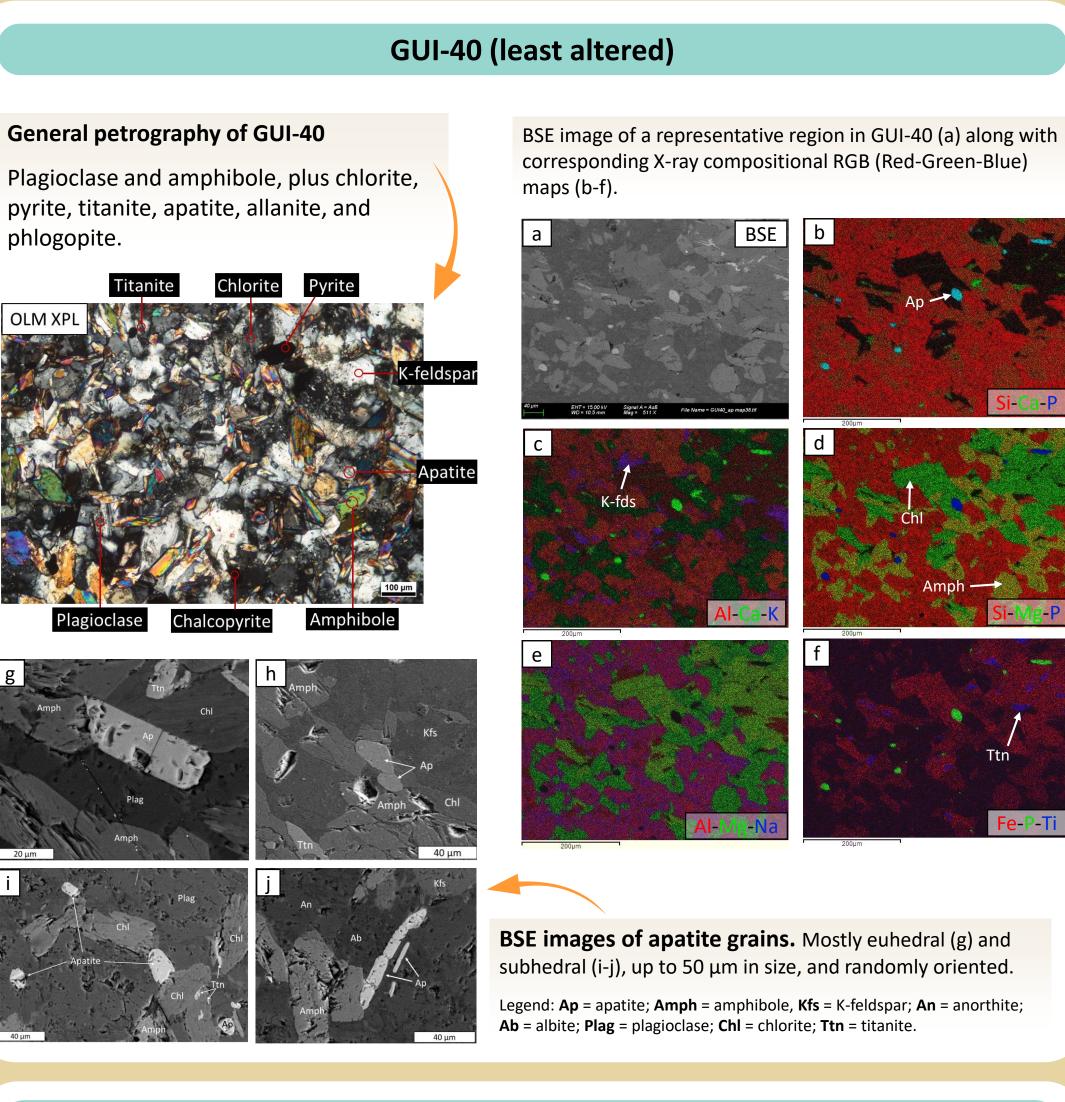
2. GEOLOGIC SETTING Les Guilleries Location of LGL dykes Late Variscan magmatism Les Guilleries lower series (Osor block) Variscan igneous rocks a) Location of les Guilleries Massif and Les Guilleries Carboniferous flysch deposits Post-Paleozoic rocks LD within the Catalan Coastal Ranges. b) Geological Pre-Variscan magmatism Ordovician Orthogneisses Les Guilleries Lamprophyres map of the study area and location of les Guilleries LD. Ediacaran, Cambrian and Ordovician c) Photograph of lamprophyre dykes in a host marble. Marble levels Source: Modified after Mellado et al. (2021).

3. ANALYTICAL TECHNIQUES

We have examined two polished thin sections of the lamprophyre dykes emplaced in marble (Fig. 1), numbered GUI-40 and GUI-41, using:

- Optical Light Microscopy (OLM) for general petrography.
- Scanning Electron Microscopy (SEM) for BSE imaging and EDS Xray analysis of interest sites and apatite grains.
- Electron probe micro-analysis (EPMA) for chemical analyses of apatite grains.
- Transmission electron microscopy (TEM) for BF TEM and DF STEM imaging, SAED patterns, and HR TEM for examination of apatite texture and composition at (sub)micron scales.

4. RESULTS AND DISCUSSION

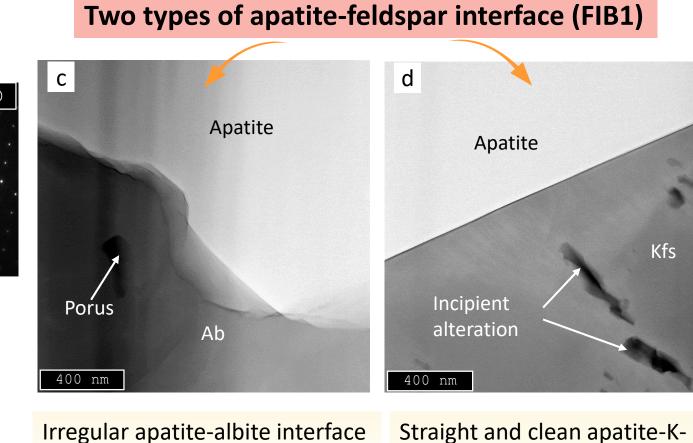


TEM observations (Apatite in GUI-40) DF STEM Apatite [001] a) BSE image showing the location of the FIB cut. b) HAADF mosaic of the FIB section. Apatite Polysynthetic is oriented down the c axis ([001] **twinning** zone axis) and associated with Ab and Kfs. Inclusions follow the crystallographic

orientation of apatite

Apatite

[001] zone axis



- Dissolution (alteration).

Apatite

High-resolution TEM image

of the low-Z phase = voids

DF STEM line profile of a bimodal

H₂S gas?

Straight and clean apatite-Kfeldspar interface. Growth under equilibrium conditions.

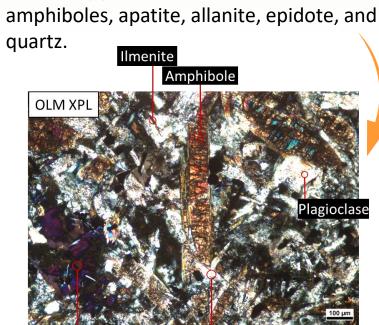
Nano-inclusions within host apatite

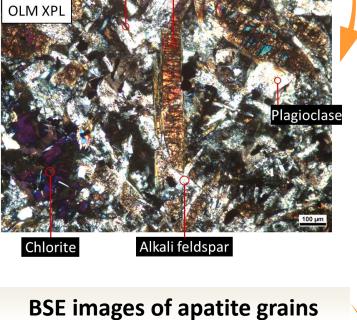
- ❖Inclusions only at the apatite core (e,l).
- Inclusions are abundant and similarly oriented (f,g,m), following the **negative crystal** of apatite.
- Inclusions consist of pyrrhotite (h), pyrrhotite and void (i), and/or void space (j).

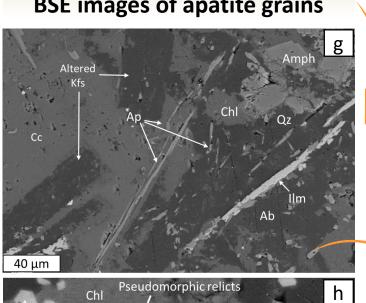
Line profile 1

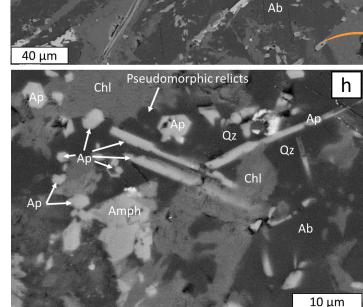
GUI-41 (more altered)

General petrography of GUI-41 Altered amphiboles (~300 μm sized) and plagioclase, plus chlorite and calcite (0.5 mm sized). Albite dominates the groundmass. Other phases include titanite, ilmenite, pyrite, chromite, smaller amphiboles, apatite, allanite, epidote, and

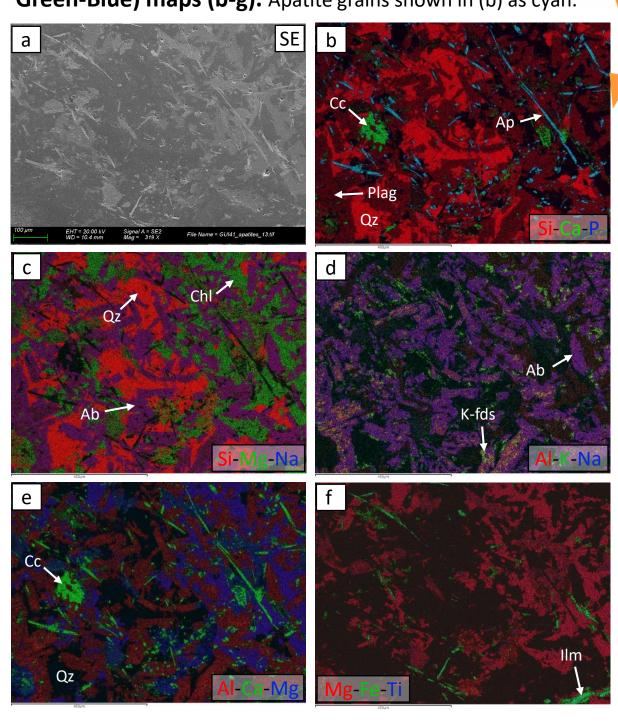








BSE image of a representative region in GUI-41 (a) along with corresponding X-ray compositional RGB (Red-**Green-Blue**) maps (b-g). Apatite grains shown in (b) as cyan.

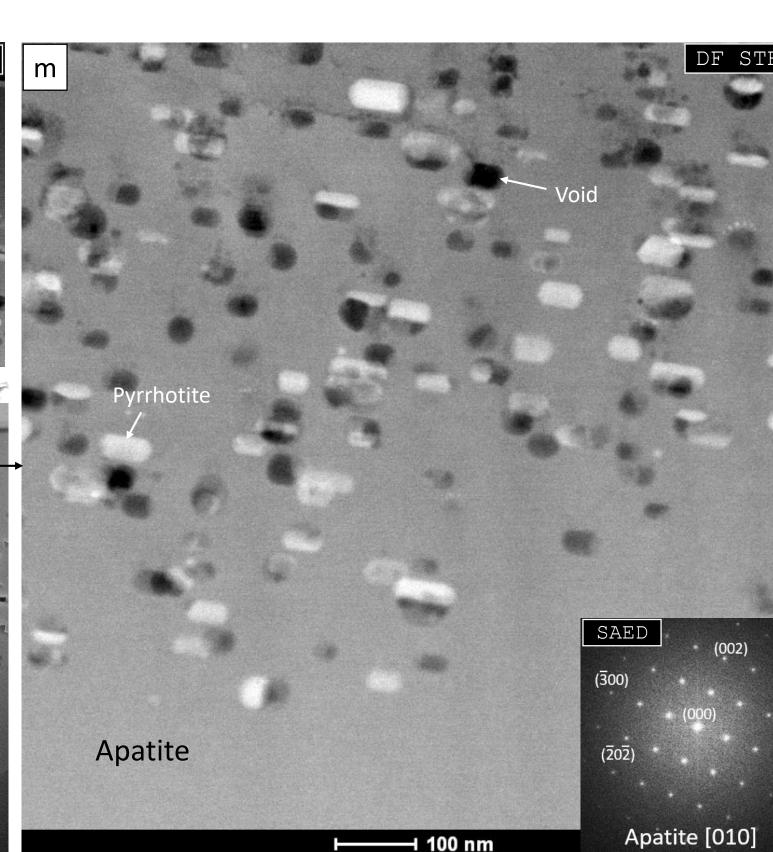


- Apatite grains (g,h):
- Range from a few-to-ten μm in width and up to 150 μm in length (highly acicular).
- Randomly oriented, heterogeneously distributed, and appear to cut all pre-existing phases, including chlorite and calcite.
- Abbreviations: **Ap** = apatite; **Amph** = amphibole; **Plag** = Ca,Na-plagioclase; **Ab** = albite; **K-fds** = K-feldspar; **Chl** = chlorite; **Ilm** = ilmenite; **Qz** = quartz; **Cc** = calcite.

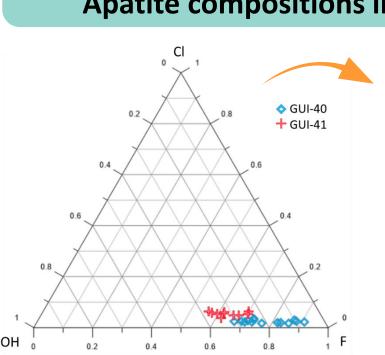
inclusion showing the drop of elements at the low-Z phase and the peaks in S and Fe on the high-Z phase. k) SE image showing the Albite location of the FIB cut. I) SEM in-lens Silicate-sulfide melt immiscibility image of the FIB section during FIB preparation. Apatite is oriented subparallel to the c axis. Pt strip Prismatic/tubular FIB2 Nano-inclusions Nano-porosity Apatite

Nano-inclusions

Porosity —



Apatite compositions in GUI-40 and GUI-41 (EPMA)



Ternary diagram for CI-F-OH Primary apatites (GUI-40) are F-richer (and Cl-poorer) than

 Hydroxyl content is higher in secondary apatites

secondary apatites (GUI-41).

Little to no overlap between the two datasets.

Primary versus secondary apatites Pyrrhotite nano-inclusions Primary Secondary FeO (wt%)

1 μm

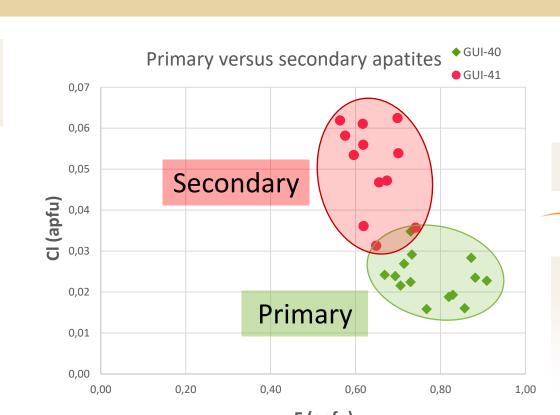
Feldspar

Individual EMPA analysis of apatite

grains in GUI-40 and GUI-41

FeO versus SO₃ (wt%) plot

GUI-40 apatite: much larger variability in S and Fe contents in GUI-40 apatites compared to GUI-41 apatites. Attributed to the presence of **pyrrhotite** inclusions observed by TEM.



F versus CI (apfu) plot

Secondary apatites contain slightly more CI than primary apatites, although the CI content is low in all of them.

5. CONCLUSIONS

- Apatite grains from GUI-40 are primary, formed by igneous crystallization from the parental lamprophyre magma, whereas apatite grains from GUI-41 are secondary products of hydrothermal alteration.
- The parental magma contained some Fe and S. Sulfur was not incorporated into the apatite structure to form ellestadite domains (as seen in Ferraris et al., 2005) due to the presence of Fe, because Fe is incompatible in apatite.
- ❖TEM work reveals nano-inclusions of pyrrhotite and void space within host apatite, following the negative crystal of apatite. The parental magma experienced melt immiscibility and apatite trapped droplets of a sulfide melt.
- ❖ From this study, we conclude that lamprophyres formed under reducing conditions.
- This work demonstrates that (sub)micrometer studies of apatite using TEM can give valuable insights into parental magmas and processes that are overlooked at the SEM scales (i.e., Martínez et al., 2023a,b).

6. ACKNOWLEDGMENTS

Nano-inclusions <

Clean interface

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7. REFERENCES

Ferraris, C., White, T. J., Plévert, J., and Wegner, R. (2005) Nanometric modulation in apatite. Physics and chemistry of minerals 32, 485-492

Martínez, M., Shearer, C.K., and Brearley, A.J. (2023a) Nanostructural domains in martian apatites that record primary subsolidus exsolution of halogens: Insights into nakhlite petrogenesis. American Mineralogist. Preprint. 10.2138/am-2022-

Martínez, M., Shearer, C. K., and Brearley, A. J. (2023b) Ferro-chloro-winchite in Northwest Africa (NWA) 998 apatite-hosted melt inclusion: New insights into the nakhlite parent melt. Geochimica et Cosmochimica Acta 344, 122-133.

Mellado, E., Corbella i Cordomí, M., Navarro Ciurana, D., and Kylander, A. (2021). The enriched Variscan lithosphere of NE Iberia: data from postcollisional Permian calc-alkaline lamprophyre dykes of Les Guilleries. Geologica Acta 19, 1-23.

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